

Radiation measurements

Motivation (Energy Balance)

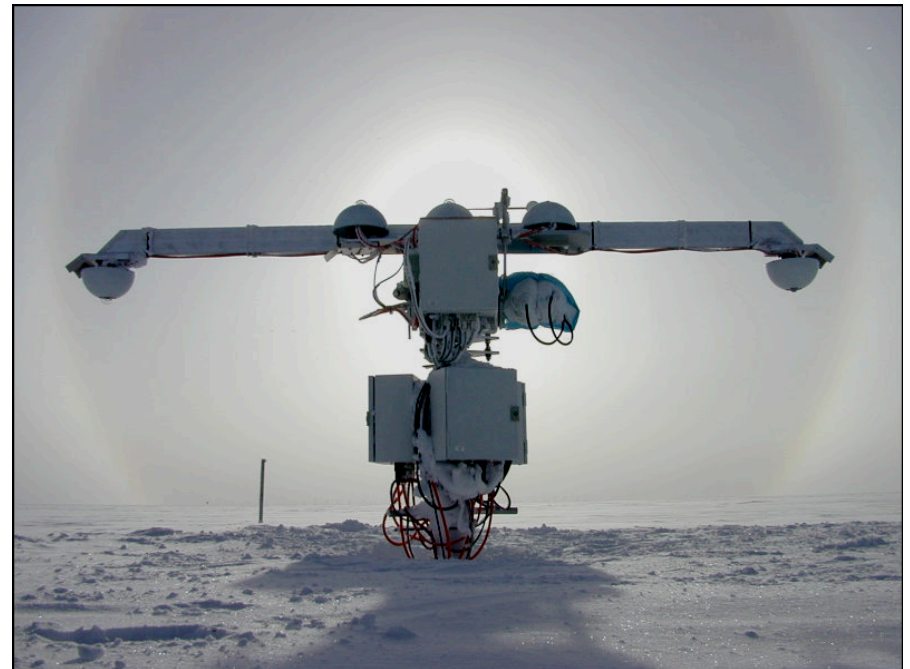
Background

Radiation Quantities & Terms in Radiation Budget

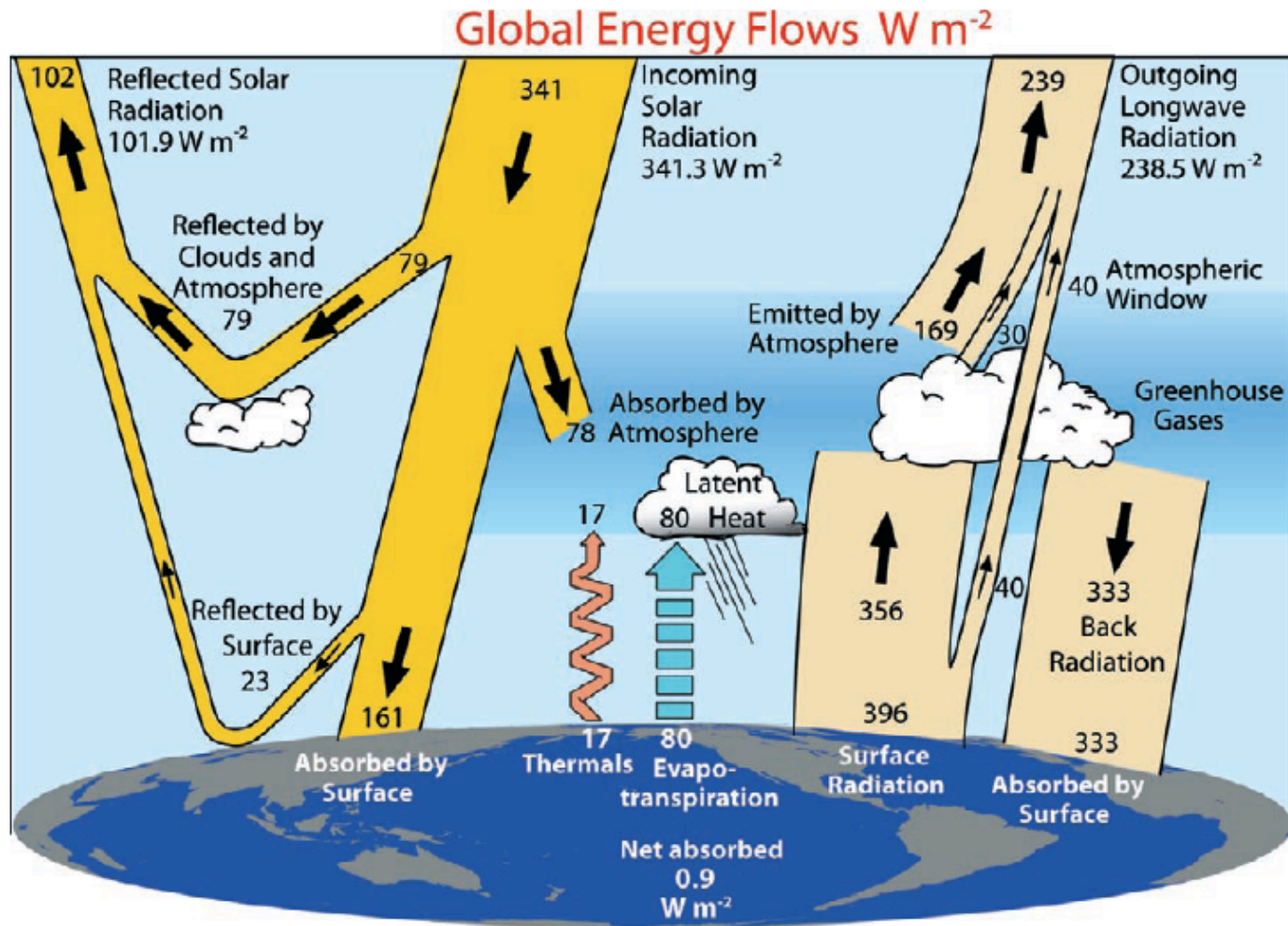
Instrumentation & Measurement Principles

Radiation Balance in different climates

Sebastian W. Hoch
485 INSCC

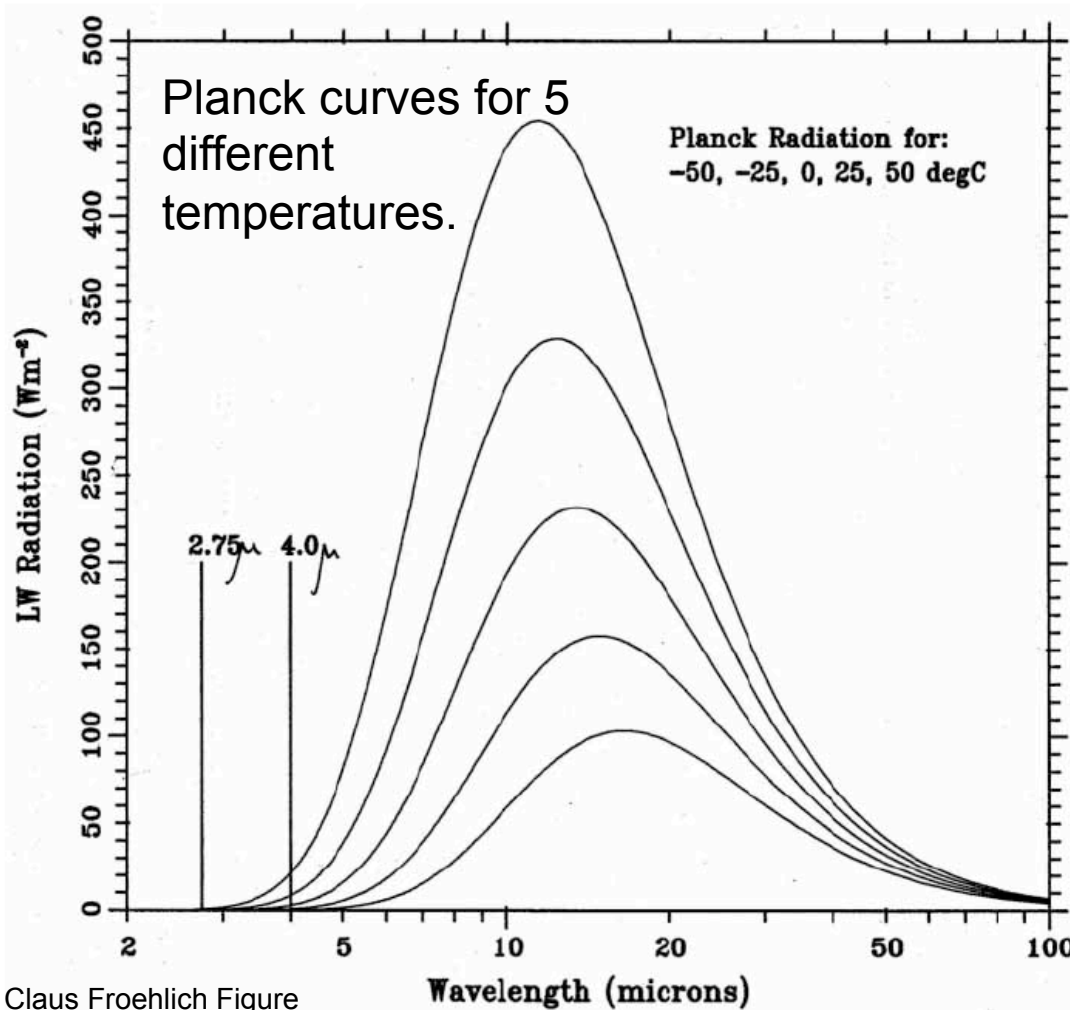


Radiation and the Energy Budget



Trenberth et al. 2009 BAMS

Why **Shortwave** Radiation and **Longwave** Radiation?



