

# Radiation measurements

Motivation (Energy Balance)

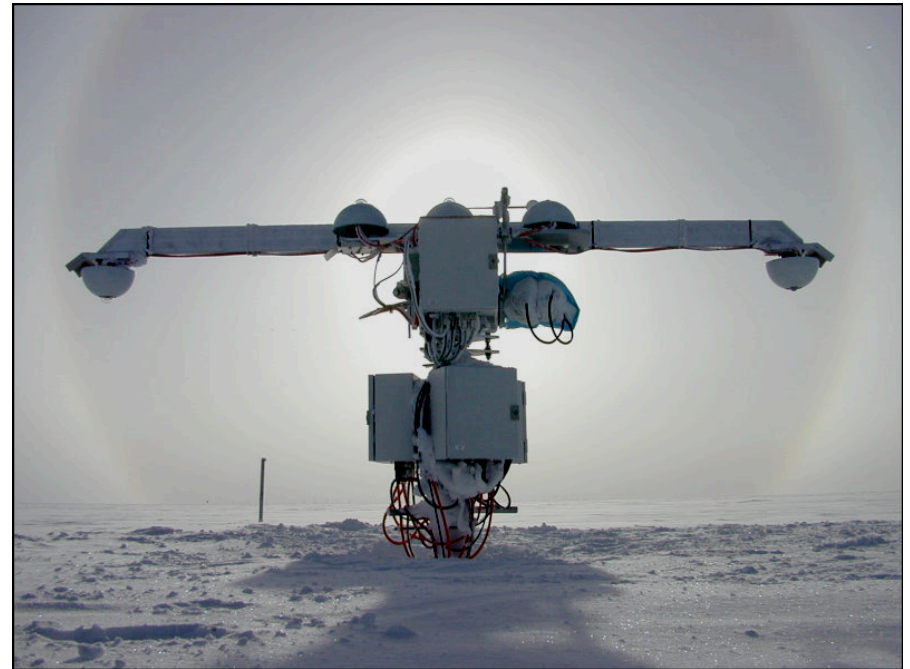
Background

Radiation Quantities & Terms in Radiation Budget

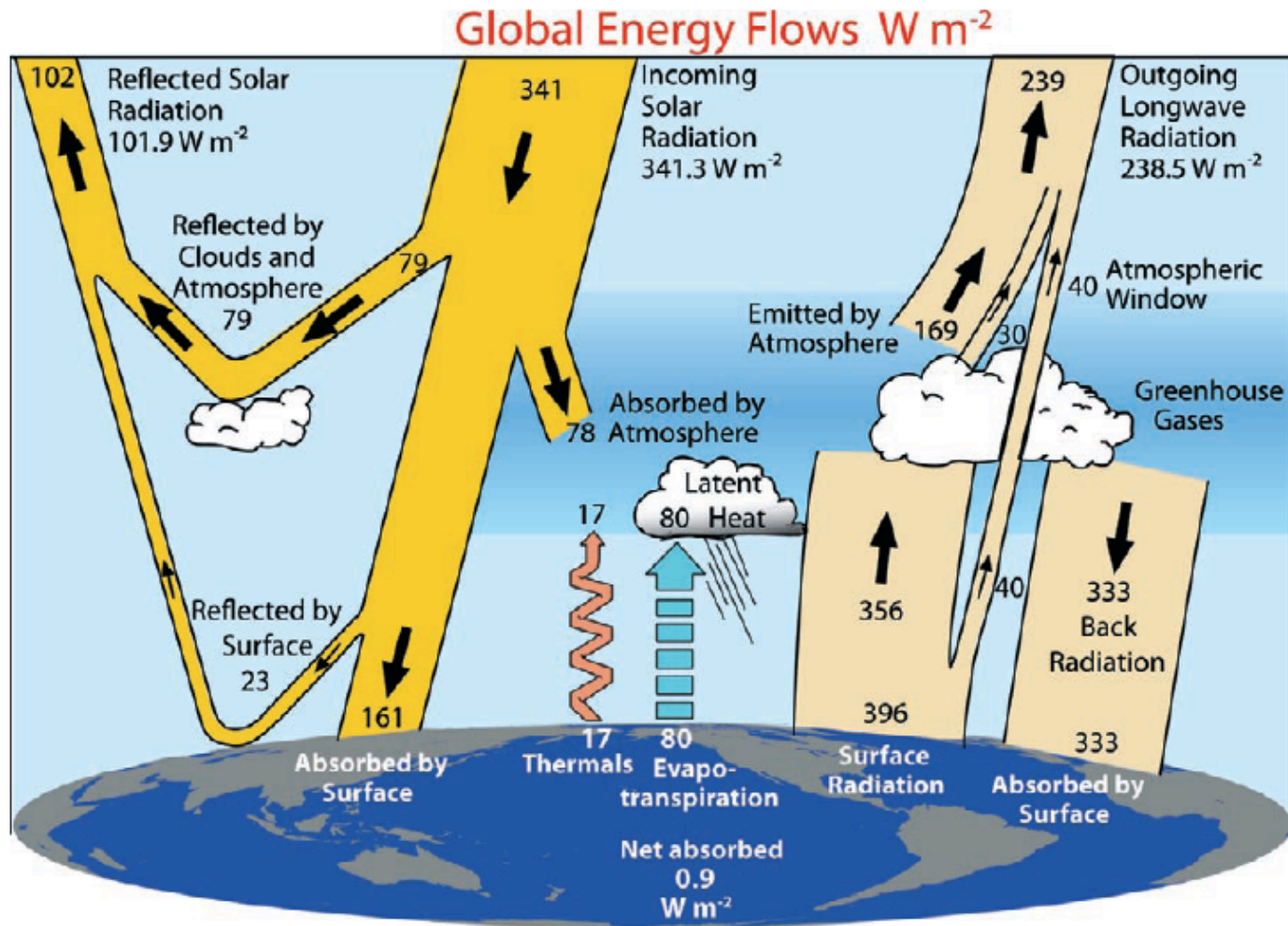
Instrumentation & Measurement Principles

Radiation Balance in different climates

Sebastian W. Hoch  
485 INSCC



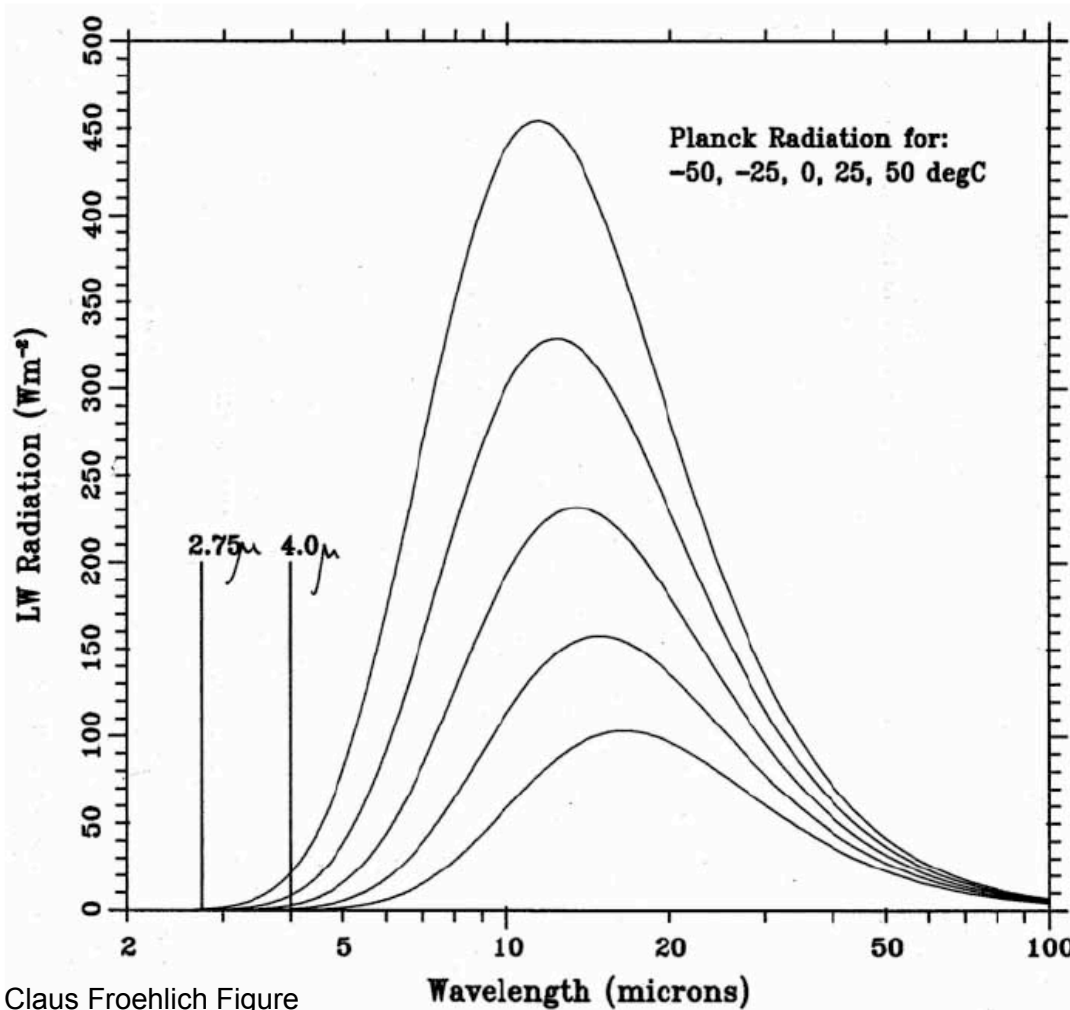
# Radiation and the Energy Budget



Trenberth et al. 2009 BAMS



# Why Shortwave Radiation and Longwave Radiation?



Planck curves for 5 different temperatures.

Planck Function

$$B_{\lambda}(T) = \frac{2hc^2}{\lambda^5 (e^{hc/k\lambda T} - 1)}$$

$h = 6.626\ 068\ 96(33) \times 10^{-34}$  J s

$c$ : speed of light ...

