

Lake-Effect

Characteristics, Processes and Lake-Orographic Interactions

Atmos 5210
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Learning Objectives

- After this class you should be able to
 - Recognize several ways that lakes and complex terrain affect the morphology and intensity of lake-effect storms
 - Use this recognition to better analyze and predict lake-effect storms



Outline

- Introduction
- The Great Salt Lake (GSL)
- The GSL-effect
 - Morphology and climatology
 - Environmental conditions
 - Shoreline band dynamics
- Orographic effects
- Japan
- Mechanisms of Orographic Enhancement



Introduction



What Is Lake Effect Snow?

- “Lake effect snow is produced when cold winds move across long expanses of warmer lake water, *picking up water vapor*, which freezes and is deposited on the lee shores”
 - Wikipedia.com (2006)
- “Precipitation occurring near or downwind from the shore of a lake resulting from the *warming (destabilization) and moistening* of relatively cold air during passage over a warm body of water”
 - Glossary of Meteorology (2000)

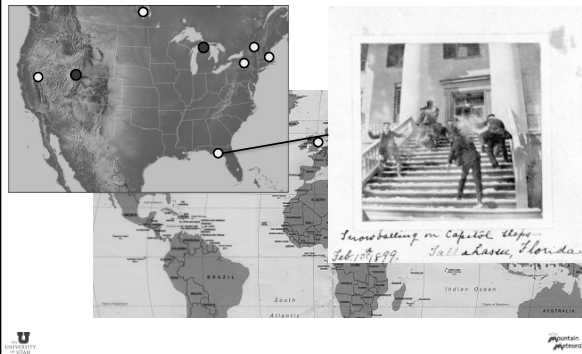


Additional Factors

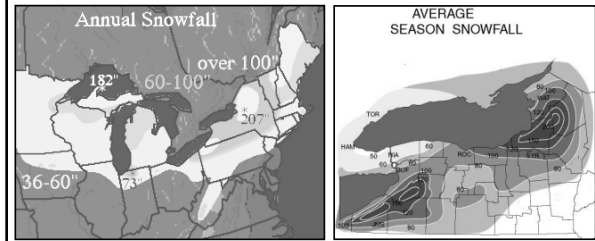
- Lake-lake interactions & aggregate effects (Great Lakes)
- Boundary layer & thermally driven circulations
- Orography
 - Mountains, hills, plateaus, and windward/leeward coastal geometry
- Surface roughness contrasts
- Ice cover



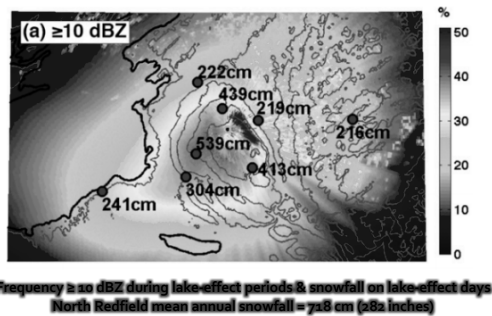
Lake-Effect Regions



Great Lakes Climatology



Lake Ontario and Tug Hill

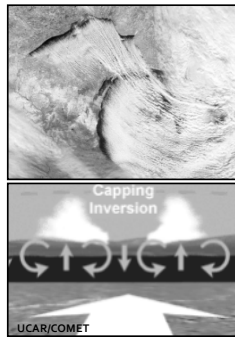


Great Lakes Snowfall Records

- 24 h: 77" @ Montague, NY
- Storm: 95" @ Montague, NY (10-14 Jan 97)
- Month: 192" @ Bennett Bridges, NY (Jan 1978)
- Winter: 466.9" @ Hooker, NY (1976-77)

Event Types

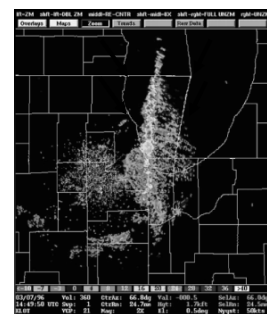
- Wide-area coverage
 - Wind parallel bands (a.k.a. cloud streets)
 - Horizontal roll convection
 - Rolls oriented along boundary layer shear vector



NASA, UCAR/COMET

Event Types

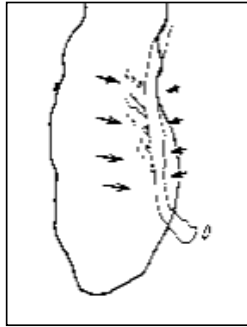
- Shore-parallel bands
 - Type I: MidLake
 - Flow along major lake axis
 - Large lake-land temperature difference
 - Land-breeze fronts form one convergence zone



Paul Markowski, Penn State University

Event Types

- Shore-parallel bands
 - Type II: Shoreline bands
 - Form near lee shore
 - Land breeze opposes large-scale flow



Hjelmfelt (1990)