Planck Function

$$B_{\lambda}(T) = \frac{2hc^2}{\lambda^5 (e^{hc/k\lambda T} - 1)}$$

$$h = 6.626\ 068\ 96(33) \times 10^{-34} \text{ J s}$$

c : speed of light

λ : wavelength

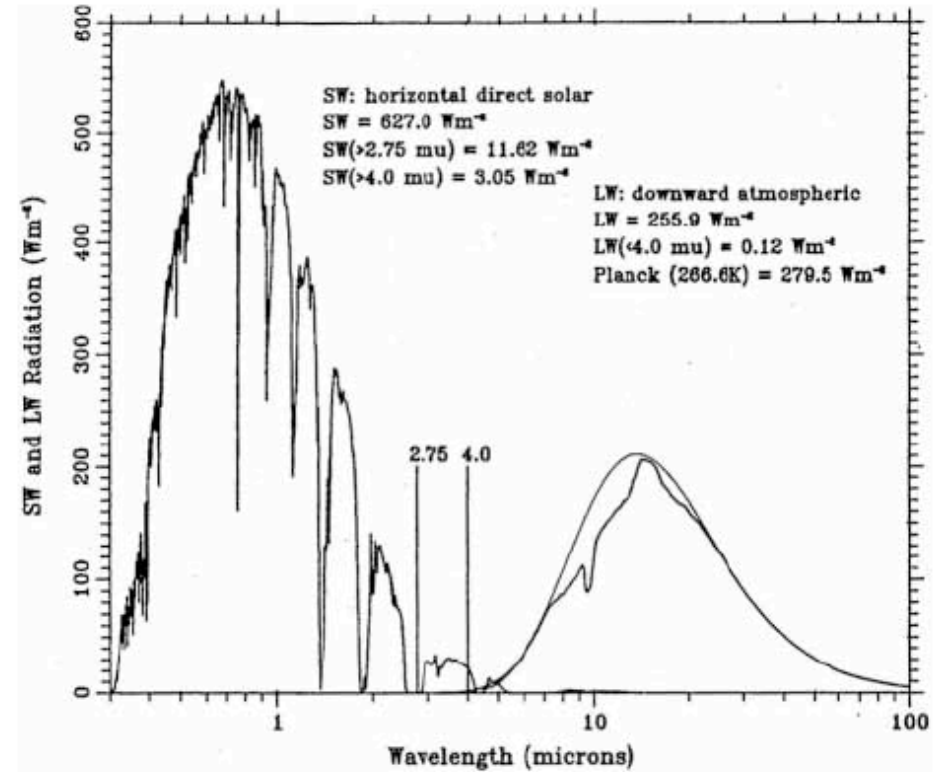
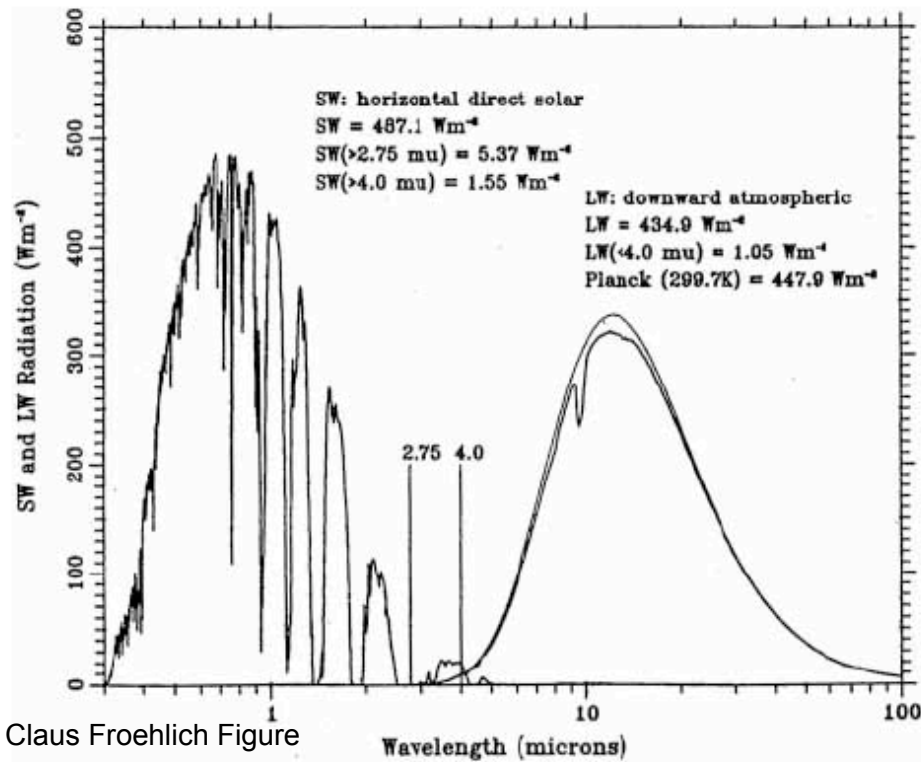
Everything emits radiation – depending on the temperature!

Wien's Displacement Law

$$\lambda_{\max} = \frac{a}{T} \text{ mit } a = 2.89776 \times 10^{-3} \text{ m K}$$

Longwave or Terrestrial or Infrared Radiation and

Shortwave or Solar Radiation



Tropical atmosphere

Midlatitude Summer, 1600 m (ASL)

Other quantities defined by spectral range:

- UV Radiation (A, B, C)
- PAR: Photosynthetically Active Radiation; 400 - 700 nm

There's an overlap at times ...

Radiation Quantities

Quantity	Symbol	SI unit	Abbr.	Notes
Radiant energy	Q	joule	J	energy
Radiant flux	Φ	watt	W	radiant energy per unit time, also called <i>radiant power</i>
Radiant intensity	I	watt per steradian	$\text{W}\cdot\text{sr}^{-1}$	power per unit solid angle
Radiance	L	watt per steradian per square metre	$\text{W}\cdot\text{sr}^{-1}\cdot\text{m}^{-2}$	power per unit solid angle per unit <i>projected</i> source area. called <i>intensity</i> in some other fields of study.
Irradiance	E, I	watt per square metre	$\text{W}\cdot\text{m}^{-2}$	power incident on a surface. sometimes confusingly called "intensity".
Radiant exitance / Radiant emittance	M	watt per square metre	$\text{W}\cdot\text{m}^{-2}$	power emitted from a surface.
Radiosity	J or J_λ	watt per square metre	$\text{W}\cdot\text{m}^{-2}$	emitted plus reflected power leaving a surface
Spectral radiance	L_λ or L_ν	watt per steradian per metre ³ or watt per steradian per square metre per hertz	$\text{W}\cdot\text{sr}^{-1}\cdot\text{m}^{-3}$ or $\text{W}\cdot\text{sr}^{-1}\cdot\text{m}^{-2}\cdot\text{Hz}^{-1}$	commonly measured in $\text{W}\cdot\text{sr}^{-1}\cdot\text{m}^{-2}\cdot\text{nm}^{-1}$
Spectral irradiance	E_λ or E_ν	watt per metre ³ or watt per square metre per hertz	$\text{W}\cdot\text{m}^{-3}$ or $\text{W}\cdot\text{m}^{-2}\cdot\text{Hz}^{-1}$	commonly measured in $\text{W}\cdot\text{m}^{-2}\cdot\text{nm}^{-1}$

The Radiation Balance – the terms (Irradiances $W\ m^{-2}$)

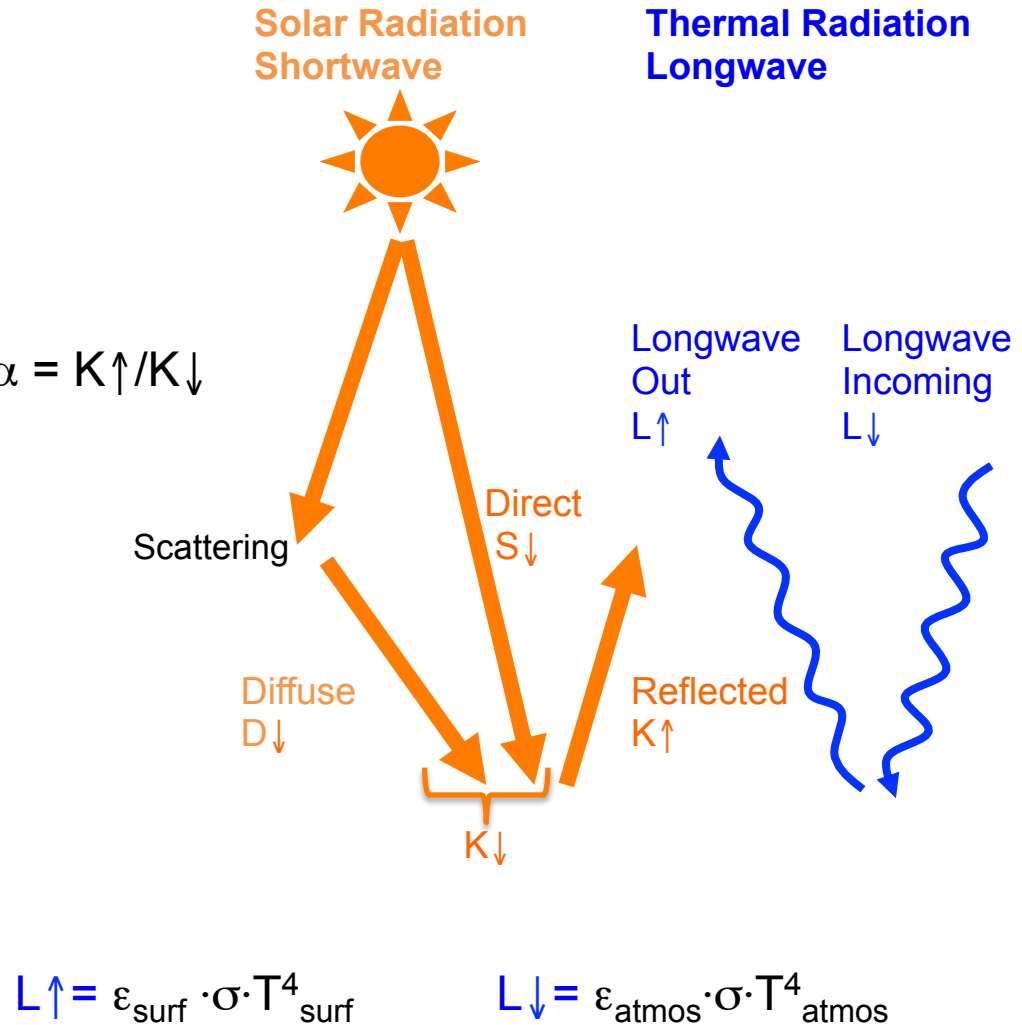
Direct Solar Radiation $S\downarrow$
 Diffuse (Solar) Radiation $D\downarrow$
 Global Radiation ($K\downarrow$, G_I) = $S\downarrow + D\downarrow$
 Shortwave Reflected Radiation $K\uparrow$
 Shortwave Net Radiation K^*

Longwave Incoming $L\downarrow$
 Longwave Outgoing Radiation $L\uparrow$
 Longwave Net Radiation L^*

Net Radiation Q^*

$$\begin{aligned}
 Q^* &= K^* + L^* \\
 &= K\downarrow - K\uparrow + L\downarrow - L\uparrow \\
 &= (1 - \alpha) * K\downarrow + L^*
 \end{aligned}$$

$$\text{Albedo } \alpha = K\uparrow / K\downarrow$$



$$L\uparrow = \epsilon_{\text{surf}} \cdot \sigma \cdot T_{\text{surf}}^4$$

$$L\downarrow = \epsilon_{\text{atmos}} \cdot \sigma \cdot T_{\text{atmos}}^4$$

Stefan-Boltzmann Constant
 $\sigma: 5.67 \cdot 10^{-8} \text{ J s}^{-1} \text{ m}^{-2} \text{ K}^{-4}$

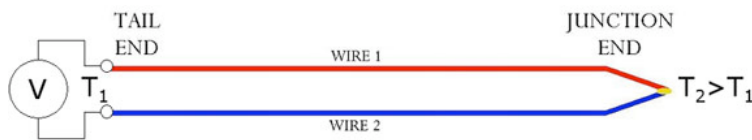
Measurement Principle

Thermopile

- converts thermal energy into electrical energy
- composed of thermocouples (usually in series)
- output voltage proportional to a local temperature difference
- range of tens or hundreds of millivolts.

