# The Slice Method

## Atmospheric Sciences 6150

Consider an ensemble (or population) of cumulus clouds consisting of cloudy (saturated) updrafts with temperature  $T_c$  at height  $z_0$ , and speed  $w_c$  and fractional area  $\sigma$  at heights  $z \ge z_0$ , but below cloud top height,  $z_t$ . The clouds are embedded in a clear (unsaturated) environment with temperature  $T_e$  at height  $z_0$ , and initial lapse rate  $\gamma = -dT/dz$  at heights between  $z_0$  and  $z_t$ . Assume that the air between the clouds at height  $z_0$  remains unaffected by mixing with air from the cloudy updrafts during the time period of interest. This is a good assumption when the cumulus cloud top level  $z_t$  is sufficiently far above  $z_0$ , and the time period of interest is sufficiently small.



1. Cumulus cloud updrafts are non-hydrostatic, and are primarily driven by buoyancy, B, so that we can write

$$\frac{dw_c}{dt} \approx B.$$

If we ignore contributions to B from water vapor and condensate loading, B can be expressed mathematically in terms of  $T_c$  and  $T_e$ . Do so.

Answer:

$$g\frac{T_c - T_e}{T_e}$$

2. The horizontally averaged vertical velocity (averaged over cloudy and clear regions) is  $\bar{w}$  at heights at heights between  $z_0$  and  $z_t$ . Express the vertical velocity of the air between the clouds,  $w_e$ , in terms of  $\bar{w}$ ,  $w_c$ , and  $\sigma$ .

Answer:

Conservation of mass requires

$$\bar{w} = \sigma w_c + (1 - \sigma) w_e,$$

where  $w_e$  is the vertical velocity in the clear air between the cloudy updrafts. Solving for  $w_e$  gives

$$w_e = \frac{\bar{w} - \sigma w_c}{1 - \sigma}.\tag{1}$$

3. Assume that the cloudy updrafts originate in a boundary layer whose properties do not change with time, such as might be encountered over an ocean region. Under these conditions, show that

$$\frac{d(T_c - T_e)}{dt} = -\left(\frac{g}{c_p} - \gamma\right)\frac{\sigma w_c - \bar{w}}{1 - \sigma},$$

where g is the acceleration of gravity and  $c_p$  is the specific heat capacity at constant pressure.

#### Answer:

In this case, the temperature of a cloudy updraft at a given height does not change with time, so that  $T_c$  is constant, and  $d(T_c - T_e)/dt = -dT_e/dt$ . The diagram shows that the air between the clouds at height  $z_0$  and time t has descended from a height  $z_0 + h'(t)$  and warmed due to adiabatic compression:

$$T_e(t) = T_0 + (\Gamma_d - \gamma)h'(t),$$

where  $T_0 = T_e(0)$  and  $\Gamma_d = g/c_p = 9.8$  K/km is the dry adiabatic lapse rate. Take the time derivative to get

$$\frac{dT_e}{dt} = (\Gamma_d - \gamma) \frac{dh'(t)}{dt}.$$

Recognizing that  $dh'(t)/dt = -w_e$ , this becomes

$$\frac{dT_e}{dt} = -(\Gamma_d - \gamma)w_e.$$
(2)

This could also be obtained by applying the adiabatic thermodynamic equation,

$$\frac{\partial T}{\partial t} = -w(\Gamma_d + \frac{\partial T}{\partial z}),$$

to the air between the clouds.

Now use (1) to substitute for  $w_e$  in (2) to get

$$\frac{dT_e}{dt} = (\Gamma_d - \gamma) \frac{\sigma w_c - \bar{w}}{1 - \sigma}.$$
(3)

Finally, use  $d(T_c - T_e)/dt = -dT_e/dt$  and  $\Gamma_d = g/c_p$  in (3) to get

$$\frac{d(T_c - T_e)}{dt} = -\left(\frac{g}{c_p} - \gamma\right)\frac{\sigma w_c - \bar{w}}{1 - \sigma}.$$

4. Assume that a sufficient condition for the existence of cumulus clouds at height  $z_0$  is  $T_c > T_e$ . What is the lifetime of the cumulus ensemble (time until  $T_e = T_c$ ) when  $T_c > T_e$  initially and  $\bar{w} = 0$ ?

Answer:

Set  $\bar{w} = 0$  in (3) to obtain

$$\frac{dT_e}{dt} = (\Gamma_d - \gamma)\frac{\sigma w_c}{1 - \sigma}.$$

Integrate  $T_e(t)$  from  $T_0$  to  $T_c$  to get the time, t. The r.h.s. is a constant, so we get

$$t = (T_0 - T_c) \left(\frac{dT_e}{dt}\right)^{-1}.$$

5. Observations show that a typical value of  $\sigma$  is about 0.1. Based on your knowledge of the typical characteristics of cumulus clouds and of the lower atmosphere when cumulus clouds are present, estimate the characteristic lifetime of a cumulus ensemble (not of an individual cumulus cloud) when  $\bar{w} = 0$ .

### Answer:

Typical values are  $T_c - T_e = 1$  K,  $w_c = 2$  m/s, and  $\gamma = 6.5$  K/km. Plugging these values into the answer for Problem 4, we get 1364 s = 23 minutes.

6. Assuming that, initially,  $T_c > T_e$ , under what conditions could a cumulus ensemble persist indefinitely. Give an example, and suggest values of the relevant parameters.

#### Answer:

To obtain  $dT_e/dt = 0$ , we must have either  $\gamma = \Gamma_d$  or  $w_e = 0$ . Rarely is  $\gamma = \Gamma_d$  through a deep layer of the troposphere, so this condition is not the relevant one. The condition for  $w_e = 0$  is

$$\bar{w} = \sigma w_c$$

For  $\sigma = 0.1$  and  $w_c = 2$  m/s,  $\bar{w} = 0.2$  m/s. For  $\sigma = 0.01$  and  $w_c = 20$  m/s,  $\bar{w} = 0.2$  m/s.

7. Express  $dT_e/dt$ , as given by (3),

$$\frac{dT_e}{dt} = (\Gamma_d - \gamma) \frac{\sigma w_c - \bar{w}}{1 - \sigma},$$

in terms of the dry static energy of the environment,  $s_e = c_p T_e + gz$ , when  $\sigma \ll 1$ .

Answer:

$$\frac{1}{c_p}\frac{ds_e}{dt} = \frac{dT_e}{dt},$$
$$\frac{1}{c_p}\frac{\partial s_e}{\partial z} = \frac{\partial T_e}{\partial z} + \frac{g}{c_p} = \Gamma_d - \gamma,$$

and

$$\frac{\sigma w_c - \bar{w}}{1 - \sigma} \approx \sigma w_c - \bar{w},$$

so (3) becomes

$$\frac{ds_e}{dt} = -\bar{w}\frac{\partial s_e}{\partial z} + \sigma w_c \frac{\partial s_e}{\partial z}.$$