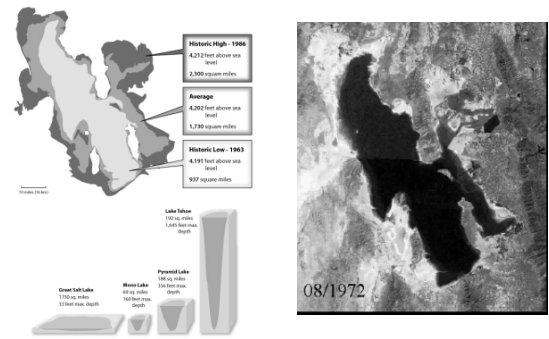
The Great Salt Lake

GSL Physiography

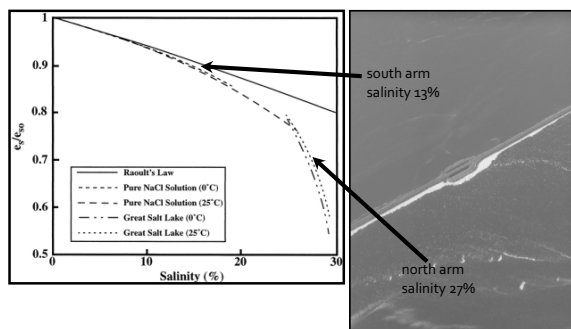
- Length: 120 km
 - Lake Ontario: 355 km
- Width: 45 km
 - Lake Ontario: 85 km
- Area: 2500–5900 km²
 - Lake Ontario: 59,000 km²
- Ave depth: 4.5 m
 - Lake Ontario: 86 m
- Surrounded by intense, steeply sloped mountains
 - Lake Ontario: Tug Hill (500 m above lake level)
- Salinity
 - Gilbert Bay: 6–12‰
 - Gunnison Bay: 27‰



Elevation, Depth, and Area

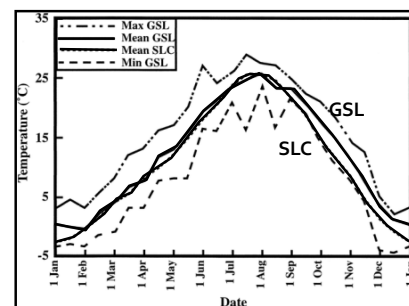

http://learn.genetics.utah.edu/content/gsl/physical_char/

Salinity



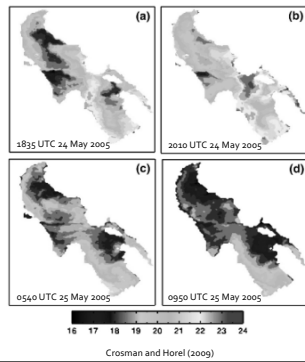
Dickson et al. (1965), Steenburgh et al. (2000)

Temperature Climatology



Steenburgh et al. (2000)

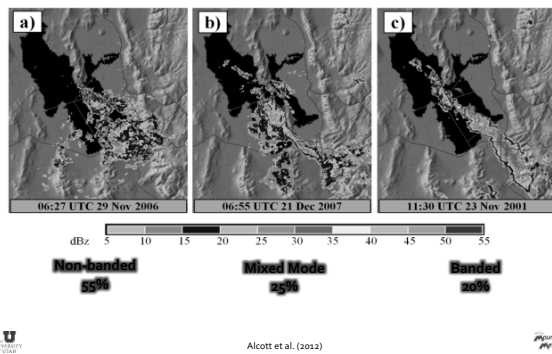
Spatio-Temporal Variability



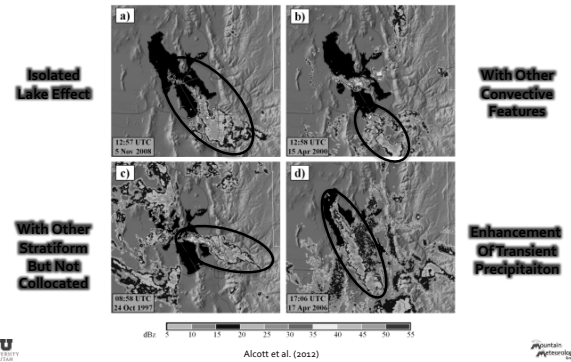
The GSL-Effect

Morphology and Climatology

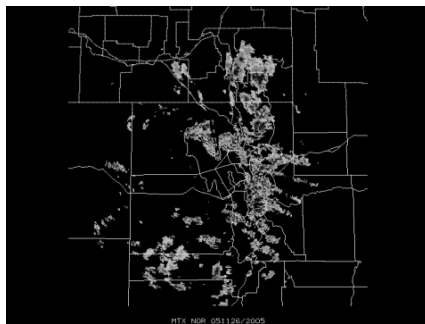
Morphology



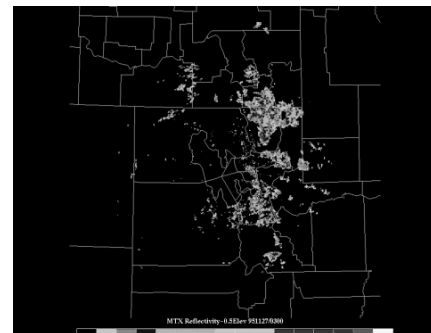
Pure and "Contaminated"



Weakly or Non Banded Event



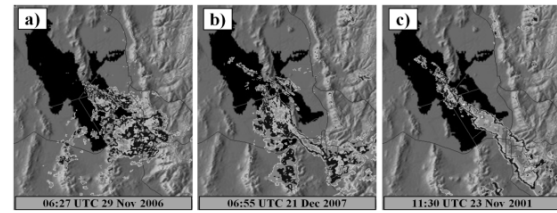
Banded Event



Cloud Band



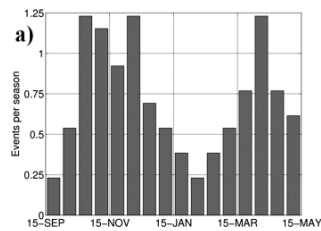
Discussion



Non-banded 55%
Mixed Mode 25%
Banded 20%

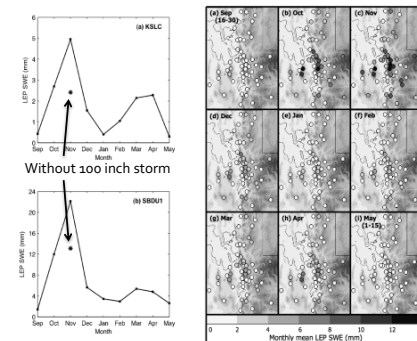
What controls the morphology of lake-effect systems?