$\lambda$ : wavelength

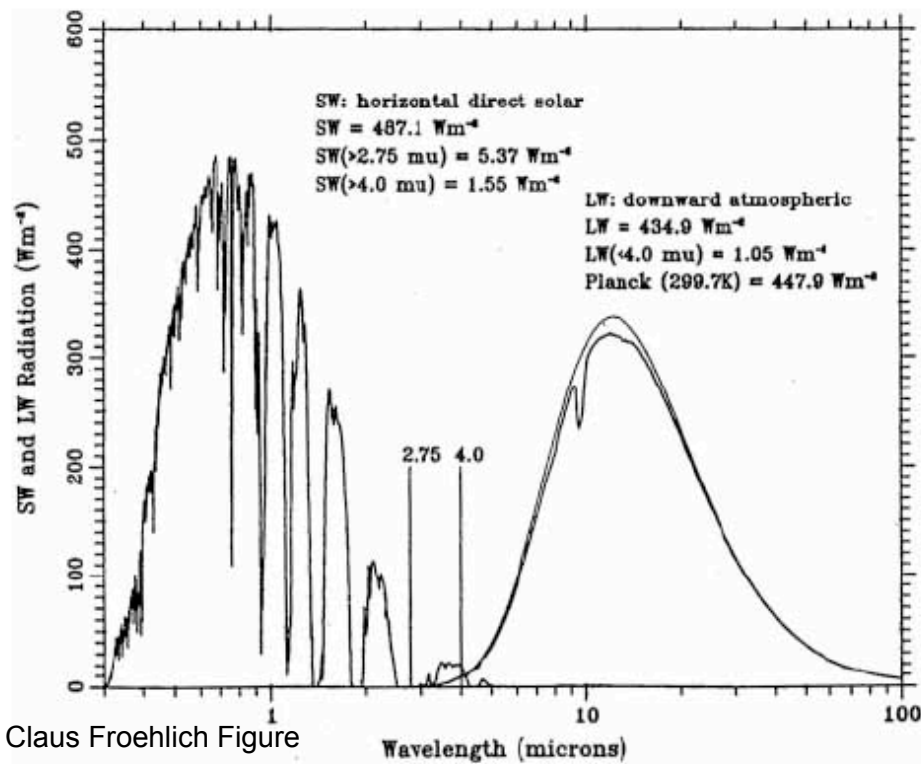
Everything emits radiation – depending on the temperature!

Wien's Displacement Law

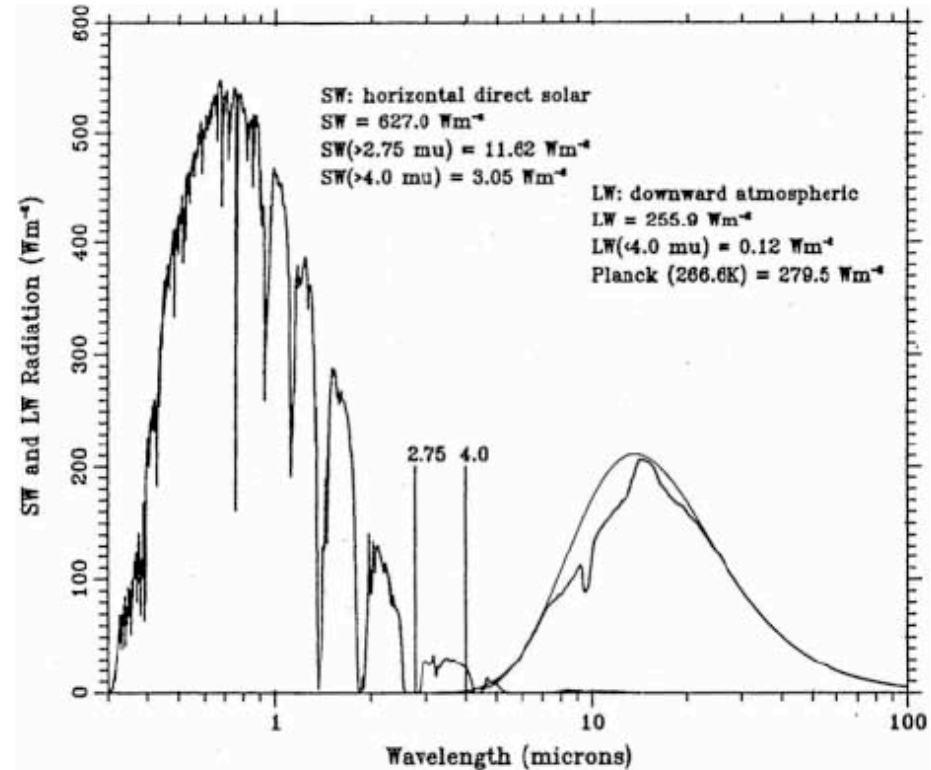
$$\lambda_{\max} = \frac{a}{T} \text{ mit } a = 2.89776 \times 10^{-3} \text{ mK}$$

# Longwave or Terrestrial or Infrared Radiation and

# Shortwave or Solar Radiation



Tropical atmosphere



Midlatitude Summer, 1600 m (ASL)

There's an overlap at times ...

Other quantities defined by spectral range:

- UV Radiation (A, B, C)
- PAR: Photosynthetically Active Radiation; 400 - 700 nm

# Radiation Quantities

Quantity	Symbol	SI unit	Abbr.	Notes
Radiant energy	$Q$	joule	J	energy
Radiant flux	$\Phi$	watt	W	radiant energy per unit time, also called <i>radiant power</i>
Radiant intensity	$I$	watt per steradian	$W \cdot sr^{-1}$	power per unit solid angle
Radiance	$L$	watt per steradian per square metre	$W \cdot sr^{-1} \cdot m^{-2}$	power per unit solid angle per unit <i>projected</i> source area. called <i>intensity</i> in some other fields of study.
Irradiance	$E, I$	watt per square metre	$W \cdot m^{-2}$	power incident on a surface. sometimes confusingly called "intensity".
Radiant exitance / Radiant emittance	$M$	watt per square metre	$W \cdot m^{-2}$	power emitted from a surface.
Radiosity	$J$ or $J_\lambda$	watt per square metre	$W \cdot m^{-2}$	emitted plus reflected power leaving a surface
Spectral radiance	$L_\lambda$ or $L_\nu$	watt per steradian per metre <sup>3</sup> or watt per steradian per square metre per hertz	$W \cdot sr^{-1} \cdot m^{-3}$ or $W \cdot sr^{-1} \cdot m^{-2} \cdot Hz^{-1}$	commonly measured in $W \cdot sr^{-1} \cdot m^{-2} \cdot nm^{-1}$
Spectral irradiance	$E_\lambda$ or $E_\nu$	watt per metre <sup>3</sup> or watt per square metre per hertz	$W \cdot m^{-3}$ or $W \cdot m^{-2} \cdot Hz^{-1}$	commonly measured in $W \cdot m^{-2} \cdot nm^{-1}$

# The Radiation Balance – the terms (Irradiances $W m^{-2}$ )

Direct Solar Radiation  $S_{\downarrow}$

Diffuse (Solar) Radiation  $D_{\downarrow}$

Global Radiation ( $K_{\downarrow}$ ,  $G_I$ ) =  $S_{\downarrow} + D_{\downarrow}$

Shortwave Reflected Radiation  $K_{\uparrow}$

Shortwave Net Radiation  $K^*$

$$\text{Albedo } \alpha = K_{\uparrow} / K_{\downarrow}$$

Longwave Incoming  $L_{\downarrow}$

Longwave Outgoing Radiation  $L_{\uparrow}$

Longwave Net Radiation  $L^*$

Net Radiation  $Q^*$

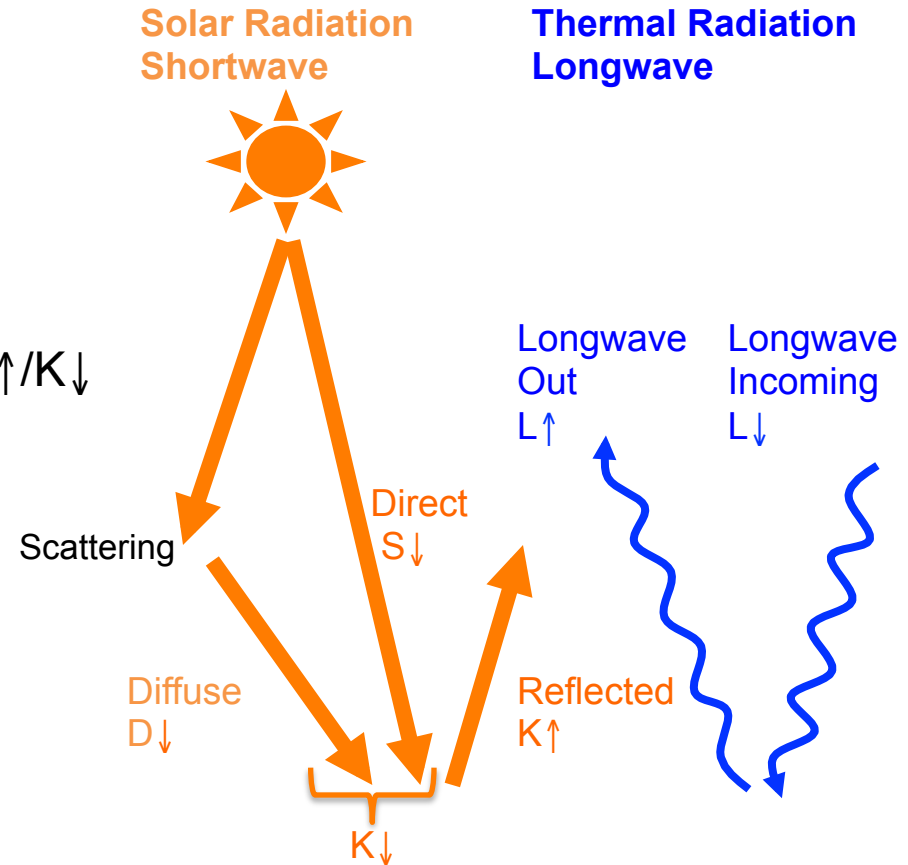
$$Q^* = K^* + L^*$$

$$= K_{\downarrow} - K_{\uparrow} + L_{\downarrow} - L_{\uparrow}$$

$$= (1 - \alpha) \cdot K_{\downarrow} + L^*$$

$$L_{\uparrow} = \epsilon_{\text{surf}} \cdot \sigma \cdot T_{\text{surf}}^4$$

$$L_{\downarrow} = \epsilon_{\text{atmos}} \cdot \sigma \cdot T_{\text{atmos}}^4$$



Stefan-Boltzmann Constant  
 $\sigma: 5.67 \cdot 10^{-8} J s^{-1} m^{-2} K^{-4}$



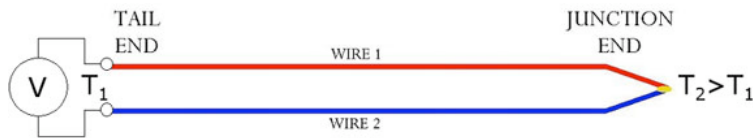
# Measurement Principle

## Thermopile

- converts thermal energy into electrical energy
- composed of thermocouples (usually in series)
- output voltage proportional to a local temperature difference
- range of tens or hundreds of millivolts.