Thermocouple

- temperature measurement based on the *Seebeck Effect*: a result of a difference in *thermoelectric power* of two materials



$$Emf = \int_{T_1}^{T_2} S_{12} \cdot dT = \int_{T_1}^{T_2} (S_1 - S_2) \cdot dT$$

- Emf is the Electro-Motive Force or Voltage; T_1 and T_2 : Temperatures of reference (T_1) and measuring end (T_2)
- S_{12} , S_1 , S_2 : **Seebeck coefficients** of the thermocouple and thermo-elements
- null voltage:
 - same materials
 - no temperature difference

Radiation observations in Climate Science - Instrumentation

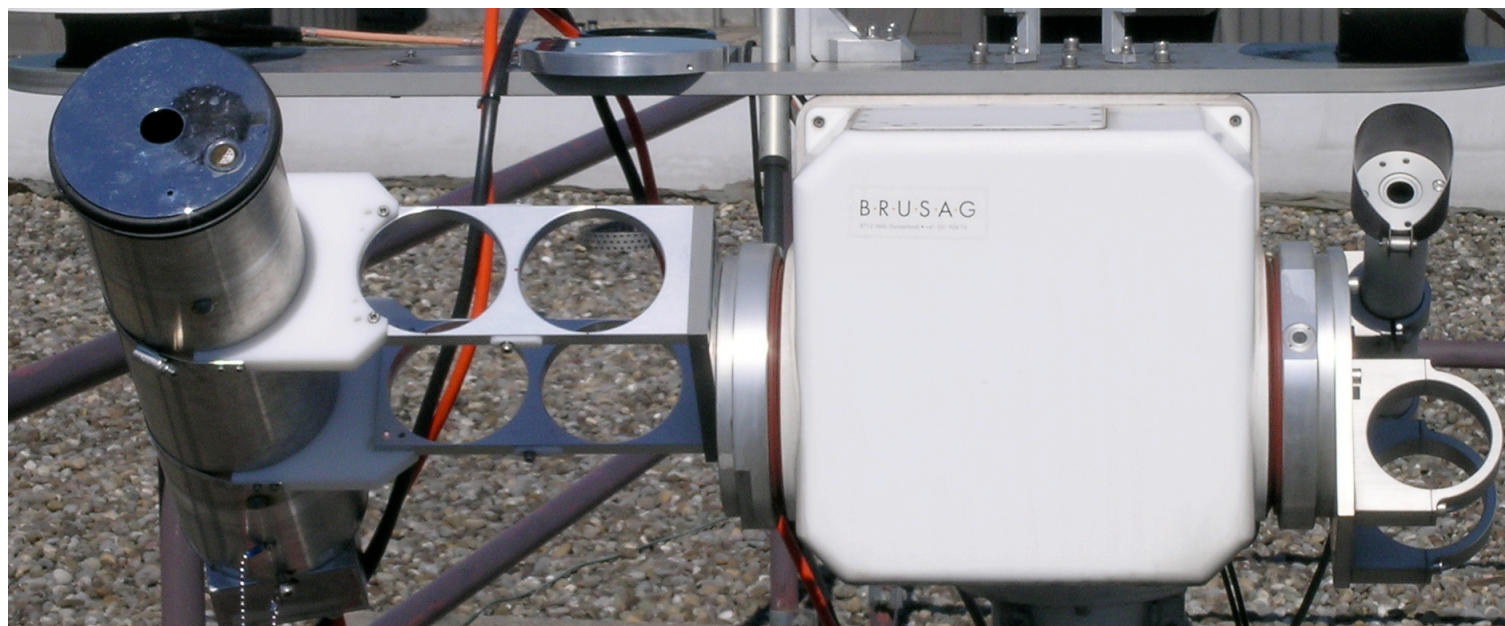
1. Pyrheliometer

Direct Solar Radiation
World Standard Instruments
(Compensation Type / Thermopile)
Open / with window ...

pyro-, pyr- +
(Greek: fire, burn; heat,
produced by heating; and
sometimes "fever")

"ἥλιος (Helios) is derived
from the noun ἥλιος,
"sun" in ancient Greek

PMO-6



Kipp & Zonen
CH1

PMO-6 Absolute Cavity Radiometer

$$S = k * (P_{\text{closed}} - P_{\text{open}})$$



pmod/wrc

Other System:
Eppley Hickley-Frieden (HF)

Thermopile Pyrheliometer (NIP / CH1)

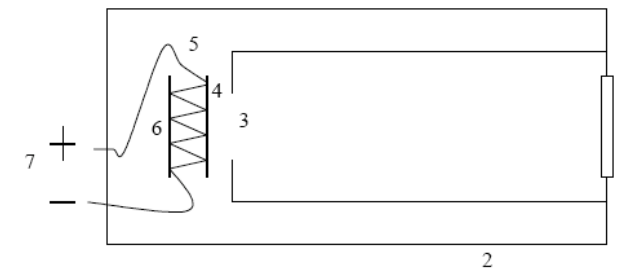
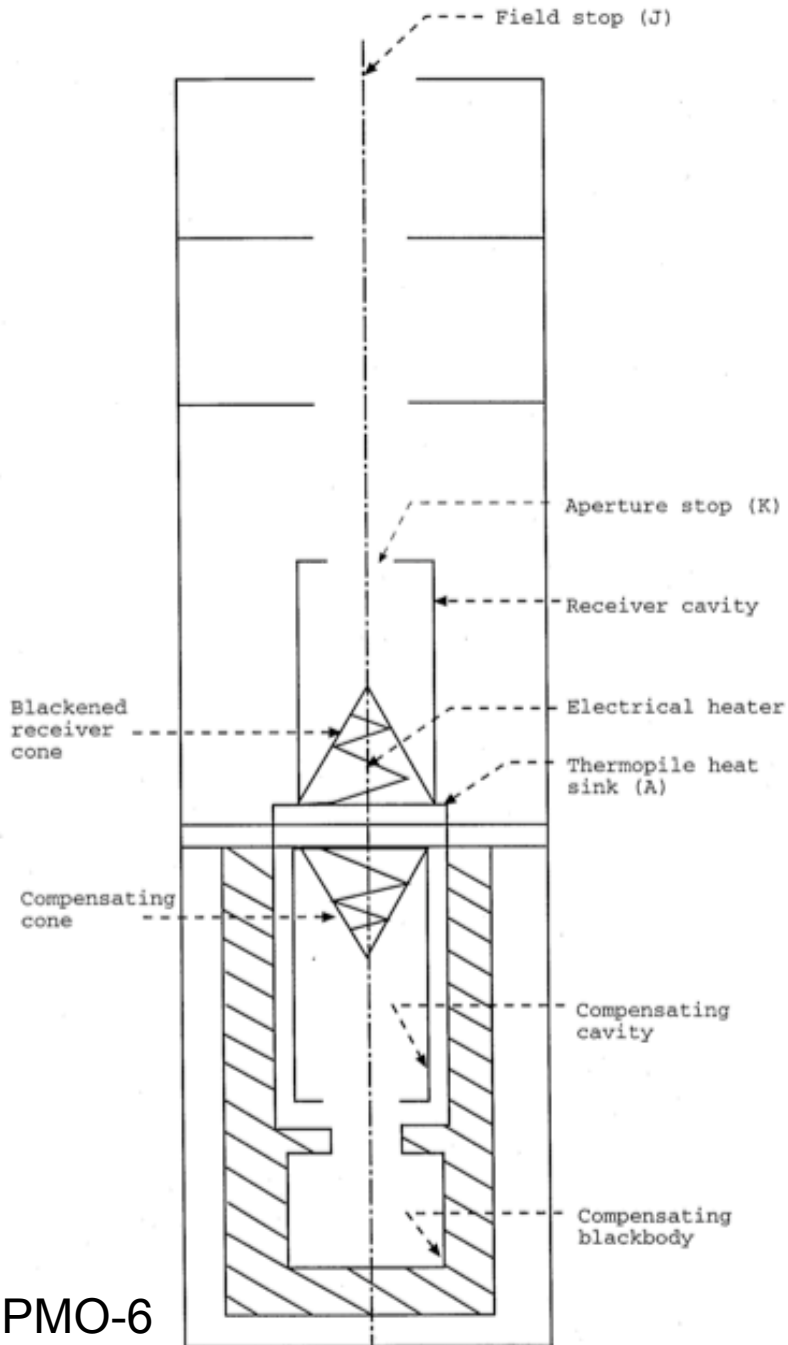


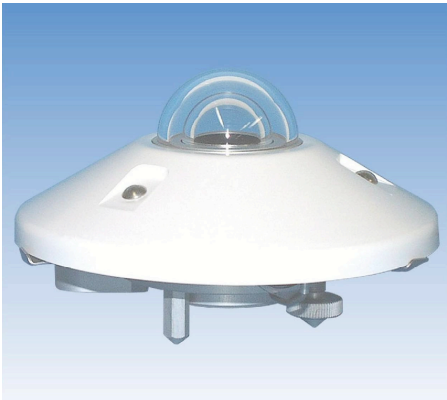
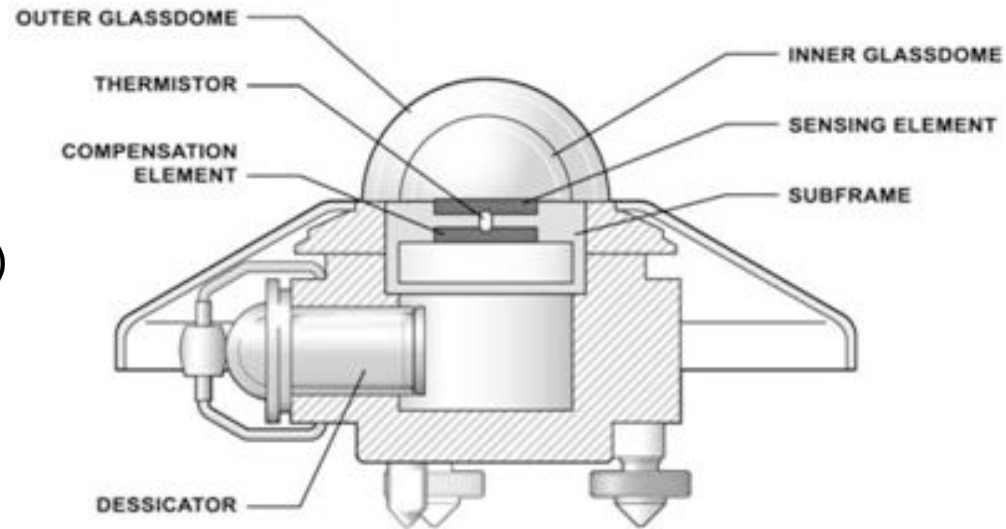
Fig. 1. Pyrheliometer schematic showing entrance window (1), thermal shield (2), detector aperture (3), light absorber (4), thermopile (5), heat sink (6), and thermopile output (7).