Seasonality

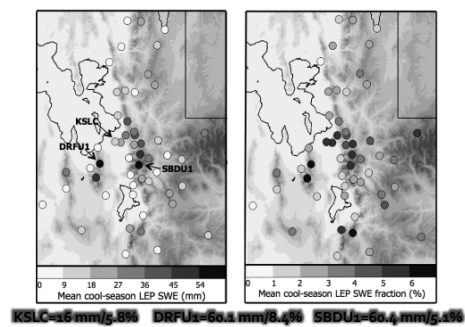


- Bimodal seasonality in event frequency
- Primary LEP SWE peak in fall

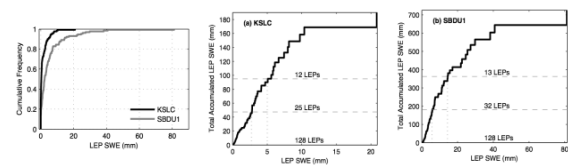
Seasonality



Mean Cool Season SWE



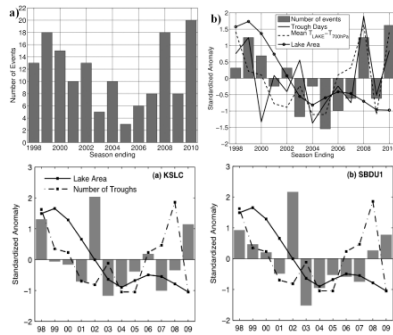
Rare but Intense Events Dominate



Lots of small events
> 10 mm very rare (<2%) @ KSLC
> 20 mm < 5% of events @ SBDU1

12(13) events produced 50% of the LEP SWE at KSLC (SBDU1) from 1998-2009
One LEP during 100 inch storm is a real outlier

Interannual Variability



Alcott et al. (2012), Yeager et al. (2013)

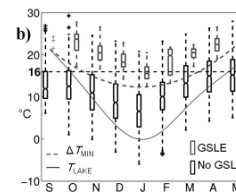
The GSL-Effect Environmental Conditions

Key Ingredients

- Instability
 - lake–700-mb $\Delta T > 16^\circ\text{C}$
 - 16°C is approximately a dry adiabatic lapse rate
 - Alcott et al. (2012) show important exceptions exist
 - Events are moist and exclusively in Dec–Feb
- Upstream Moisture
- Wind Direction/Fetch
- Land-Breeze Convergence

Steenburgh et al. (2000), Alcott et al. (2012)

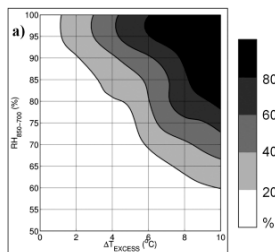
Instability



- 93% of events occur at lake–700-mb $\Delta T \geq 16^\circ\text{C}$
 - Exceptions occur exclusively during winter months (“moist” events)
 - Fall and spring events occur at higher thresholds
 - Suggests a seasonally varying threshold (ΔT_{excess}) is more appropriate than a fixed threshold

Alcott et al. (2012)

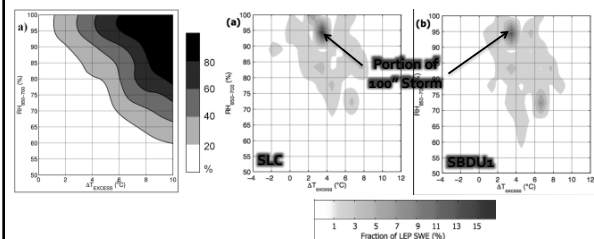
Moisture



Fraction of soundings with GSL as a function of (ΔT_{excess}) and 850–700-mb RH
Moisture matters, avoid “thermonymia”