## Thermocouple

- temperature measurement based on the *Seebeck Effect*: a result of a difference in *thermoelectric power* of two materials



$$Emf = \int_{T_1}^{T_2} S_{12} \cdot dT = \int_{T_1}^{T_2} (S_1 - S_2) \cdot dT$$

- Emf is the Electro-Motive Force or Voltage;  $T_1$  and  $T_2$ : Temperatures of reference ( $T_1$ ) and measuring end ( $T_2$ )
- $S_{12}$ ,  $S_1$ ,  $S_2$ : **Seebeck coefficients** of the thermocouple and thermo-elements
- null voltage:
  - same materials
  - no temperature difference

# Radiation observations in Climate Science - Instrumentation

## 1. Pyrheliometer

Direct Solar Radiation

World Standard Instruments

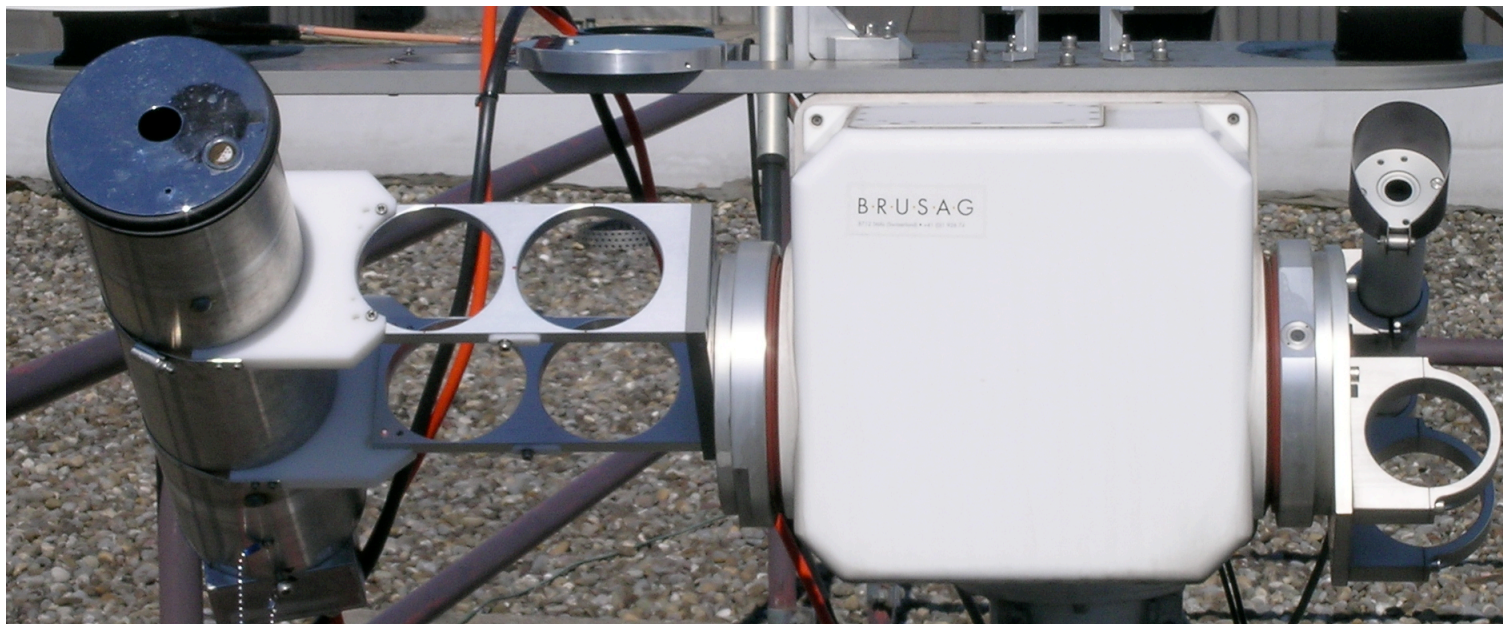
(Compensation Type / Thermopile)

Open / with window ...

pyro-, pyr- +  
(Greek: fire, burn; heat,  
produced by heating; and  
sometimes "fever")

"ἥλιος (Helios) is derived  
from the noun ἥλιος,  
"sun" in ancient Greek

PMO-6



Kipp & Zonen  
CH1

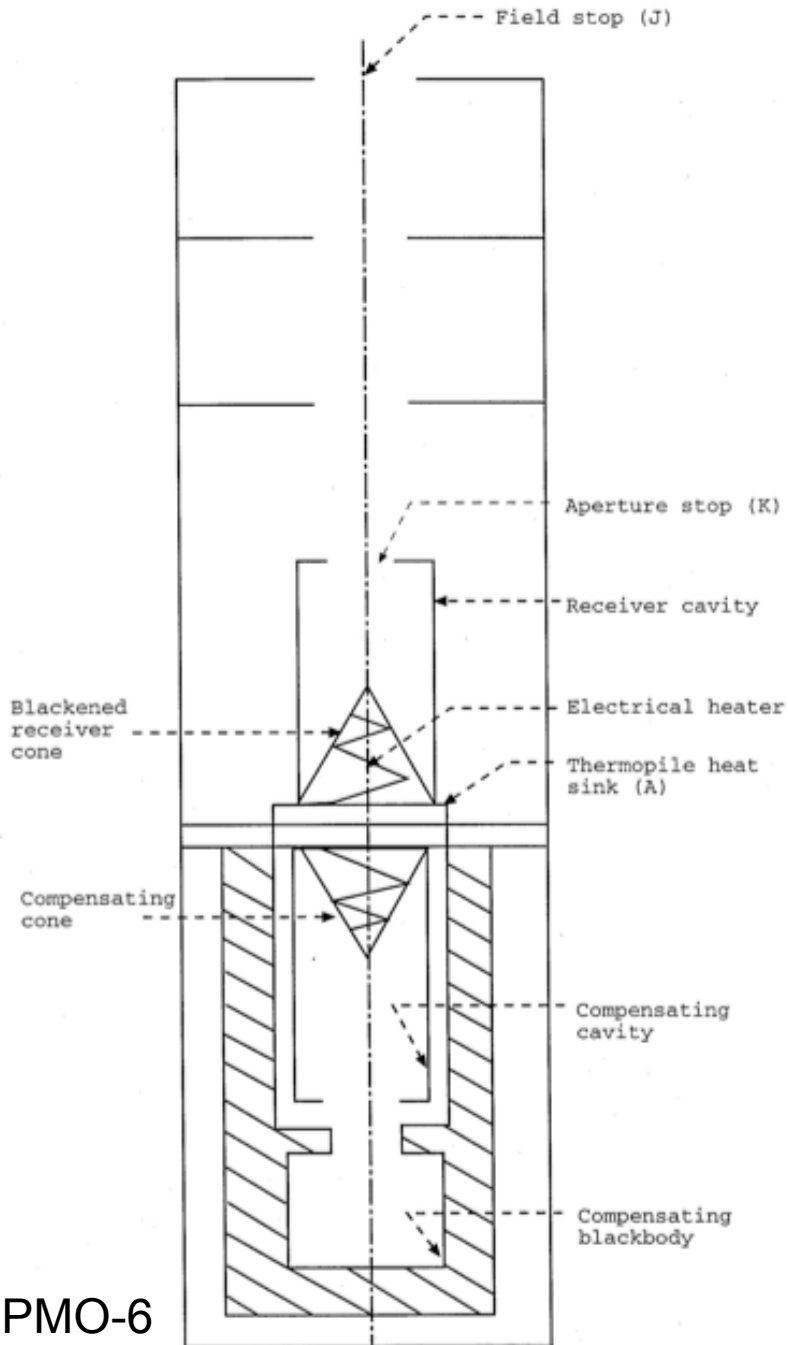
# PMO-6 Absolute Cavity Radiometer

$$S = k * (P_{\text{closed}} - P_{\text{open}})$$



pmod/wrc

Other System:  
Eppley Hickley-Frieden (HF)



PMO-6

NREL-TP-463-20619

# Thermopile Pyrheliometer (NIP / CH1)

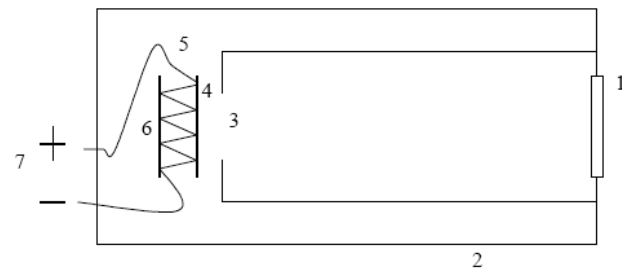
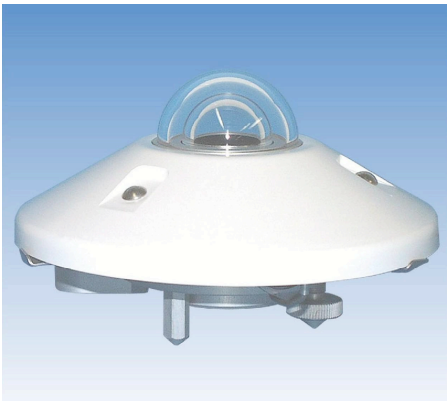
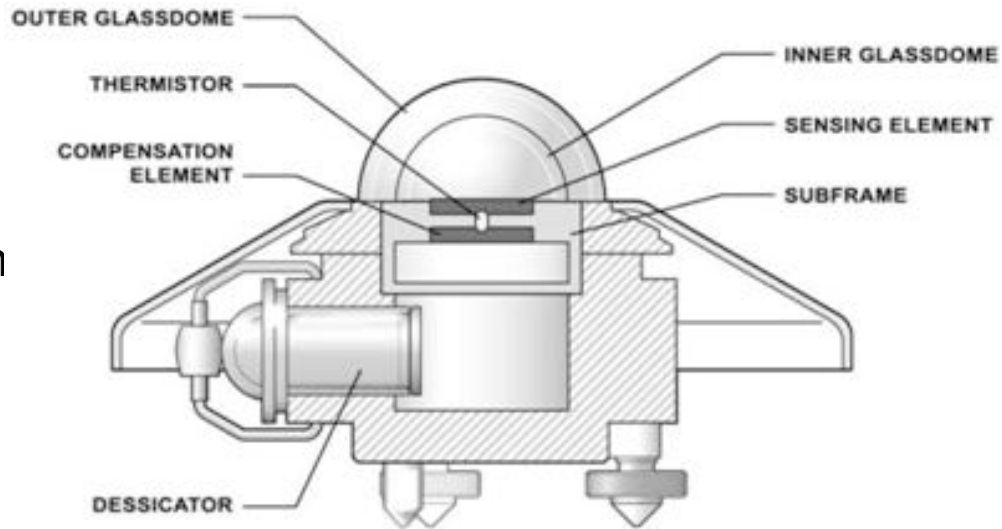


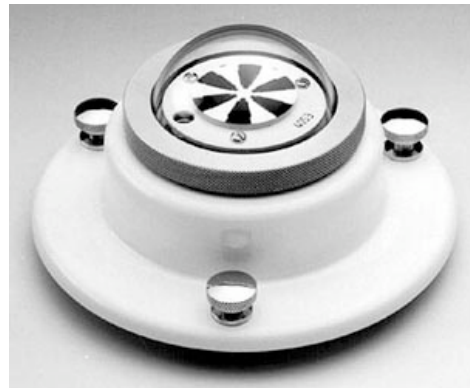
Fig. 1. Pyrheliometer schematic showing entrance window (1), thermal shield (2), detector aperture (3), light absorber (4), thermopile (5), heat sink (6), and thermopile output (7).