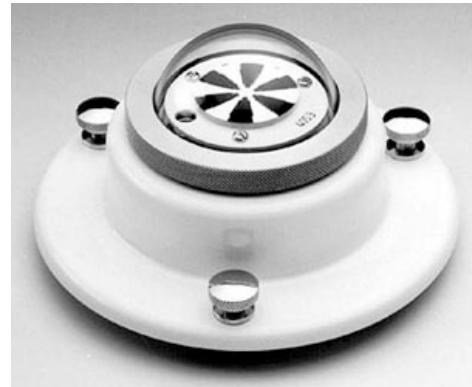
PMO-6

2. Pyranometer

- Global Radiation
- Shortwave Reflected Radiation
- Diffuse Radiation (in conjunction with a shading disk or shadow-band)
- Glass or quartz dome



Standard



Black & White Type



Photodiode Type

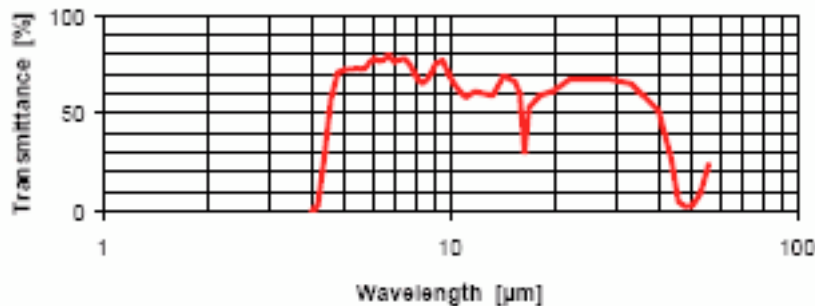
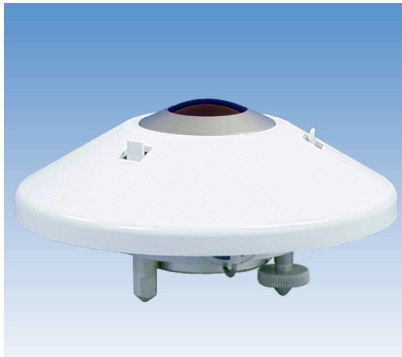
Shading – Shadowbands and Shading disks



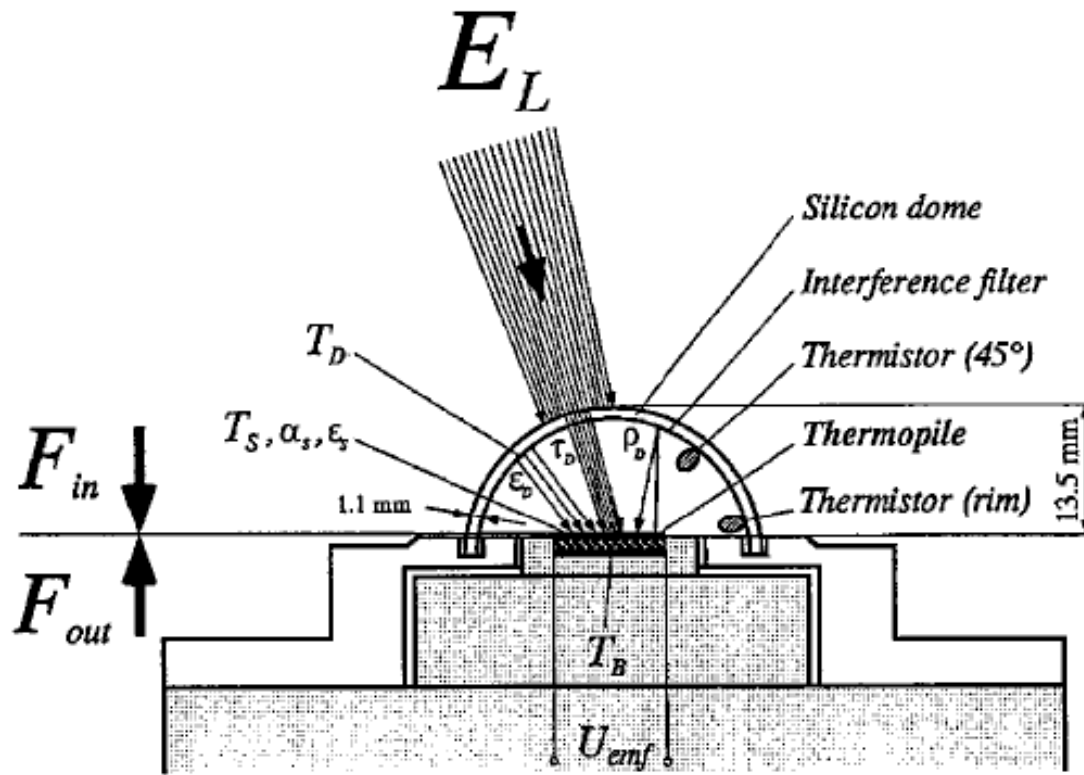
3. Pyrgeometer

Longwave Radiation
Thermopile, Silicon (Si) dome

geo-, ge- +
(Greek: earth, land,
soil; world)



Si-Dome and interference filters



Schematic view of Eppley PIR (Philipona et al. 1995)

Pyrgeometer Formula:

$$E_L = \frac{U_{emf}}{C} \left(\underbrace{1}_{LWin_a} + \underbrace{k_1 \sigma T_B^3}_{LWin_b} \right) + \underbrace{k_2 \sigma T_B^4}_{LWin_b} - \underbrace{k_3 \sigma (T_D^4 - T_B^4)}_{LWin_c}.$$

We neglect k_1 , set k_2 to 1.0, and k_3 to a mean value of 3.5.

4. Pyrradiometer

- “All-wave” Radiation
- Thermopile measurements
- Polyethylene Dome
- Double domes: Net-Radiometer
- “Wind Speed Error”



Different response
to short- and
longwave fluxes!

Birds like to destroy them, too...

5. Heliograph / Sunshine Duration Sensor



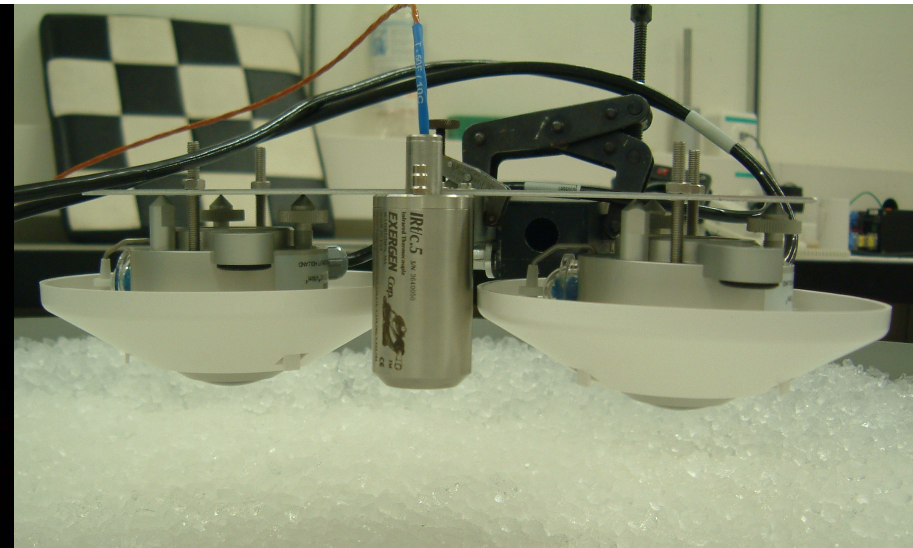
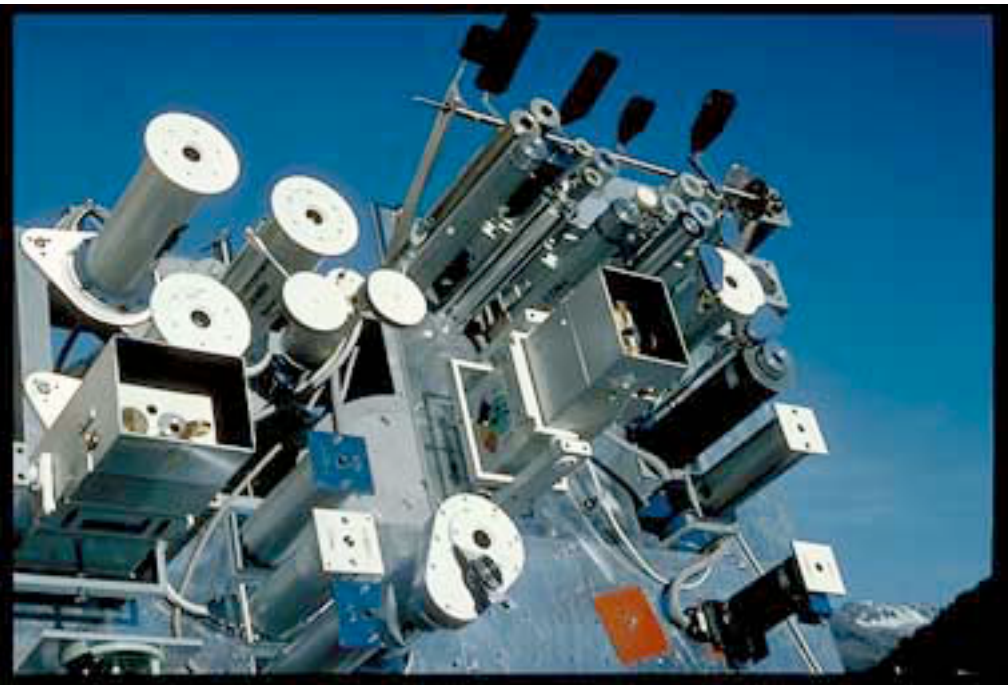
Campbell-Stokes Sunshine Recorder

“Sunshine”: Flux > 120 Wm^{-2}



One end of an optical fiber revolves around the sun axis. The opening angle is limited by an optical diaphragm. At the other end, a photovoltaic detector receives the light pulse when the fiber window meets the sun. The detected signal is compared to a threshold. A pulse is generated when the radiation intensity exceeds 120 W/m^2 .