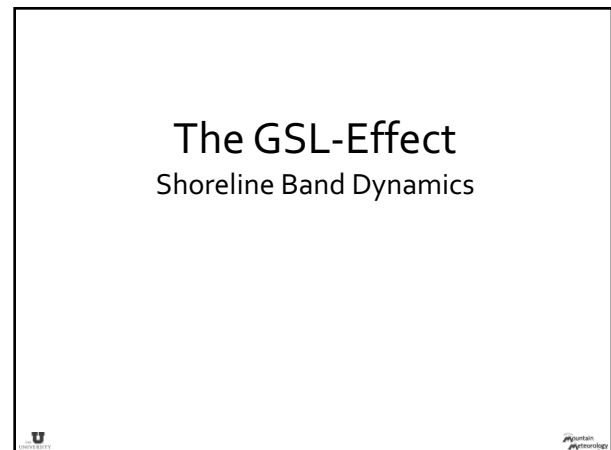
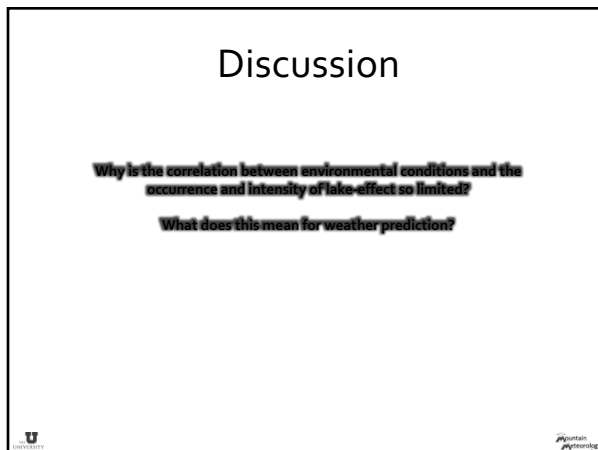
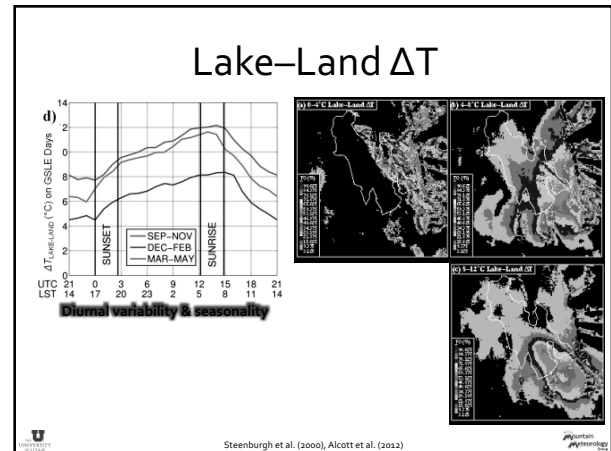
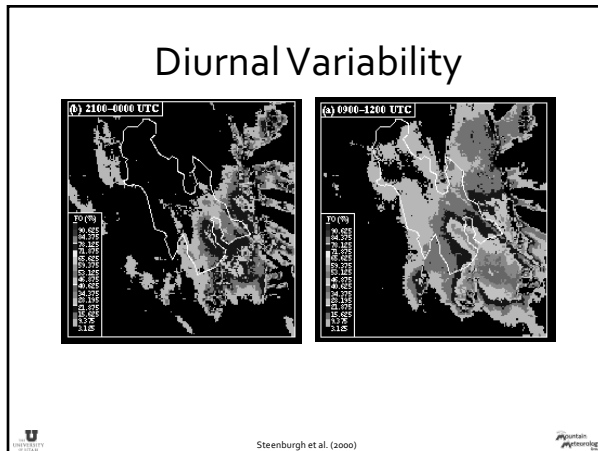
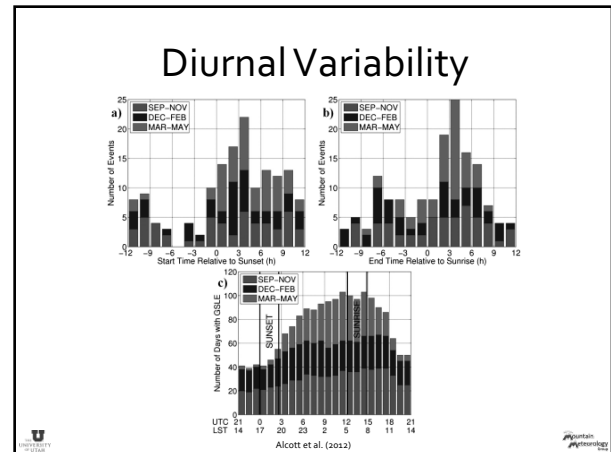
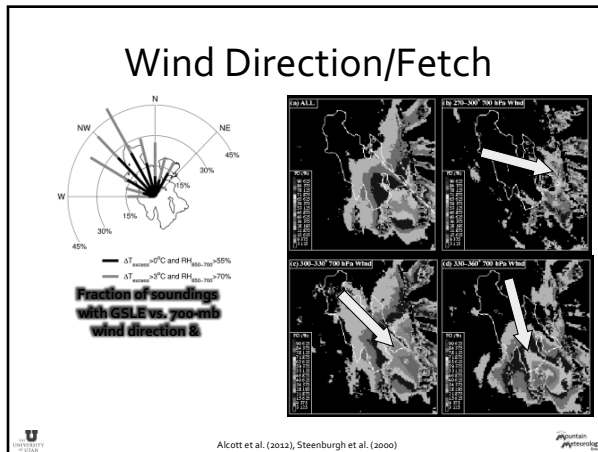
Alcott et al. (2012)

More on Instability/Moisture



Event intensity/LEP SWE not well correlated with ΔT_{excess} and moisture
Other processes, poorly resolved by current obs and analyses, are likely important

Alcott et al. (2012), Yeager et al. (2013)

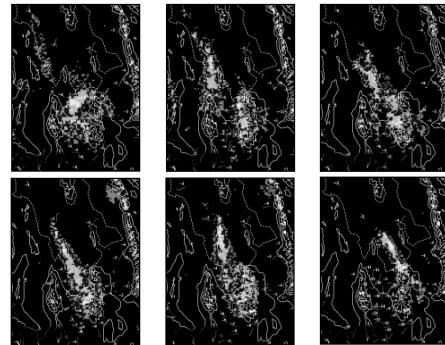


7 December 1998 Event

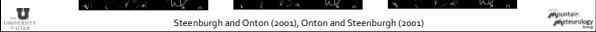
- Two lake-effect snowbands merged into a shoreline band
- Heaviest snow (up to 36 cm) concentrated in a narrow, 10-km wide band over Tooele County