## 2. Pyranometer

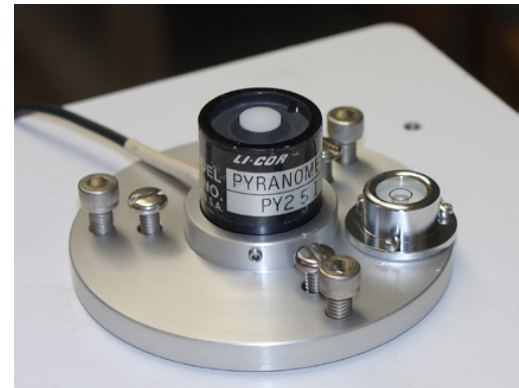
- Global Radiation
- Shortwave Reflected Radiation
- Diffuse Radiation (in conjunction with a shading disk or shadow-ban)
- Glass or Quarz dome



Standard



Black & White Type



Photodiode Type



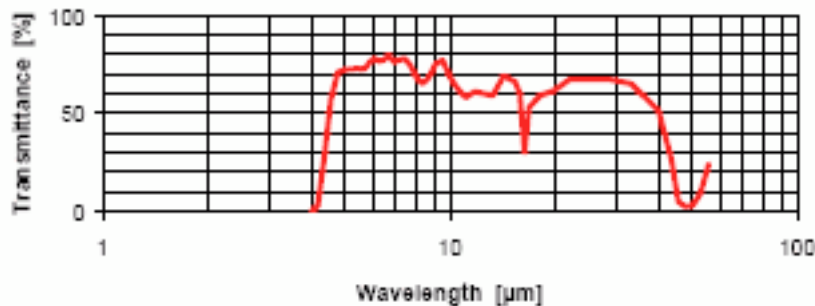
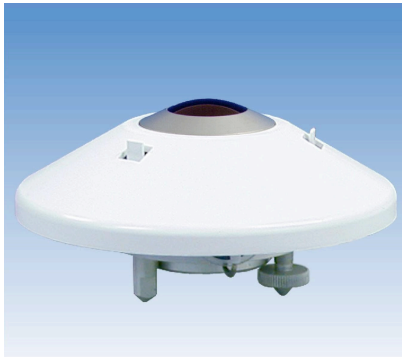
## Shading – Shadowbands and Shading disks



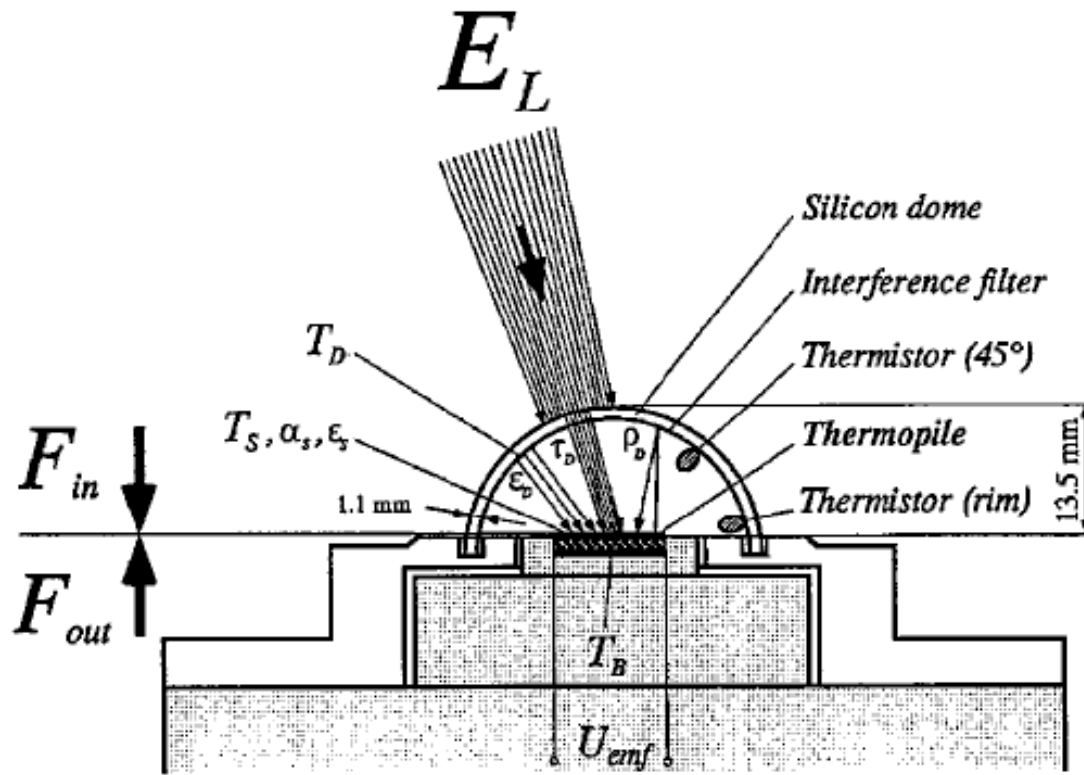
### 3. Pyrgeometer

Longwave Radiation  
Thermopile, Silicon (Si) dome

geo-, ge- +  
(Greek: earth, land,  
soil; world)



Si-Dome and interference filters



Schematic view of Eppley PIR (Philipona et al. 1995)

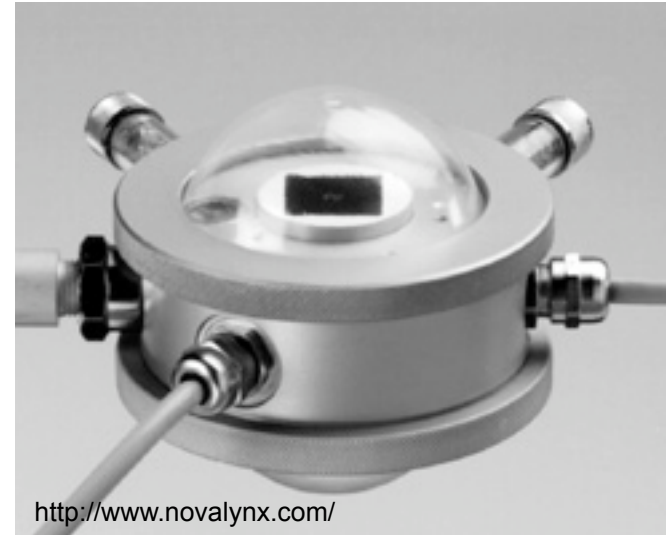
Pyrgeometer Formula:

$$E_L = \frac{U_{emf}}{C} \left( \underbrace{1}_{LWin\_a} + \underbrace{k_1 \sigma T_B^3}_{LWin\_b} \right) + \underbrace{k_2 \sigma T_B^4}_{LWin\_b} - \underbrace{k_3 \sigma (T_D^4 - T_B^4)}_{LWin\_c}.$$

We neglect  $k_1$ , set  $k_2$  to 1.0, and  $k_3$  to a mean value of 3.5.

## 4. Pyrradiometer

- “All-wave” Radiation
- Thermopile measurements
- Polyethylene Dome
- Double domes: Net-Radiometer
- “Wind Speed Error”



Different response  
to short- and  
longwave fluxes!

Birds like to destroy them, too...