Calibrations and Errors



WSG (World Standard Group)
Davos, Switzerland

- Absolute Calibration Error
(Comparison to World Standard)

Spectral Response Errors

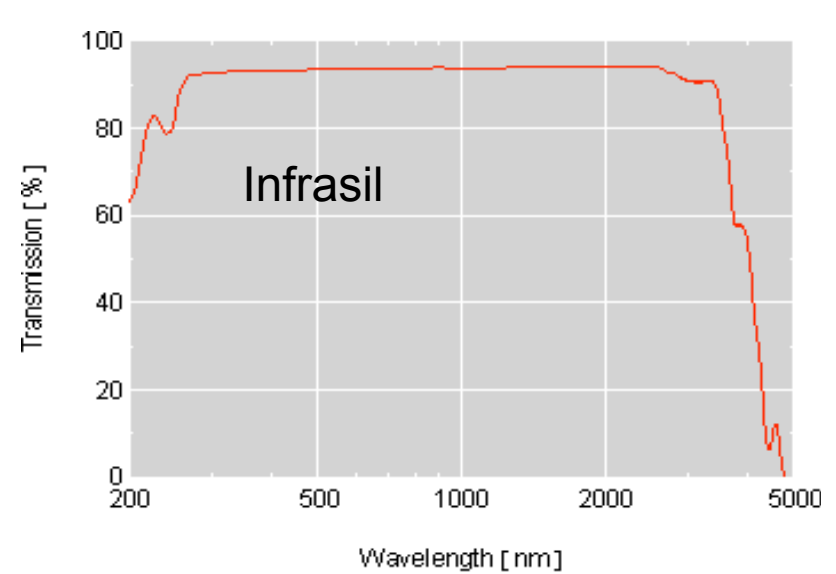
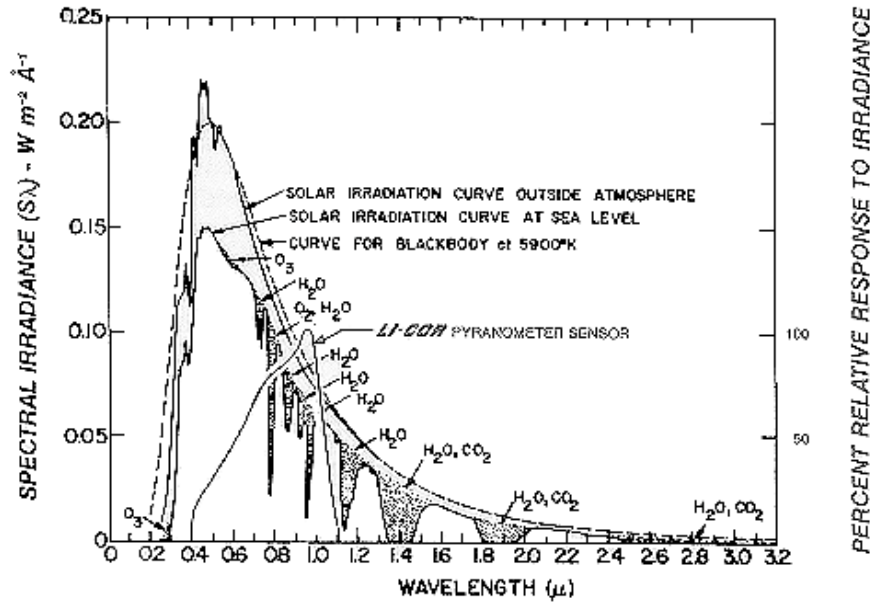
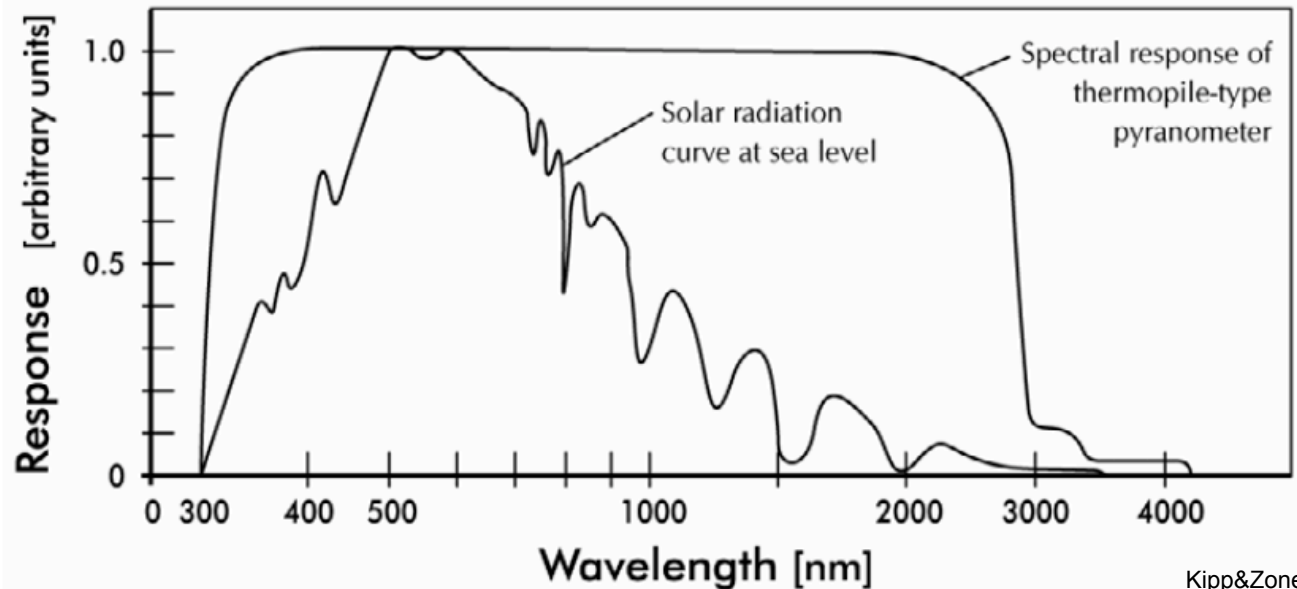
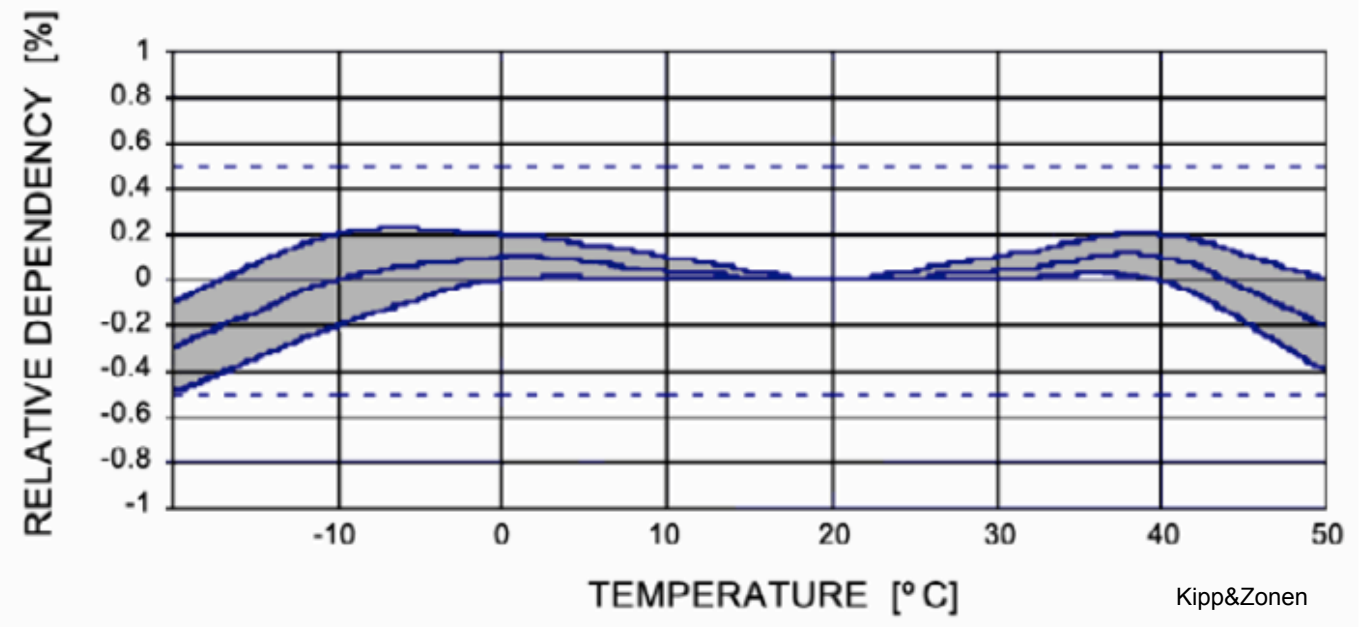


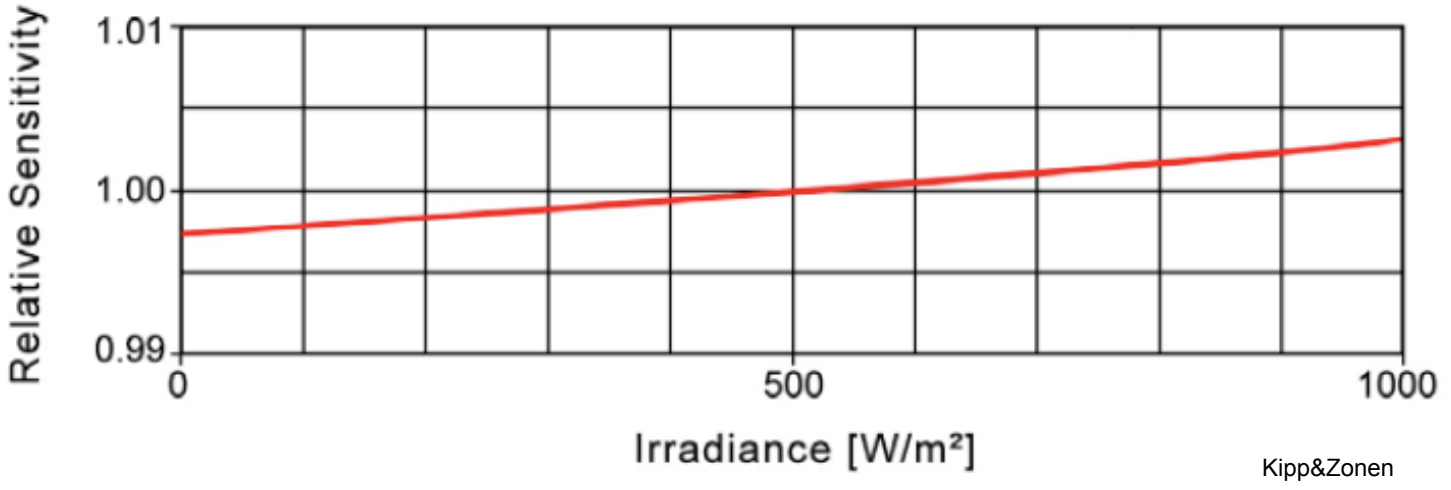
Figure 4. The LI-200SA Pyranometer spectral response is illustrated along with the energy distribution in the solar spectrum (8).