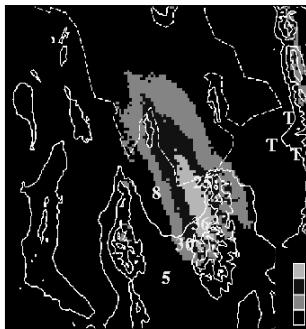
Mesoscale Evolution



Steenburgh and Onton (2003), Onton and Steenburgh (2003)



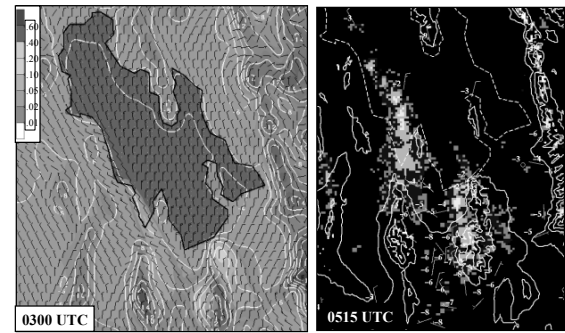
Total Snowfall (cm)



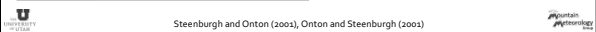
Steenburgh and Onton (2003), Onton and Steenburgh (2003)



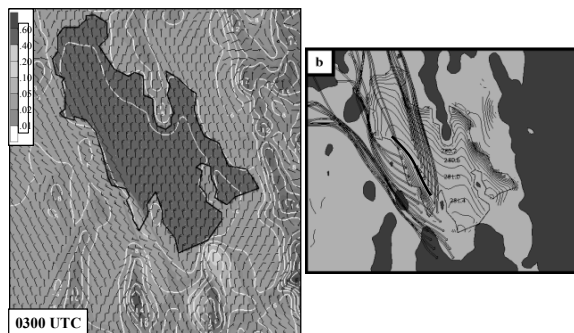
Initiation Phase



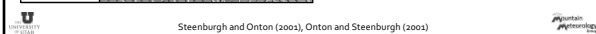
Steenburgh and Onton (2003), Onton and Steenburgh (2003)



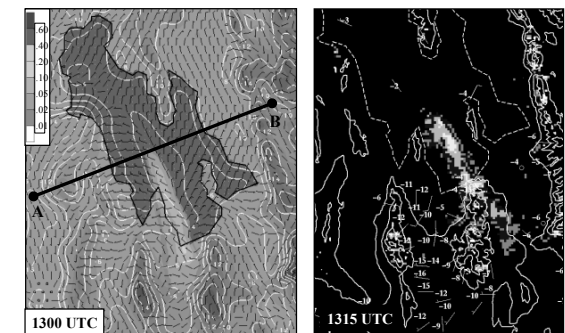
Initiation Phase



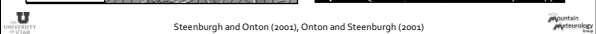
Steenburgh and Onton (2003), Onton and Steenburgh (2003)



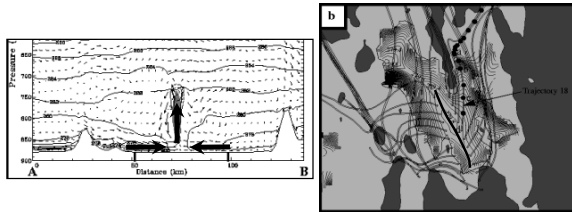
Mature Phase



Steenburgh and Onton (2003), Onton and Steenburgh (2003)

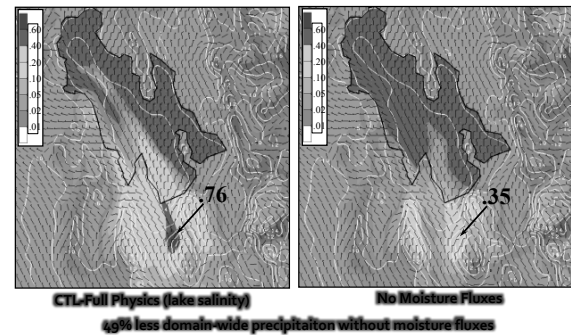


Mature Phase



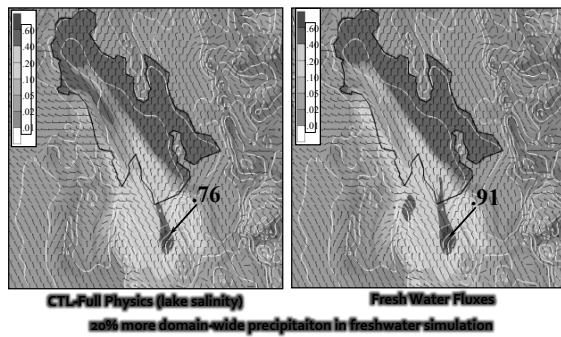
Steenburgh and Onton (2003), Onton and Steenburgh (2003)

Moisture Flux Sensitivity



Steenburgh and Onton (2003), Onton and Steenburgh (2003)