## 5. Heliograph / Sunshine Duration Sensor



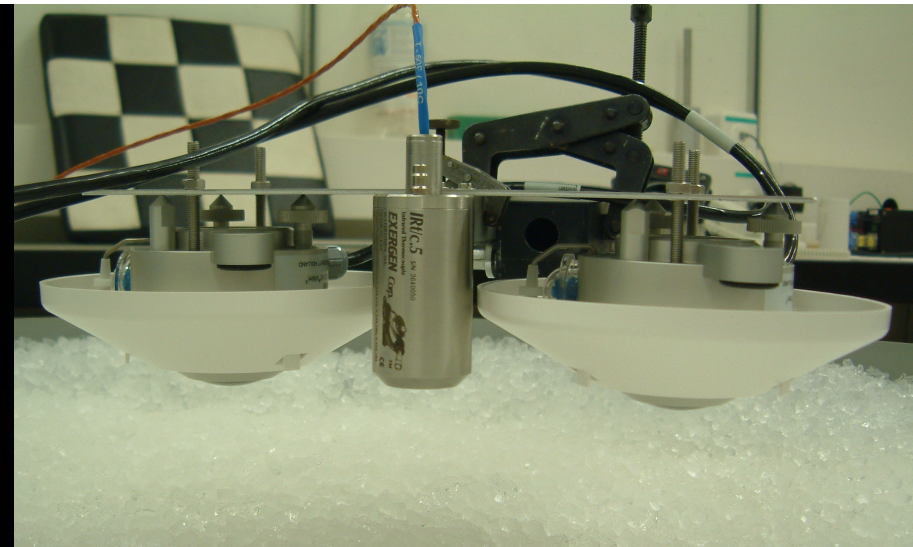
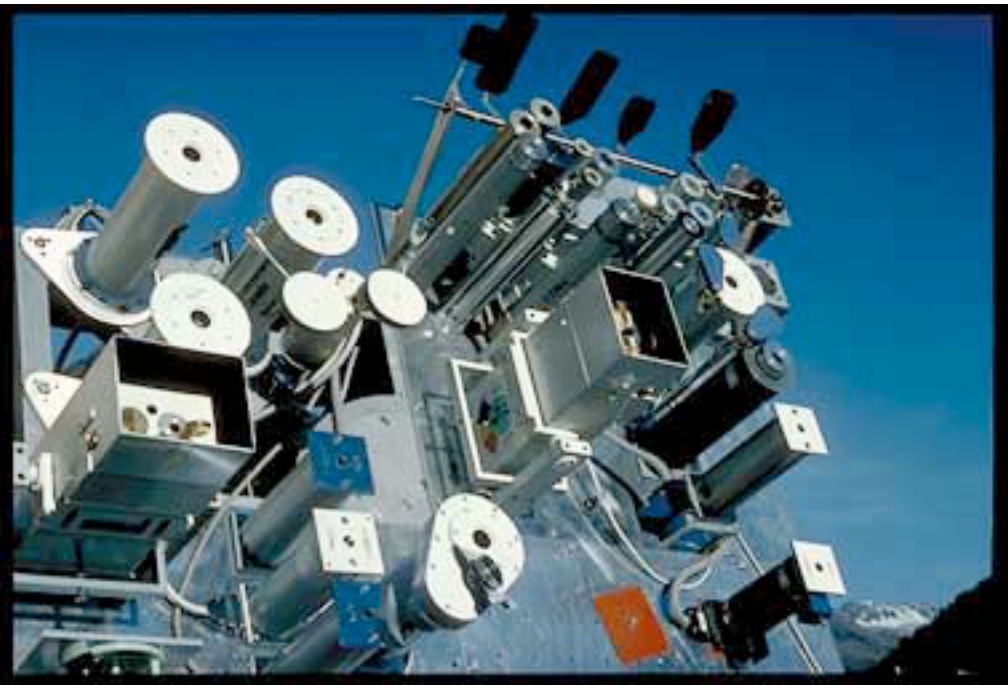
Campbell-Stokes Sunshine Recorder

“Sunshine”: Flux >  $120 \text{ Wm}^{-2}$



One end of an optical fiber revolves around the sun axis. The opening angle is limited by an optical diaphragm. At the other end, a photovoltaic detector receives the light pulse when the fiber window meets the sun. The detected signal is compared to a threshold. A pulse is generated when the radiation intensity exceeds  $120 \text{ W/m}^2$ .

# Calibrations and Errors



WSG (World Standard Group)  
Davos, Switzerland

- Absolute Calibration Error  
(Comparison to World Standard)



# Spectral Response Errors

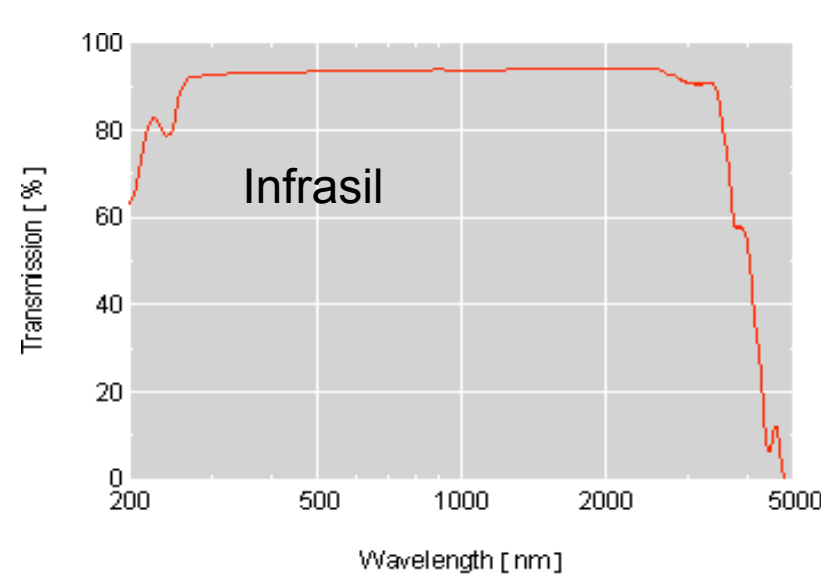
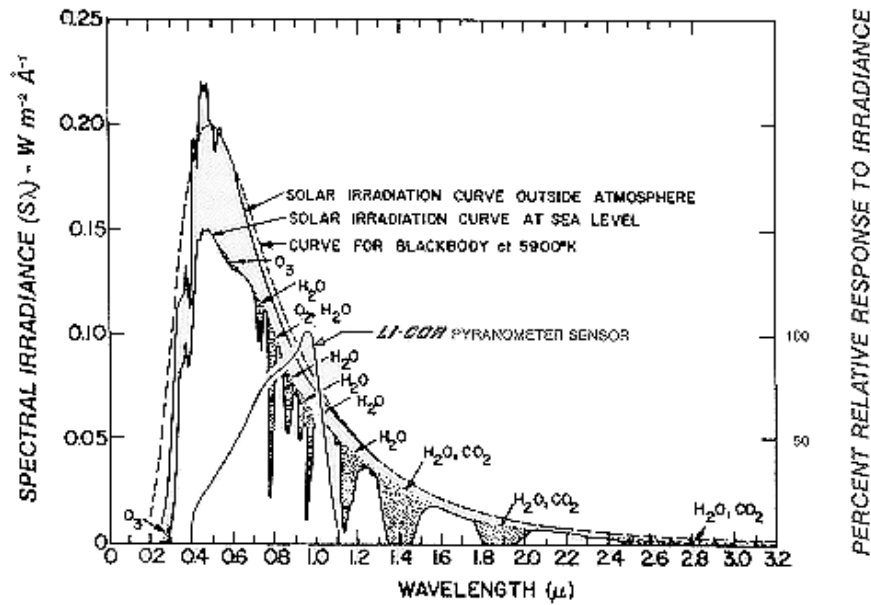
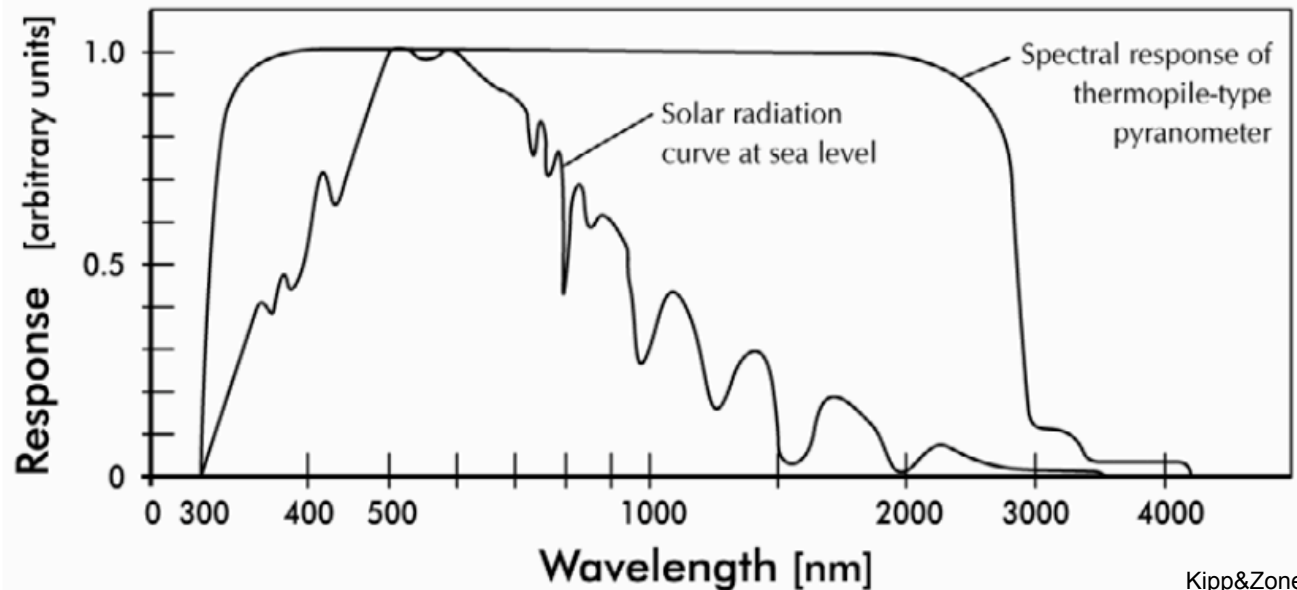
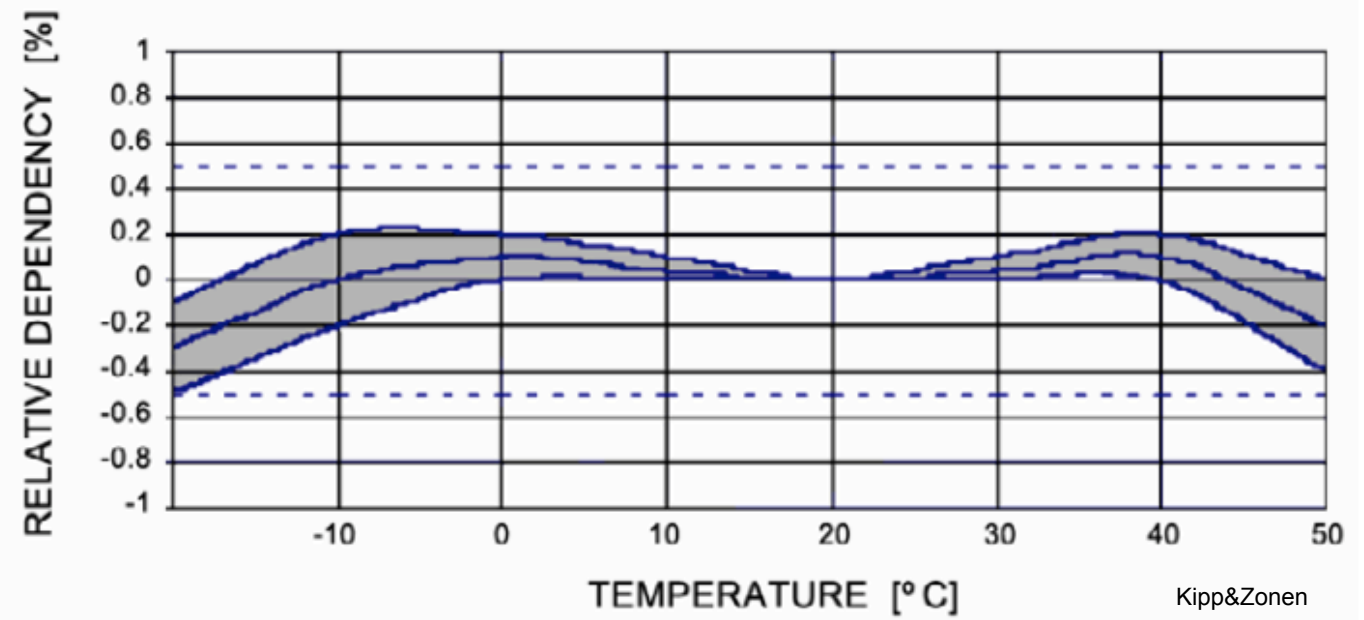


Figure 4. The LI-200SA Pyranometer spectral response is illustrated along with the energy distribution in the solar spectrum (8).