- Temperature Dependency

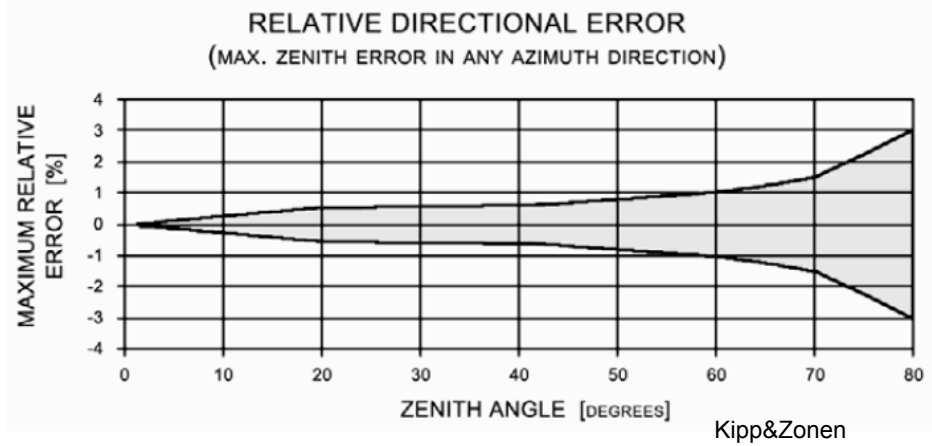
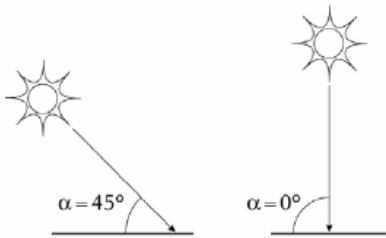


- Linearity

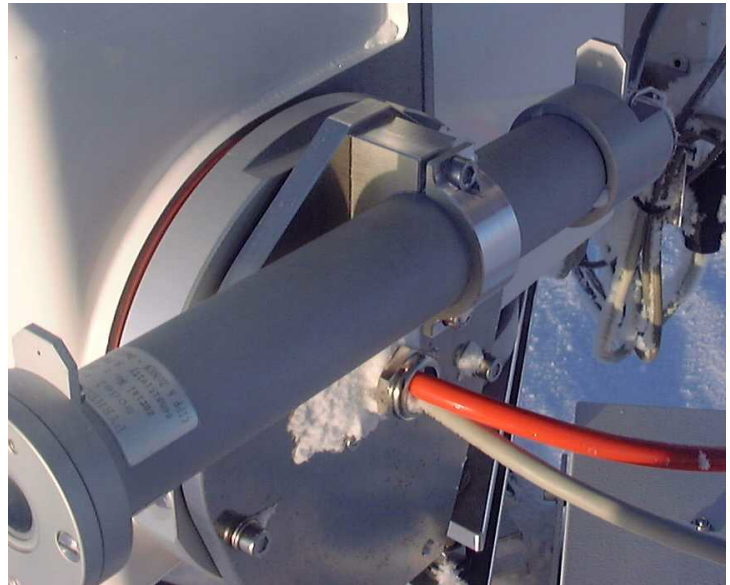


Geometric Errors:

- Cosine Response Error
(low vs high incident radiation)
- Azimuth error (sensor geometry)



- Hysteresis
- Response Time Error
- Long Term Stability (aging of thermopile / paint / resistors / etc)

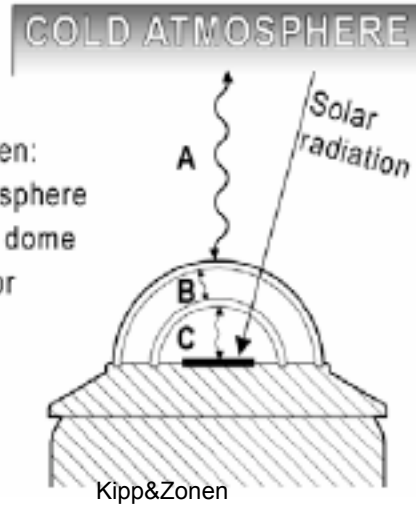


Pointing error



Condensation

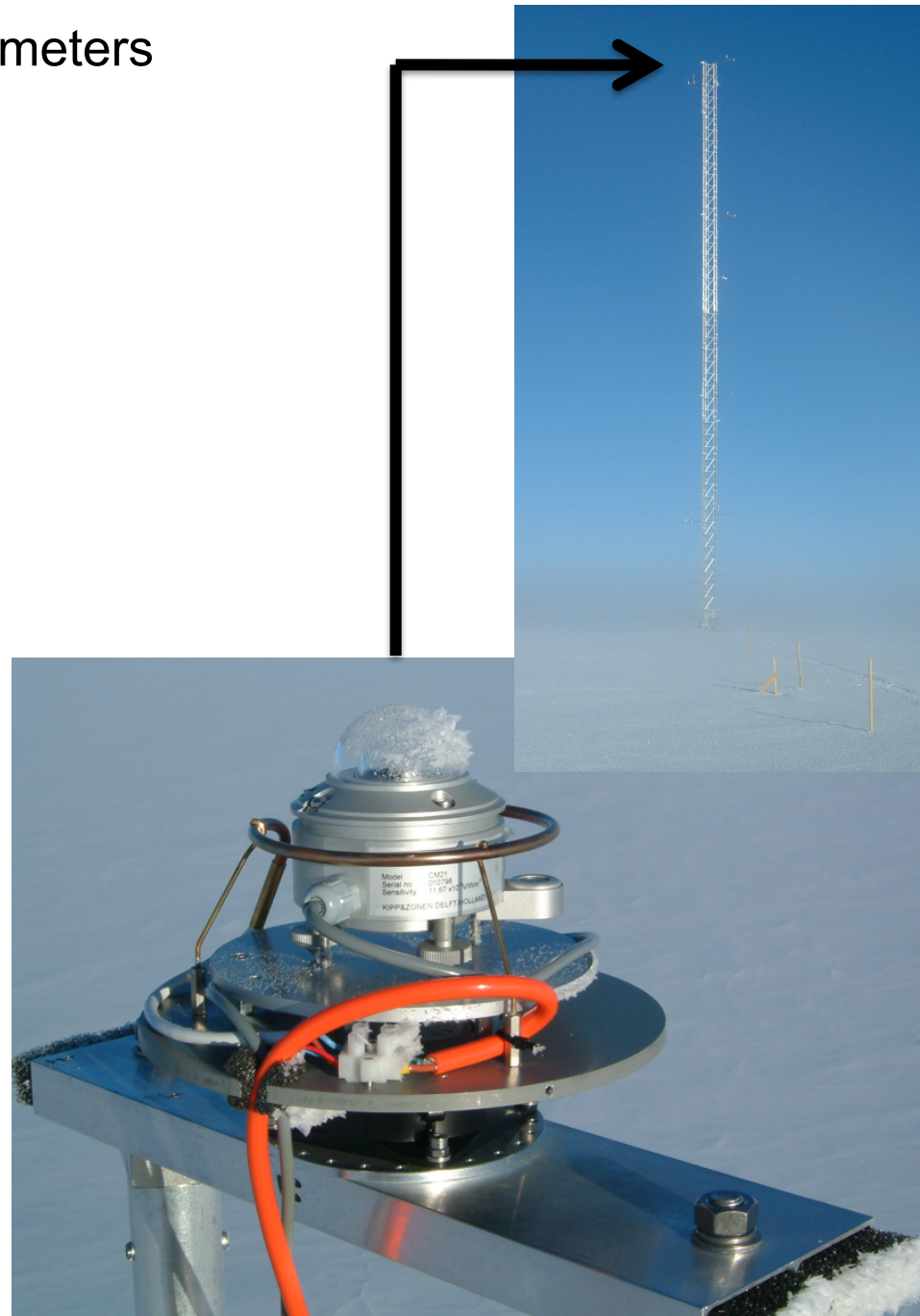
Negative-Night-Time-Offset of Pyranometers



Ventilation and Heating!

“Wind-Correction” of Pyrradiometers and Net-Radiometers

Dome material (polyethylene, lupolene) heats up. Ventilation reduces the heating effect.