Moisture Flux Sensitivity



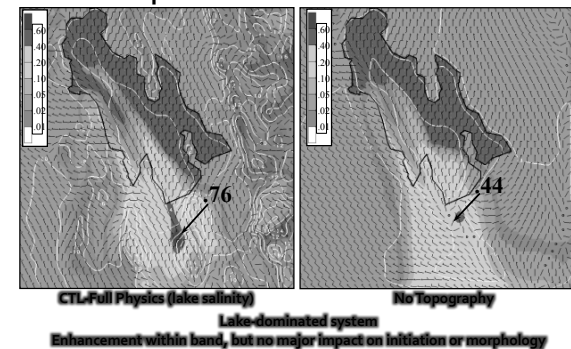
Steenburgh and Onton (2003), Onton and Steenburgh (2003)

Orographic Effects

Potential Orographic Effects

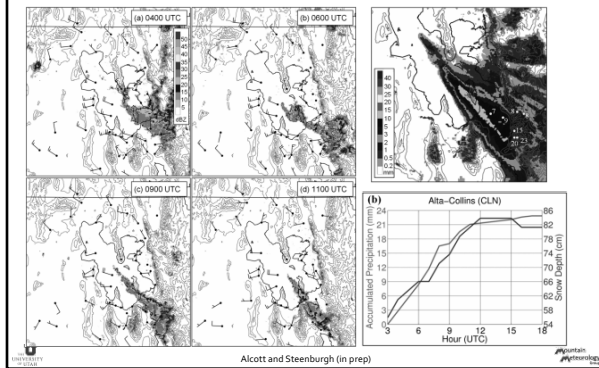
- Precipitation enhancement
- Modification of the lake-effect system
 - Initiation, intensity, morphology
- Lake-effect systems can be altered by upstream and downstream topography

Precipitation Enhancement

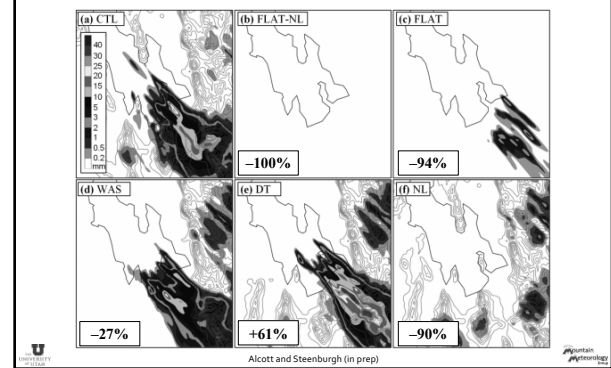


Steenburgh and Onton (2003), Onton and Steenburgh (2003)