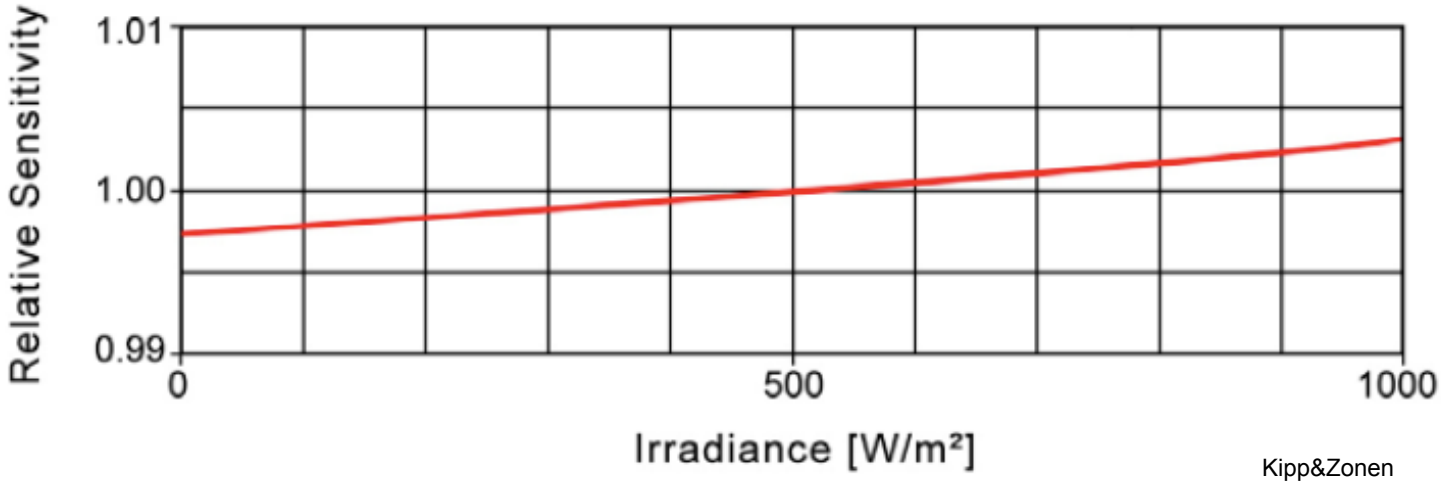




- Temperature Dependency



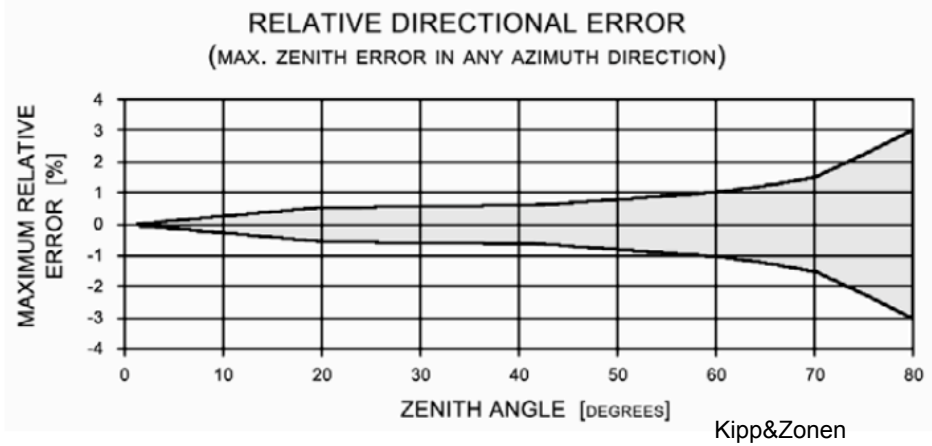
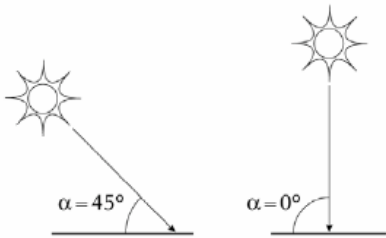
- Linearity



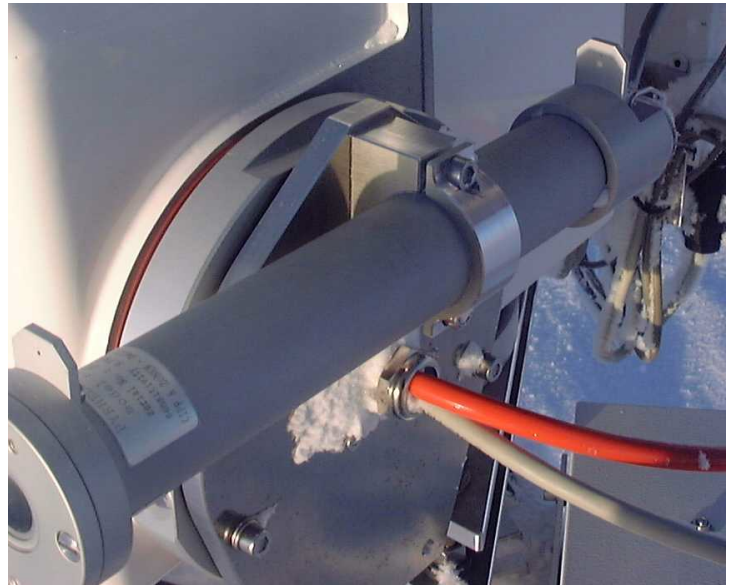


# Geometric Errors:

- Cosine Response Error  
(low vs high incident radiation)
- Azimuth error (sensor geometry)



- Hysteresis
- Response Time Error
- Long Term Stability (aging of thermopile / paint / resistors / etc)

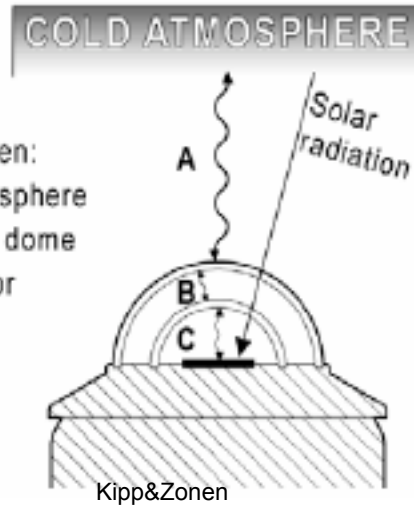


Pointing error



Condensation

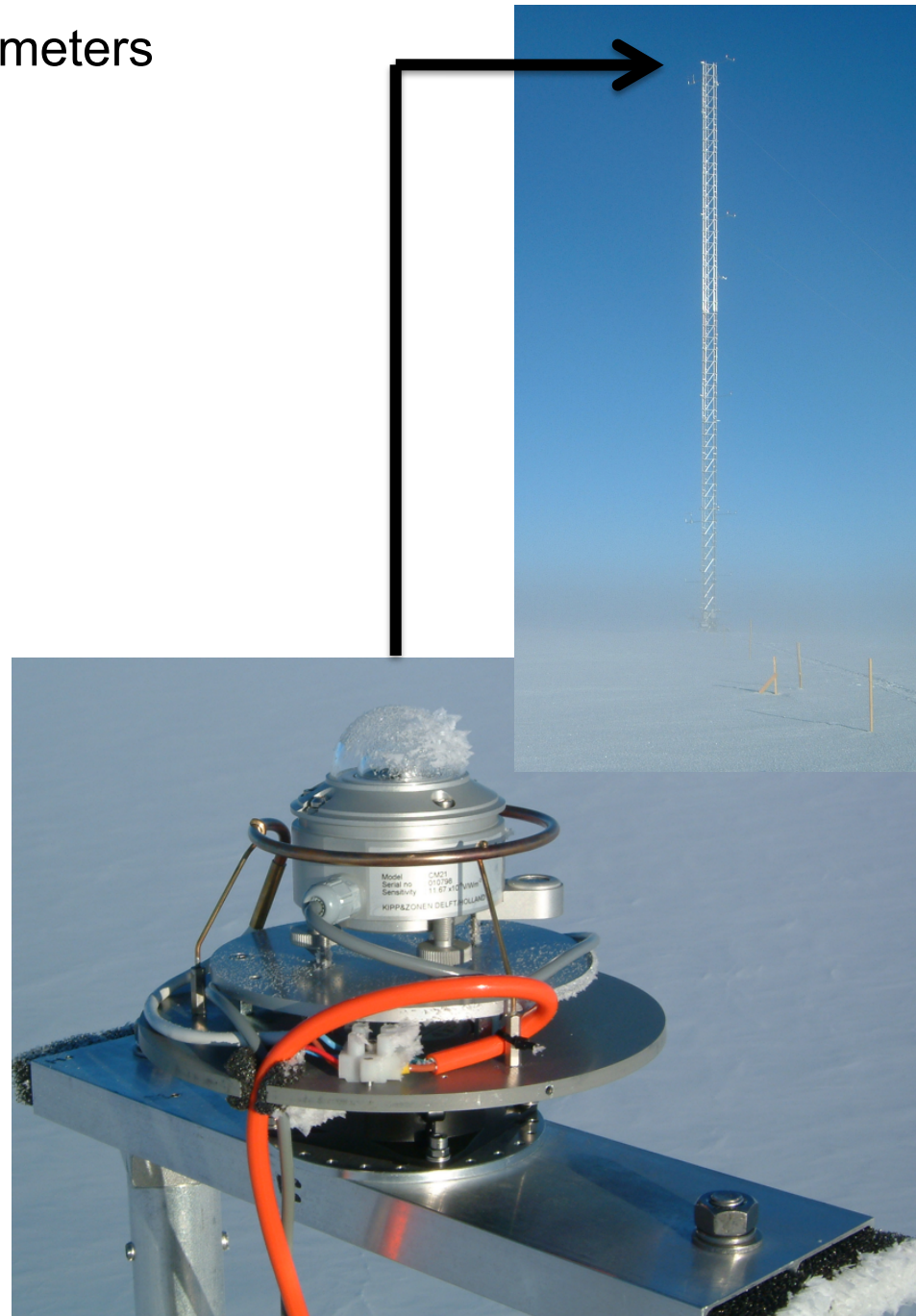
# Negative-Night-Time-Offset of Pyranometers



Ventilation and Heating!