Environmental impacts



Radiation balance measurement at Summit, Greenland

2 Pyrgeometers (Eppley PIR):

- Longwave Incoming
- Longwave Outgoing

Shadow Discs

3 Pyranometers (Kipp&Zonen CM21/11):

- Diffuse Radiation
- Global Radiation
- Shortwave Reflected

Absolute Cavity Radiometer (PMO6)
Reference Instrument

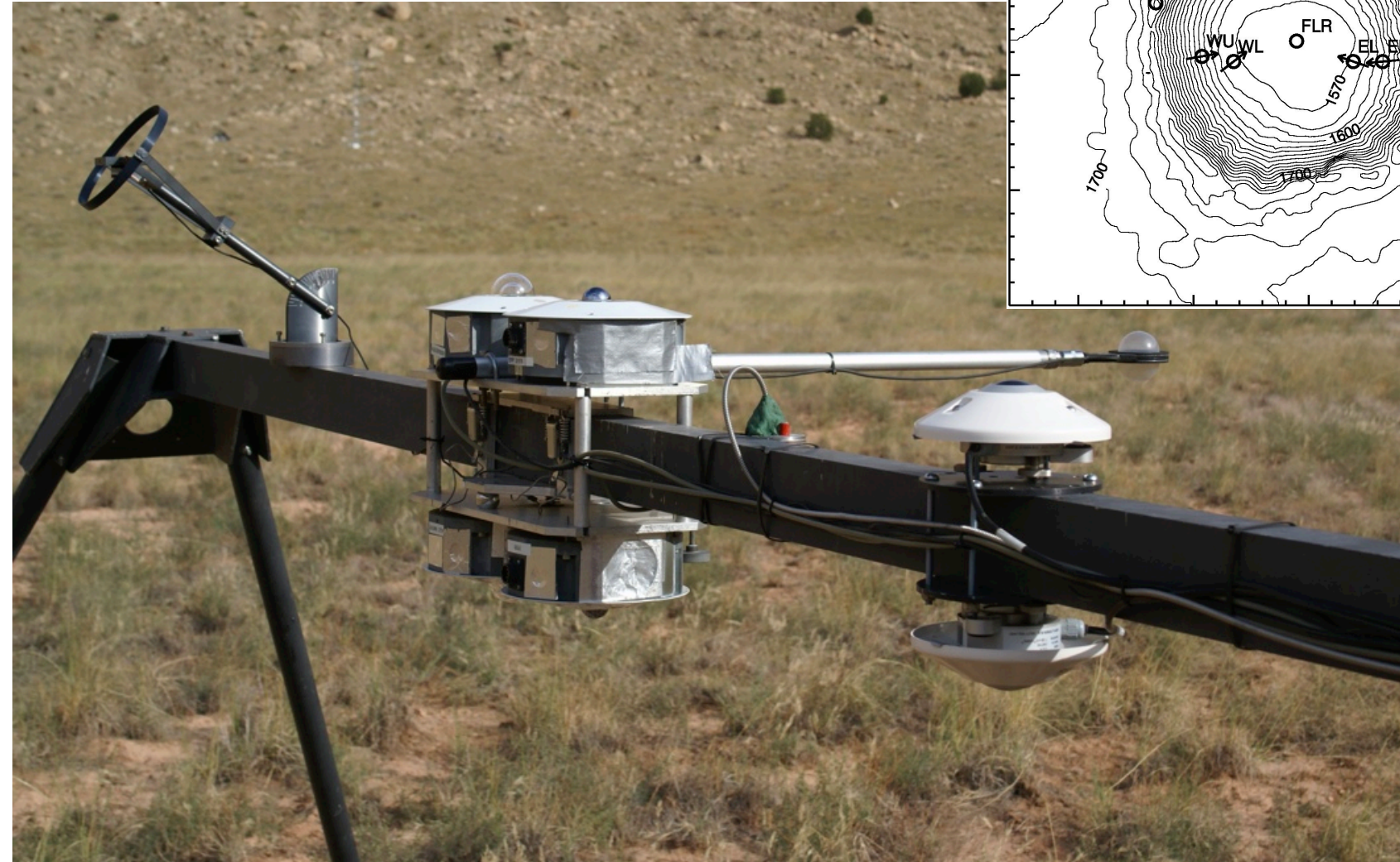
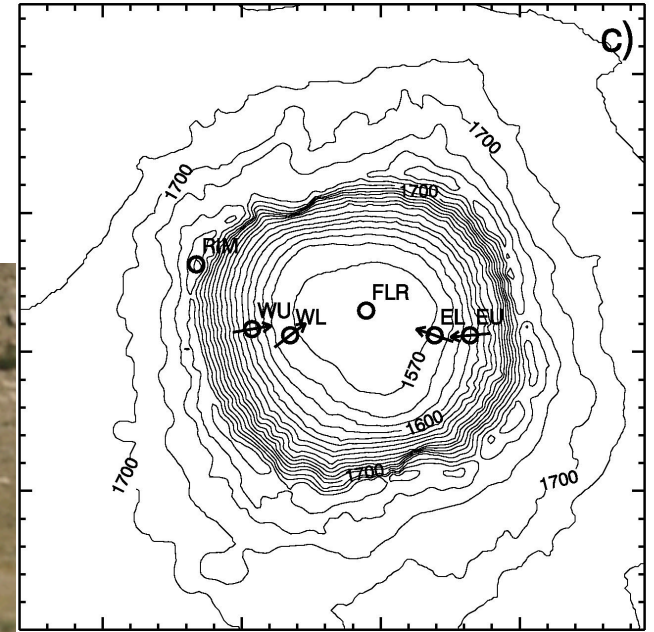
Pyrheliometer (Kipp&Zonen CH1):
• Direct Radiation

2 Precision Filter Radiometers (PFR):
Spectral Direct Radiation

- PFR-N: (368, 412, 500, 863 nm)
- PFR-UV: (305, 311, 318, 332 nm)

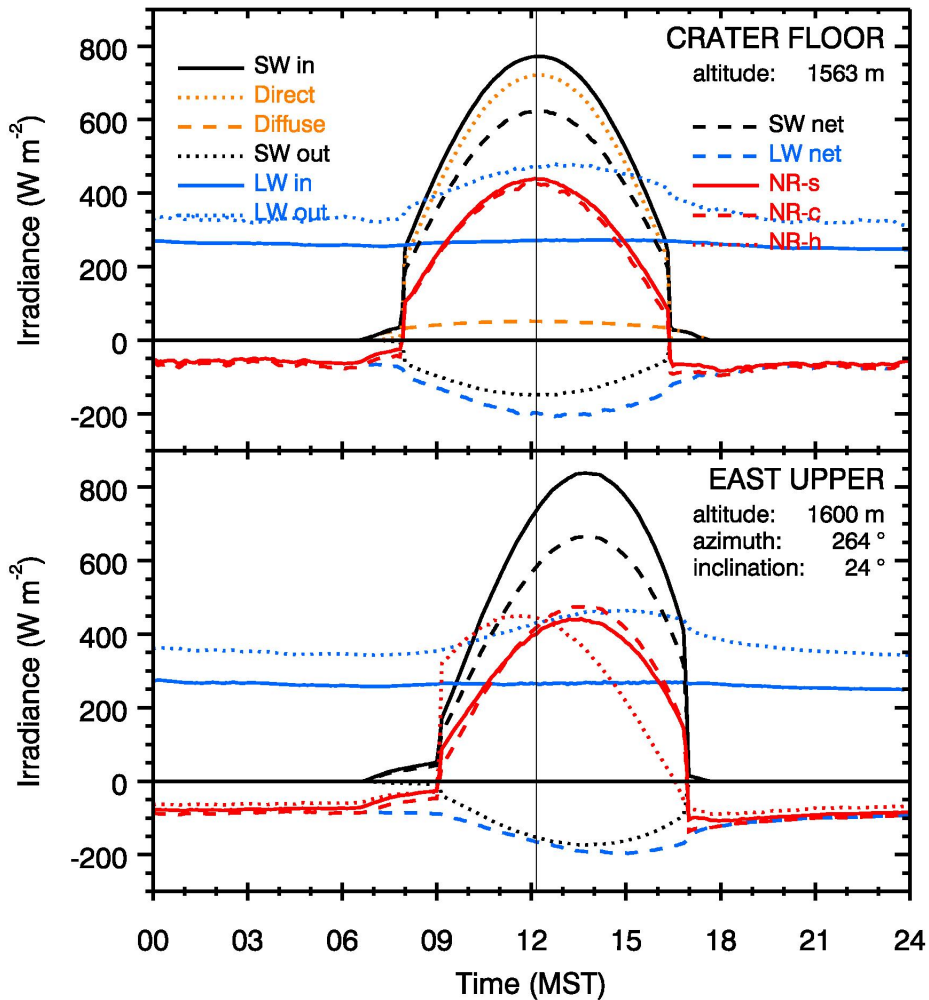


Radiation Balance Measurements during METCRAX 2006



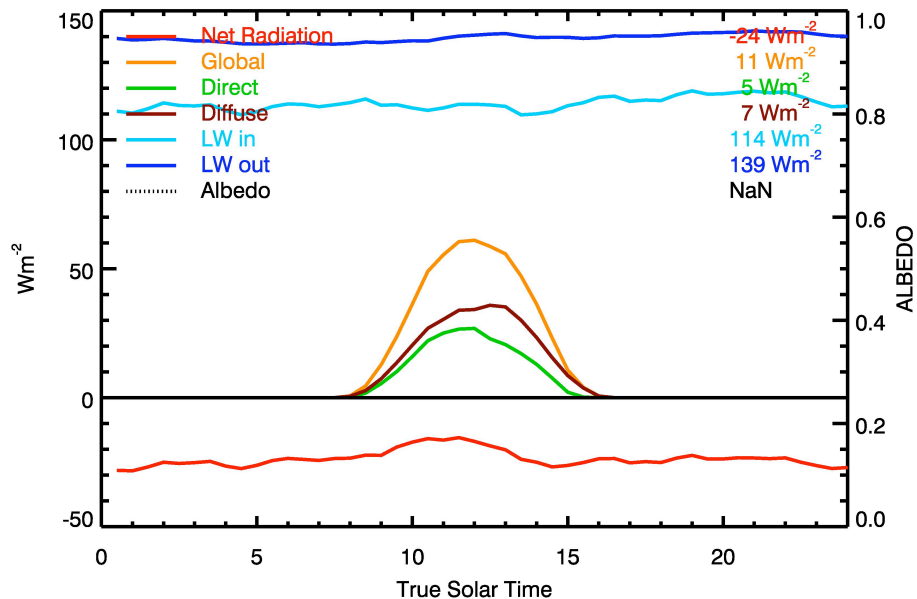
Meteor Crater, Arizona

21 October 2006

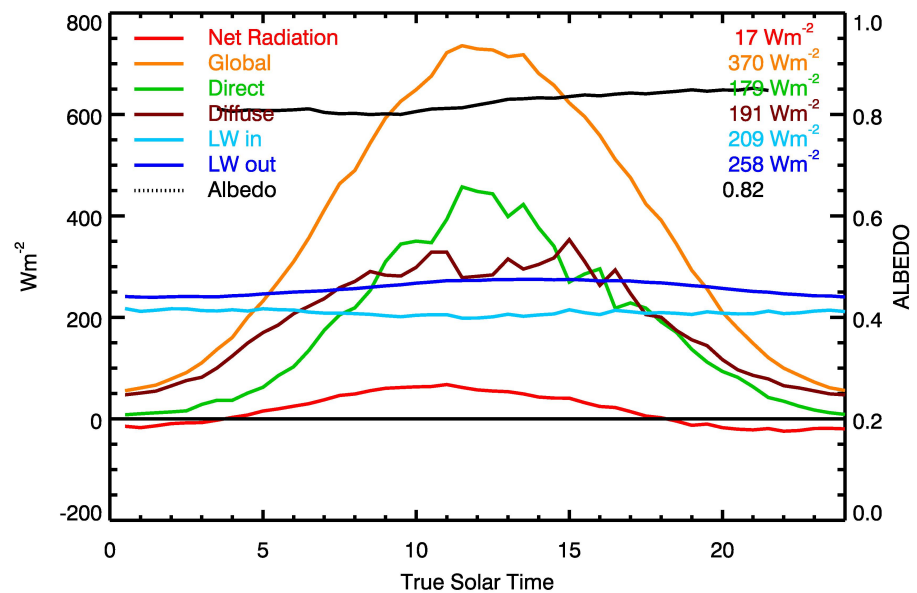


Summit, Greenland (3200 m)