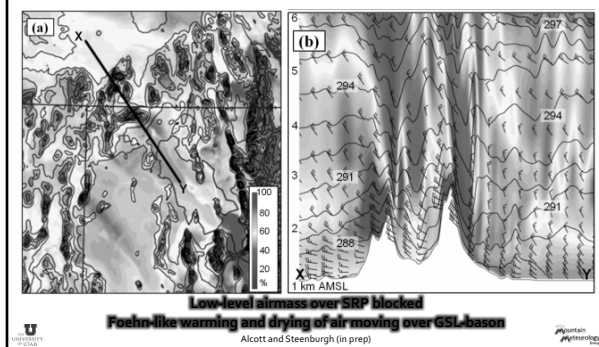
System Modification: 27 Oct 2010



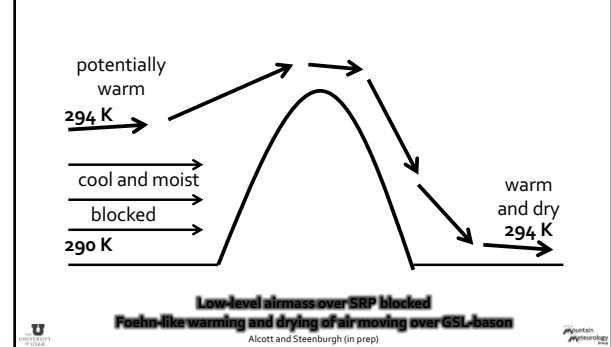
Terrain Sensitivity



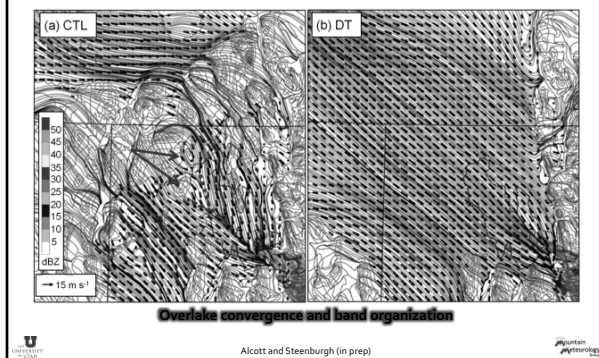
Upstream Influences



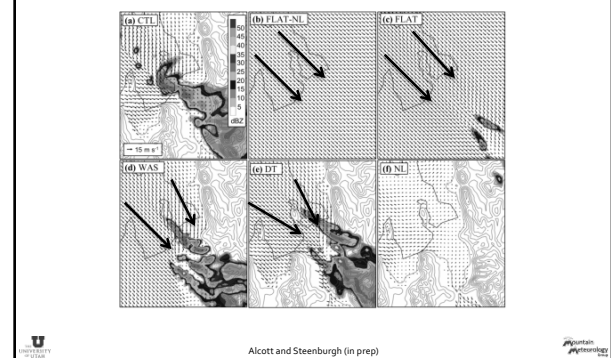
Upstream Influences



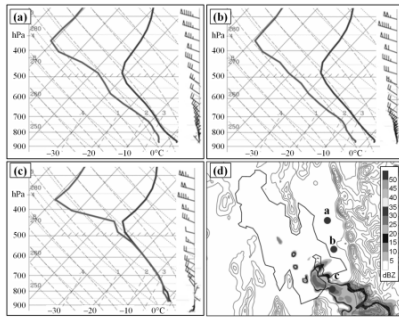
Upstream Influences



Downstream Influences



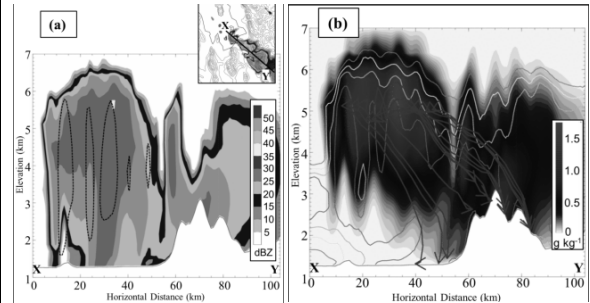
Blocking



Stable to dry motions in area around GSL

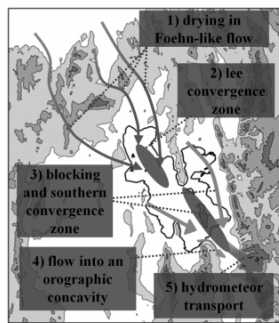
Alcott and Steenburgh (in prep)

Microphysical Considerations



Alcott and Steenburgh (in prep)

27 Oct 2010 Summary



Alcott and Steenburgh (in prep)

Discussion

7 Dec 1998

27 Oct 2010

Lake
Dominated

Orographic
Dominated

Little sensitivity to the
presence of the GSL

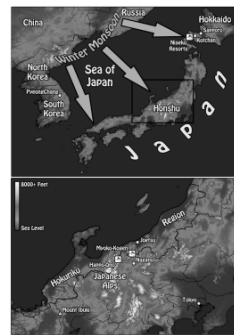
What controls the relative contribution of lake and orographic effects
on lake-effect systems?

Why are some events "lake dominated"?

Why are others influenced strongly by lake-orographic interactions?

Japan

Hokuriku District

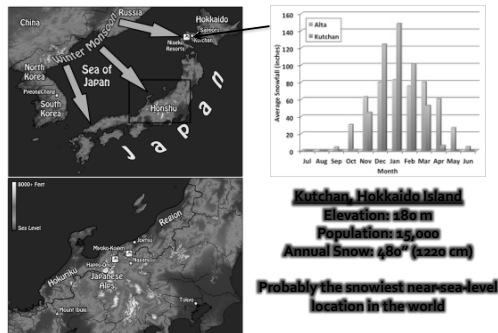


Joetsu, Hokuriku District
Elevation: sea level
Population: 205,000
Annual SWE: 108" (276 cm)
Annual Snow: 248" (630 cm)
Record Snow: 590" (1499 cm)

Probably the snowiest urban area in the world

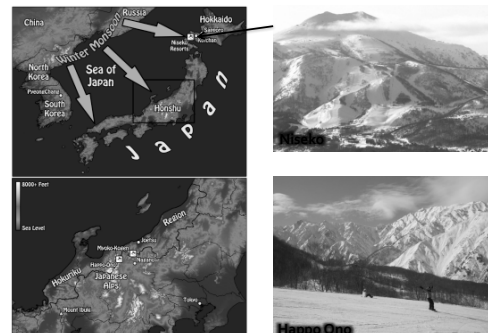
Steenburgh (2014), GoogleEarth, JMA via Wikipedia

Hokkaido Island



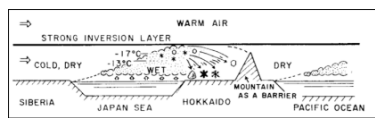
Steenburgh (2014), JMA via Wikipedia

Mountains



Steenburgh (2014), <http://www.jnto.go.jp/tourism/en/sooy.html>

Importance of the CAP



CAP=Capping Inversion or Stable Layer

Satoyuki Storms Produce more snow in the coastal lowlands
Yamayuki Storms Produce more snow in the mountains

Hypothesis Height of CAP relative to mountain crest affects orographic ratio in lake-effect storms

Magono et al. (1966)

Mechanisms of Orographic Enhancement

Conventional Wisdom

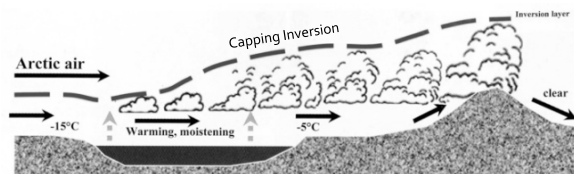
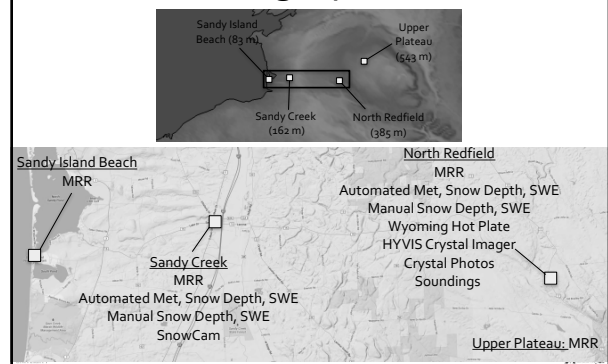


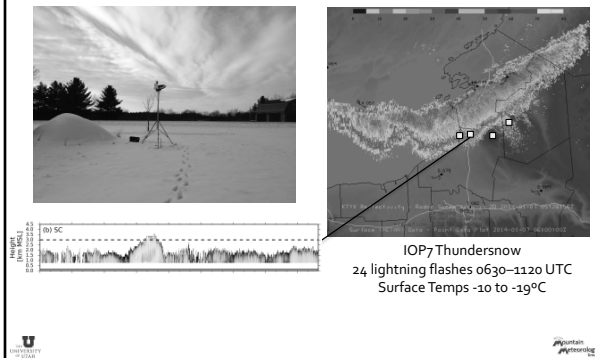
Figure 9.21. Idealized cross-sectional schematic of LES development, with downstream orographic lift. Depth of the inversion along the flow shown as a dashed red line.