## “Wind-Correction” of Pyrradiometers and Net-Radiometers

Dome material (polyethylene, lupolene) heats up. Ventilation reduces the heating effect.





# Environmental impacts



IGLOS 2002



# Radiation balance measurement at Summit, Greenland

## 2 Pyrgeometers (Eppley PIR):

- Longwave Incoming
- Longwave Outgoing

Shadow Discs

## 3 Pyranometers (Kipp&Zonen CM21/11):

- Diffuse Radiation
- Global Radiation
- Shortwave Reflected

Absolute Cavity Radiometer (PMO6)  
Reference Instrument

Pyrheliometer (Kipp&Zonen CH1):  
• Direct Radiation

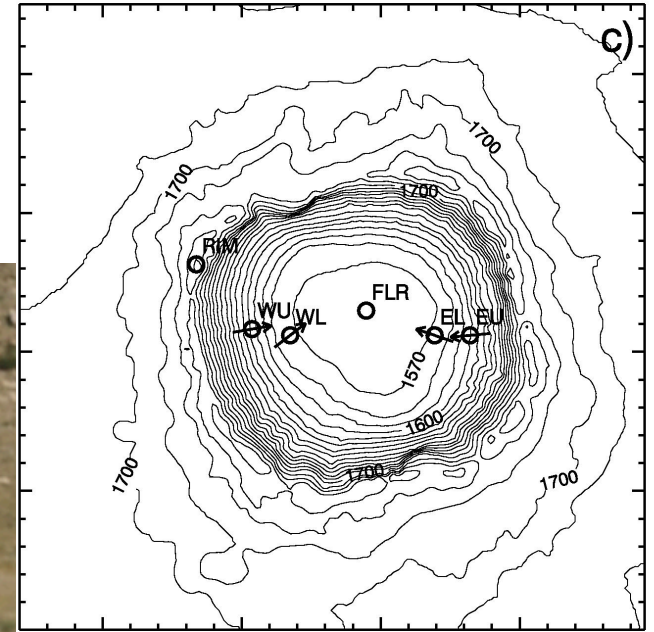
## 2 Precision Filter Radiometers (PFR):

Spectral Direct Radiation

- PFR-N: (368, 412, 500, 863 nm)
- PFR-UV: (305, 311, 318, 332 nm)



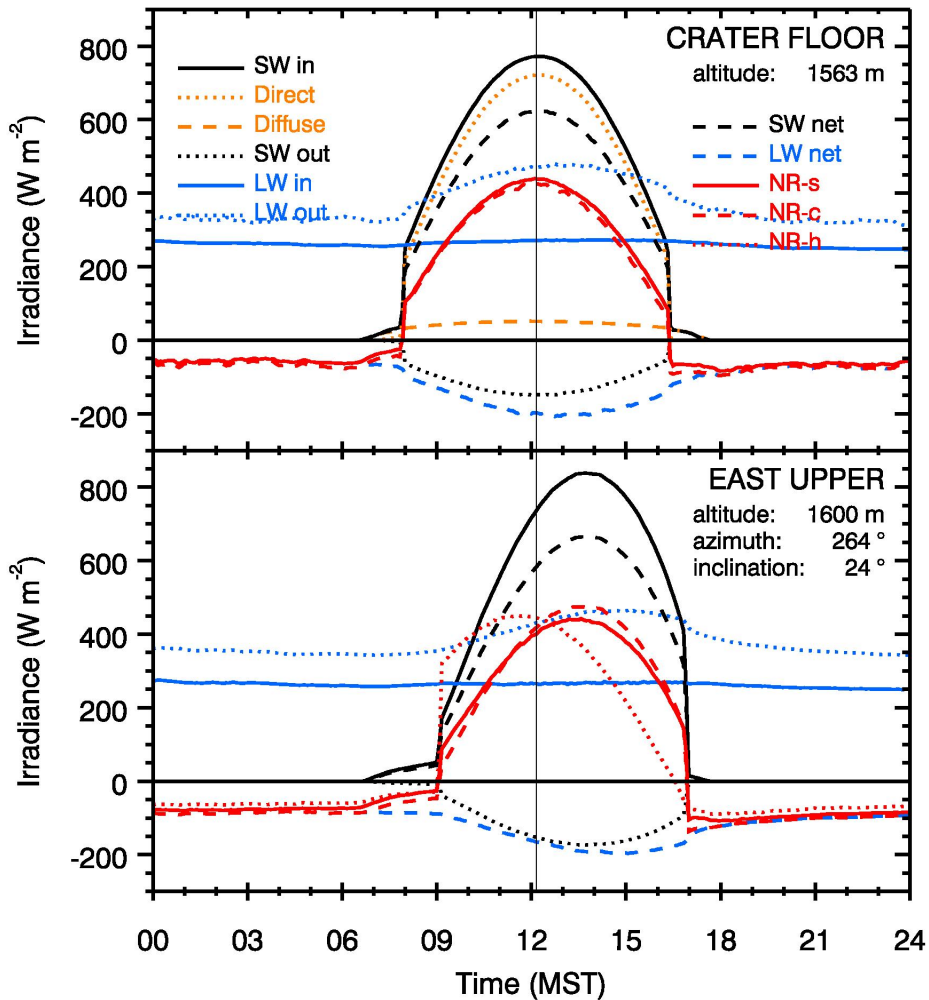
# Radiation Balance Measurements during METCRAX 2006