FEB 2002



JUN 2002



Radiation measurements during MATERHORN

EFS-Playa



EFS-Sagebrush



EFS-Slope

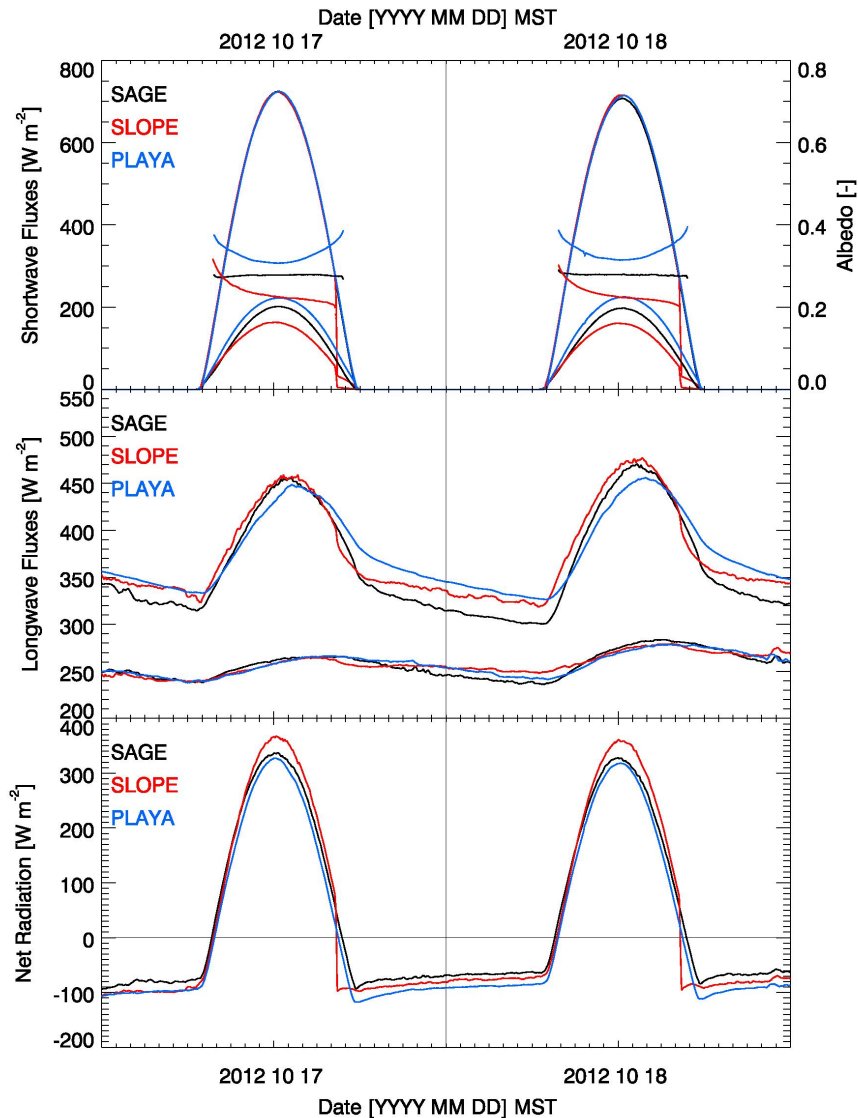


Detailed View – 4 components



Radiation Balance at EFS-Sites

$$NR = SW^{\downarrow} - SW^{\uparrow} + LW^{\downarrow} - LW^{\uparrow}$$



- Same shortwave energy input SW^{\downarrow}
- Albedo controls SW^{\uparrow}
- Same daytime NR at EFS-Sage and EFS-Playa
- Differences in SW^* are compensated by differences in LW^{\uparrow}
- NR differences (Playa – Sage) are larger at nighttime, pointing to differences in soil thermal properties.

“Wer misst, misst Mist!”

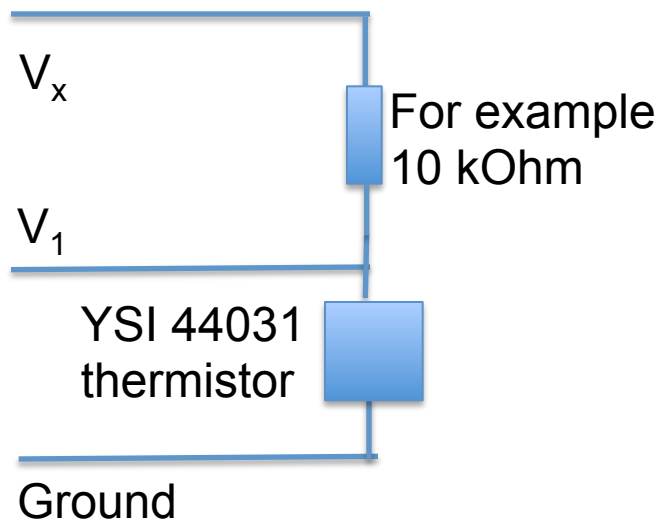


English: “Those who measure, measure m&^%\$!@# !”

Additional discussion points ...

- Build of instruments
- Thermistor measurements (YSI44031) – voltage drop across thermistor using known resistor (accuracy ?)
- Choice of voltage range / resistor / resolution / excitation voltage
- Measurements in different climates / elevations?

Considerations – LW measurements



What temperature range do I expect? – What resistance will that mean?

What voltage range do I choose?

What is my best excitation voltage?

What will my resolution ($W\ m^{-2}$) be (cc ~ $10\ \mu V / (W\ m^{-2})$)?

Temperature [°C]	Resistance [°F]	Resistance [Ohm]	Temperature [°C]	Resistance [°F]	Resistance [Ohm]	Temperature [°C]	Resistance [°F]	Resistance [Ohm]
-30	-22.0	135,200	0	32.0	29,490	30	86.0	8,194
-29	-20.2	127,900	1	33.8	28,150	31	87.8	7,880
-28	-18.4	121,100	2	35.6	26,890	32	89.6	7,579
-27	-16.6	114,600	3	37.4	25,690	33	91.4	7,291
-26	-14.8	108,600	4	39.2	24,550	34	93.2	7,016
-25	-13.0	102,900	5	41.0	23,460	35	95.0	6,752
-24	-11.2	97,490	6	42.8	22,430	36	96.8	6,500
-23	-9.4	92,430	7	44.6	21,450	37	98.6	6,258
-22	-7.6	87,660	8	46.4	20,520	38	100.4	6,026
-21	-5.8	83,160	9	48.2	19,630	39	102.2	5,805
-20	-4.0	78,910	10	50.0	18,790	40	104.0	5,592
-19	-2.2	74,910	11	51.8	17,980	41	105.8	5,389
-18	-0.4	71,130	12	53.6	17,220	42	107.6	5,193
-17	1.4	67,570	13	55.4	16,490	43	109.4	5,006
-16	3.2	64,200	14	57.2	15,790	44	111.2	4,827
-15	5.0	61,020	15	59.0	15,130	45	113.0	4,655
-14	6.8	58,010	16	60.8	14,500	46	114.8	4,489
-13	8.6	55,170	17	62.6	13,900	47	116.6	4,331
-12	10.4	52,480	18	64.4	13,330	48	118.4	4,179
-11	12.2	49,940	19	66.2	12,790	49	120.2	4,033
-10	14.0	47,540	20	68.0	12,260	50	122.0	3,893
-9	15.8	45,270	21	69.8	11,770	51	123.8	3,758
-8	17.6	43,110	22	71.6	11,290	52	125.6	3,629
-7	19.4	41,070	23	73.4	10,840	53	127.4	3,504
-6	21.2	39,140	24	75.2	10,410	54	129.2	3,385
-5	23.0	37,310	25	77.0	10,000	55	131.0	3,270
-4	24.8	35,570	26	78.8	9,605	56	132.8	3,160
-3	26.6	33,930	27	80.6	9,227	57	134.6	3,054
-2	28.4	32,370	28	82.4	8,867	58	136.4	2,952
-1	30.2	30,890	29	84.2	8,523	59	138.2	2,854

From Kipp & Zonen – CGR4 manual

Credits & Acknowledgements

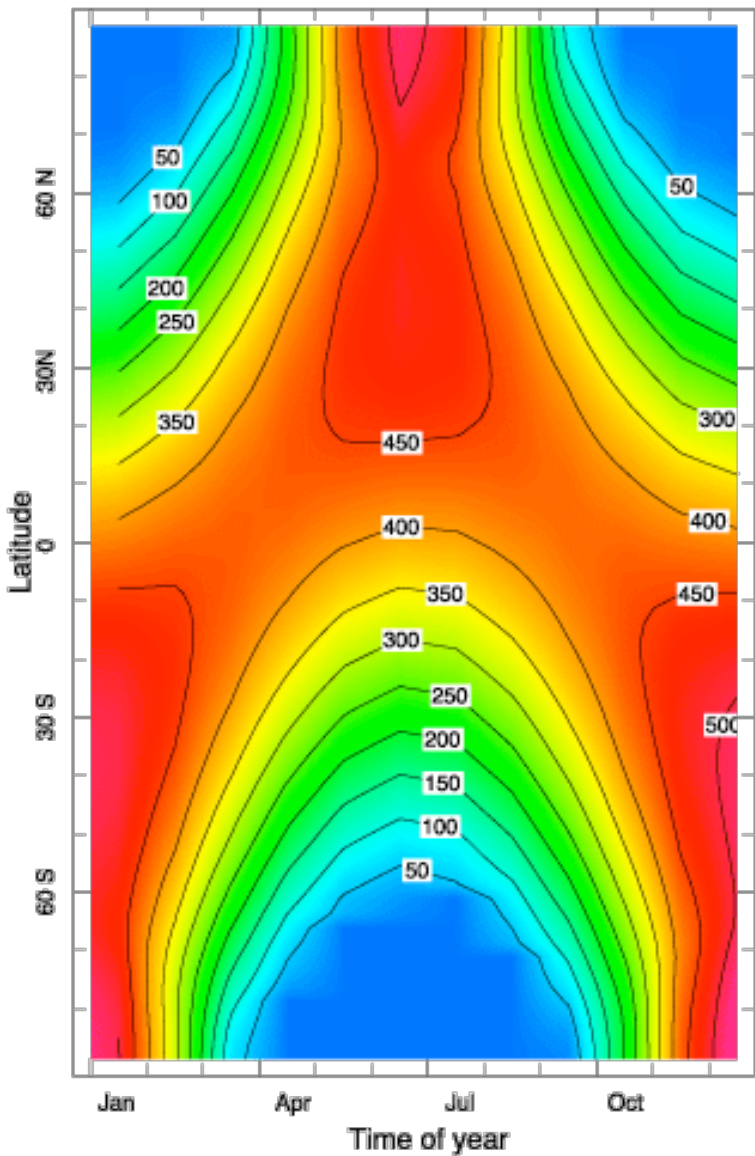
Lecture notes of Prof. Claus Froehlich, Davos: <ftp://ftp.pmodwrc.ch/pub/Claus/Vorlesung2009/>

Notes on ETH Feldkurs Rietholtzbach by Reto Stoeckli

Kipp & Zonen: <http://www.kippzonen.com/?downloadcategory/551/Pyranometers.aspx>

Wikipedia articles

Latitude-Time Distribution of Incoming Solar Radiation at the Top of the Atmosphere



Based on ERBE data. Units are W/m^2