Orographic lifting invigorates convection
 (i.e., larger updrafts speeds and cloud depths)
 Lackmann (2011)

OWLeS Orographic Transect



MicroRainRadars (MRRs)

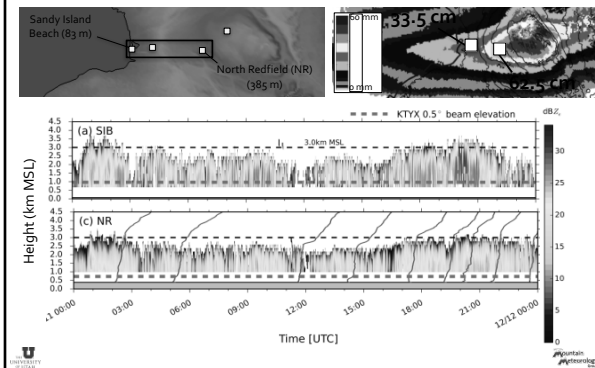


Snow and Water Equivalent Obs

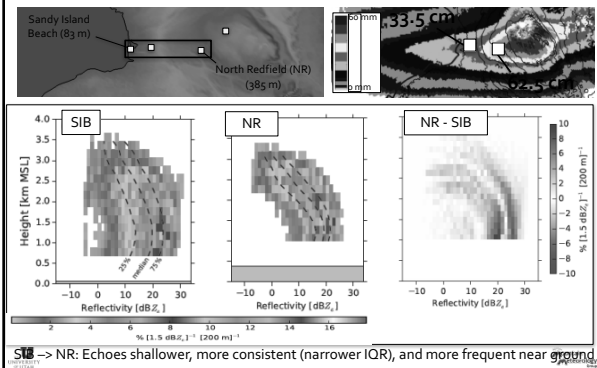
Collected @ Sandy Creek (lowland, 162 m MSL) and North Redfield (upland, 385 m MSL)



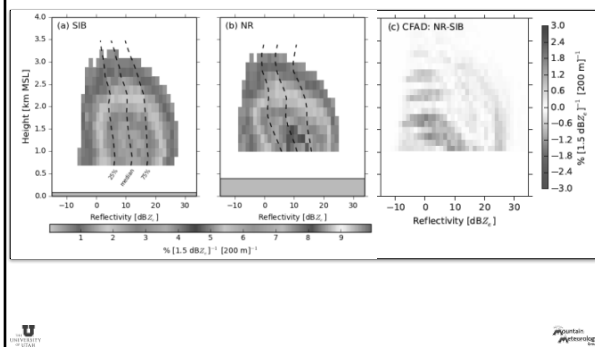
IOP2: MRRs



IOP2 CFADs



All 29 Events @ SIB & NR



Conventional Wisdom

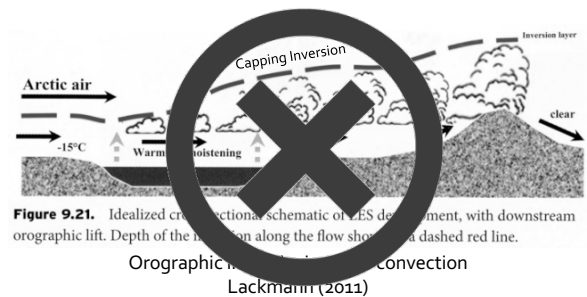


Figure 9.21. Idealized cross-sectional schematic of CES development, with downstream orographic lift. Depth of the inversion along the flow shown as a dashed red line.

Overall enhancement: See Minder et al. 2015, <http://dx.doi.org/10.1175/MWR-D-15-0117.1>
OR variations: See Leah Campbell's talk next

Summary

- Many processes influence lake-effect systems
 - Upstream instability and moisture
 - Lake conditions (surface temperature, sub-surface temperature, salinity, ice cover)
 - Land breezes and PBL circulations
 - Orography
- Orographic influences not only include precipitation enhancement, but also the initiation, intensity, and morphology of lake-effect systems



References

- Alcott, T. I., and W. J. Steenburgh, in prep: Orographic influences on a Great Salt Lake-effect snowstorm. *Mon. Wea. Rev.*, To be submitted. Latest draft available at <http://www.jrcc.utah.edu/~steenburgh/papers/OLE-orographic.pdf>.
- Alcott, T. I., W. J. Steenburgh, and N. F. Laird, 2012: Great Salt Lake-effect precipitation: Observed frequency, characteristics, and environmental factors. *Wea. Forecasting*, **27**, 954-971.
- Croaman, E., and J. Horel, 2009: MODIS-derived surface temperature of the Great Salt Lake. *Remote Sensing of Environment*, **113**, 73-81.
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