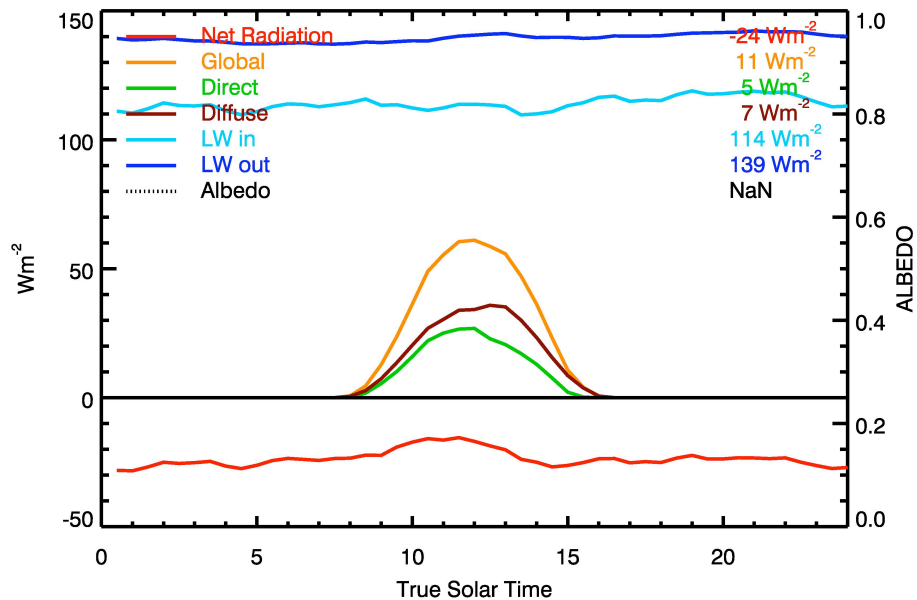
# Meteor Crater, Arizona

21 October 2006

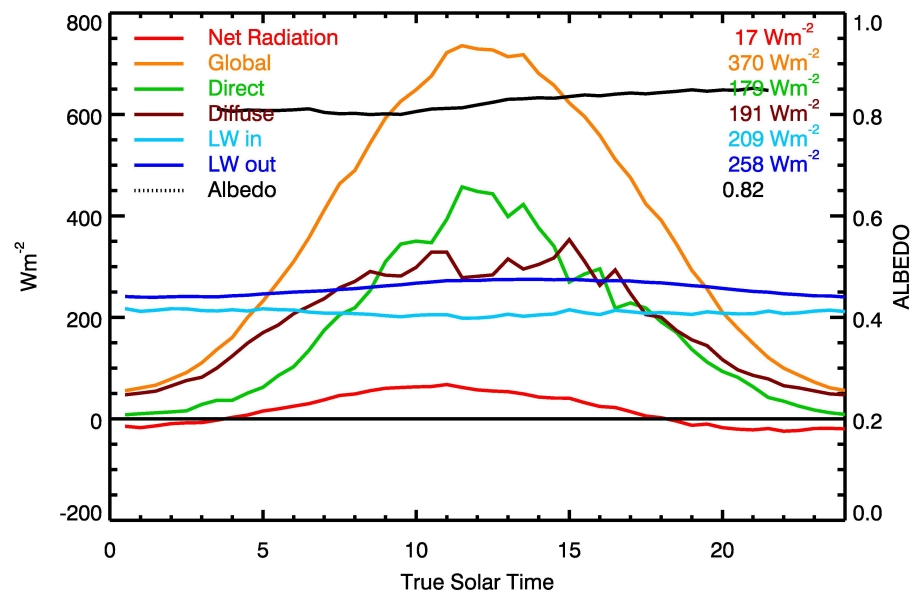


# Summit, Greenland (3200 m)

FEB 2002



JUN 2002



# Radiation measurements during MATERHORN

EFS-Playa



EFS-Sagebrush



EFS-Slope

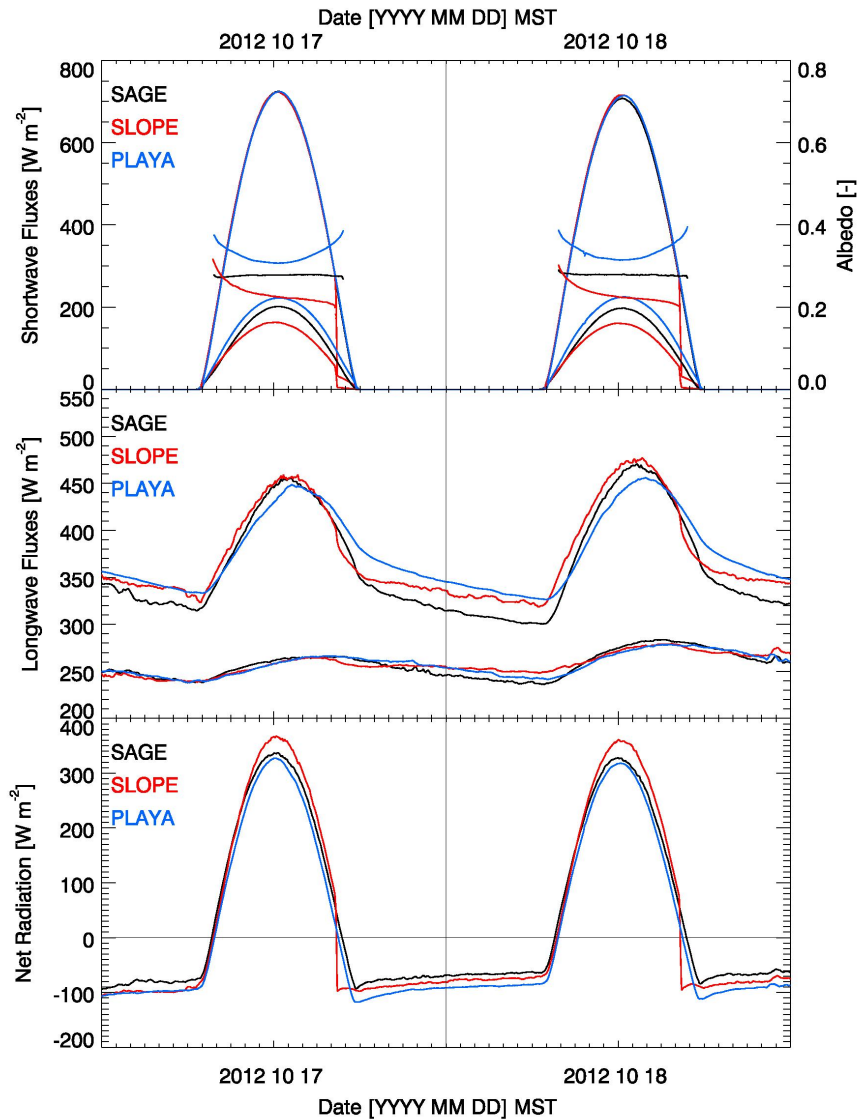


Detailed View – 4 components



# Radiation Balance at EFS-Sites

$$NR = SW^{\downarrow} - SW^{\uparrow} + LW^{\downarrow} - LW^{\uparrow}$$



- Same shortwave energy input  $SW^{\downarrow}$
- Albedo controls  $SW^{\uparrow}$
- Same daytime NR at EFS-Sage and EFS-Playa
- Differences in  $SW^*$  are compensated by differences in  $LW^{\uparrow}$
- NR differences (Playa – Sage) are larger at nighttime, pointing to differences in soil thermal properties.



Wer misst, misst Mist!

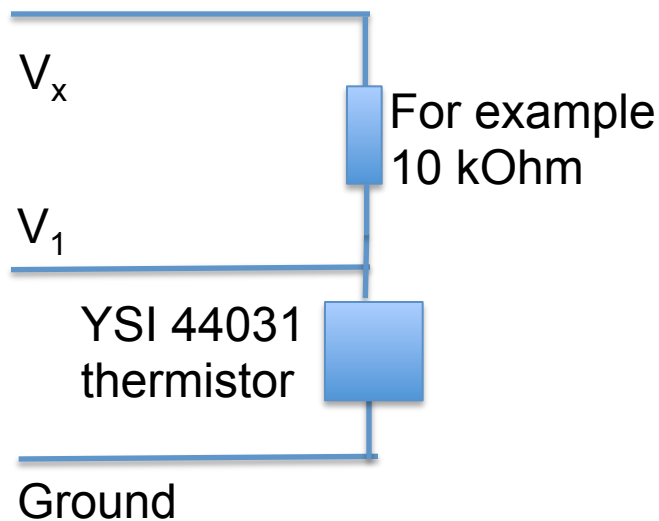


# Additional discussion points ...

- Build of instruments
- Thermistor measurements (YSI44031) – voltage drop across thermistor using known resistor (accuracy ?)
- Choice of voltage range / resistor / resolution / excitation voltage
- Measurements in different climates / elevations?



## Considerations – LW measurements



What temperature range do I expect? – What resistance will that mean?

What voltage range do I choose?

What is my best excitation voltage?

What will my resolution ( $W\ m^{-2}$ ) be (cc ~  $10\ \mu V / (W\ m^{-2})$ )?

Temperature [ °C ]	Resistance [ °F ]	Resistance [ Ohm ]	Temperature [ °C ]	Resistance [ °F ]	Resistance [ Ohm ]	Temperature [ °C ]	Resistance [ °F ]	Resistance [ Ohm ]
-30	-22.0	135,200	0	32.0	29,490	30	86.0	8,194
-29	-20.2	127,900	1	33.8	28,150	31	87.8	7,880
-28	-18.4	121,100	2	35.6	26,890	32	89.6	7,579
-27	-16.6	114,600	3	37.4	25,690	33	91.4	7,291
-26	-14.8	108,600	4	39.2	24,550	34	93.2	7,016
-25	-13.0	102,900	5	41.0	23,460	35	95.0	6,752
-24	-11.2	97,490	6	42.8	22,430	36	96.8	6,500
-23	-9.4	92,430	7	44.6	21,450	37	98.6	6,258
-22	-7.6	87,660	8	46.4	20,520	38	100.4	6,026
-21	-5.8	83,160	9	48.2	19,630	39	102.2	5,805
-20	-4.0	78,910	10	50.0	18,790	40	104.0	5,592
-19	-2.2	74,910	11	51.8	17,980	41	105.8	5,389
-18	-0.4	71,130	12	53.6	17,220	42	107.6	5,193
-17	1.4	67,570	13	55.4	16,490	43	109.4	5,006
-16	3.2	64,200	14	57.2	15,790	44	111.2	4,827
-15	5.0	61,020	15	59.0	15,130	45	113.0	4,655
-14	6.8	58,010	16	60.8	14,500	46	114.8	4,489
-13	8.6	55,170	17	62.6	13,900	47	116.6	4,331
-12	10.4	52,480	18	64.4	13,330	48	118.4	4,179
-11	12.2	49,940	19	66.2	12,790	49	120.2	4,033
-10	14.0	47,540	20	68.0	12,260	50	122.0	3,893
-9	15.8	45,270	21	69.8	11,770	51	123.8	3,758
-8	17.6	43,110	22	71.6	11,290	52	125.6	3,629
-7	19.4	41,070	23	73.4	10,840	53	127.4	3,504
-6	21.2	39,140	24	75.2	10,410	54	129.2	3,385
-5	23.0	37,310	25	77.0	10,000	55	131.0	3,270
-4	24.8	35,570	26	78.8	9,605	56	132.8	3,160
-3	26.6	33,930	27	80.6	9,227	57	134.6	3,054
-2	28.4	32,370	28	82.4	8,867	58	136.4	2,952
-1	30.2	30,890	29	84.2	8,523	59	138.2	2,854

From Kipp & Zonen – CGR4 manual

# Credits & Acknowledgements

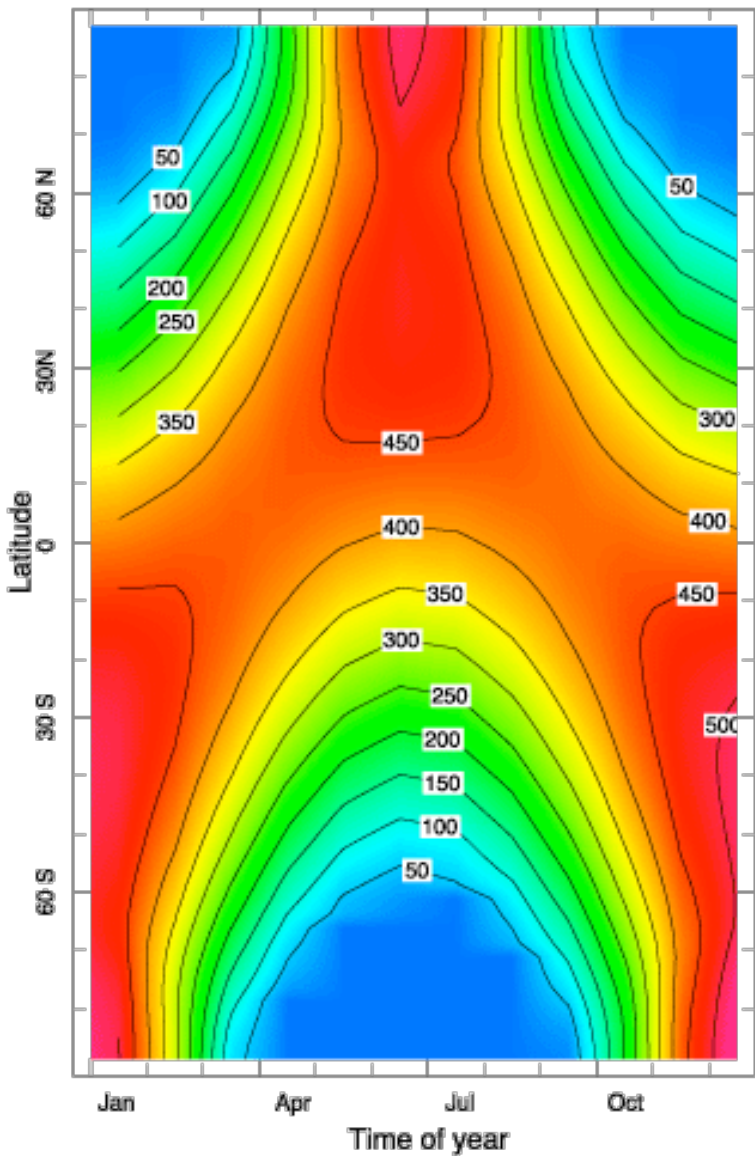
Lecture notes of Prof. Claus Froehlich, Davos: <ftp://ftp.pmodwrc.ch/pub/Claus/Vorlesung2009/>

Notes on ETH Feldkurs Rietholtzbach by Reto Stoeckli

Kipp & Zonen: <http://www.kippzonen.com/?downloadcategory/551/Pyranometers.aspx>

Wikipedia articles

# Latitude-Time Distribution of Incoming Solar Radiation at the Top of the Atmosphere



Based on ERBE data. Units are  $W/m^2$

