Observations of radiative cooling and heating under clear sky and fog conditions: Initial Results

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And updates of work by Matt Jeglum, Manuela Lehner and Dave Whiteman



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Motivation

Fog forms when the near-surface air is cooled below its dew-point temperature and when enough cloud or ice condensation nuclei are available.

Condensation conditions can be reached by different mechanisms including local cooling and mixing processes of different air masses.

MATERHORN-FOG observations were designed to directly measure the cooling contributions.

$$\frac{\partial T}{\partial t} + \vec{v_h} \nabla_h T + w \frac{\partial T}{\partial z} = \nu_T \frac{\partial^2 T}{\partial z^2} - \frac{1}{\rho c_p} \left(\frac{\partial SW}{\partial z} + \frac{\partial LW}{\partial z} + \frac{\partial H}{\partial z} \right).$$

This research investigated the **role of** *clear-air radiative cooling* or *Radiative Flux Divergence (RFD)* in the surface layer and its relative importance under different conditions.

Instrumentation and Methodology

- Small changes in LW fluxes need to be resolved.
- Kipp and Zonen CGR4 research-grade pyrgeometers at two levels





To maximize instrument accuracy, sensors are ventilated with heated air.

Careful *Relative Calibration* is necessary.

Relative Calibration - Results



Uncertainties	SLC	Heber	
LW in	± 0.38 W m ⁻²	± 0.59 W m ⁻²	
LW out	± 0.15 W m ⁻²	± 0.18 W m ⁻²	
LW net	± 0.42 W m ⁻²	± 0.61 W m ⁻²	
Radiative Heating Rate	± 5.5 K day ⁻¹	± 6.5 K day ⁻¹	

Before relative calibration: flux difference as large as 5.5 ± 2.6 Wm⁻² (PYG 8 vs 6)

Observational sites

Heber Valley (1697 m MSL) Heber Valley, Jan. 8, 2015 Salt Lake City Basin, Jan. 8, 2015

Radiative Flux Divergence (RFD) in the surface layer between 2 and ~7.5 m was investigated.

Salt Lake City Basin (1289 m MSL)



1. Clear Sky Conditions (7-9 March 2015)

Results – Radiative Fluxes – Clear Sky

SLC, clear sky, snow free

Heber, clear sky, snow free



Results – Radiative Flux Divergence (RFD) – Clear Sky

SLC, clear sky, snow free

Heber, clear sky, snow free



Results – Clear Sky

- Significant differences in the incoming, outgoing, and net longwave radiative fluxes between 2 and ~8 m AGL can be resolved after a careful relative calibration of research-grade pyrgeometers.
- A daytime convergence of the net longwave flux typically leads to a daytime radiative heating rate of up to ~100 K day⁻¹ (~4 K hr⁻¹). This magnitude corresponds to the temperature tendency observed in the morning hours. This illustrates the importance of clear-air longwave radiative heat exchange in the surface layer.
- The time lag between maximum observed heating and radiative heating indicates the important role of other processes.
- At night, under clear skies, the divergence of the outgoing flux is compensated by a convergence of the incoming flux. Zero to weak cooling results.
- Under clear-sky conditions, the divergence/convergence of the outgoing longwave flux component dominates net radiative cooling/heating.

2. Fog episodes (7-10 January 2015)

When **fog droplets** are introduced into the surface layer, radiative energy exchange changes dramatically. Radiative exchange is not limited anymore to water vapor and CO_2 spectral bands.

Results – Radiative Fluxes – with fog episodes

SLC, snow free, some fog

Heber, snow covered, some fog



Results – Radiative Flux Divergence (RFD) – with fog

SLC, snow free, some fog

Heber, snow covered, some fog



Initially, shallow fog leads to enhanced radiative cooling (Heber, 7 January). It is a strong divergence of the *downwelling* longwave flux that is responsible for the strong cooling.



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Similar behavior has been reported in the Greenlandic Arctic (Hoch et al., 2007)

Deep, but thin fog/haze in cold air pools leads to a *radiative warming* dominated by the *convergence of the incoming longwave flux* (SLC, Jan 8 ~20:00 MST, Jan 9 ~13:00 MST)



Very deep thick fog results in a *zero radiative heating/cooling*. (SLC, 9 January ~18:30 MST)



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We presented preliminary and qualitative results highlighting the complex interaction between fog and radiative flux divergence.

A rich dataset is available from MATERHORN-FOG, for the first time combining direct observations of radiative flux divergence with detailed measurements of suspended hydrometeors (fog) and aerosols.

- Radiative transfer modeling
- Heat budget calculations (sensible heat flux divergence!)

Update – Manuela Lehner

Research Assistant Prof. UU





OLUME 54

A Case Study of the Nocturnal Boundary Layer Evolution on a Slope at the Foot of a Desert Mountain

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Evening flow transition on the east slope (11/12 May 2013, IOP 4)



Evening flow transition on the east slope (11/12 May 2013, IOP 4)



Evening flow transition on the east slope (11/12 May 2013, IOP 4)



- Upper part of the slope: Weakening and stagnation of the upslope flows before the onset and increase of downslope flows
- Lower part of the slope: Gradual counter-clockwise turning and weakening of the flow

"Multiscale Forcings of Surface Winds in Complex Terrain" *Journal article to be submitted to JAMC*

Cold-Pool displacements in the Dugway basin AGU 2015









CAP Displacement Identification Criteria & Algorithm



Meteogram from PWIDS 82 (middle of the east slope)

ΔT (°C)?	Monotonic?	Recovery of 0.5∆T?	Wind from east half?	Wind from east at least 1m/s?	Overall per day count
0.5	Ν	Ν	Ν	Ν	15
2	Ν	Ν	Ν	Ν	7.2
3	Ν	Ν	Ν	Ν	4.6
3	Υ	Ν	Ν	Ν	3.4
3	Υ	Υ	Ν	Ν	2
3	Y	Y	Y	Ν	1.6
3	Y	Υ	Y	Υ	0.9
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Some numbers from PWIDS 82...

YELLOW ROW

- 63 total displacement events were observed.
- Average of 1.6 displacement events per night
- One displacement event on 29 of 40 nights (Fall campaign)
- Maximum: 5 per night.

Standardized Number Count of Displacement Events– Fall Campaign





- Flow interactions with Granite Peak appear to influence the advection processes that lead to cold pool displacements in the Dugway Basin.
- Cold Pool Displacements are relatively common on days that would be called "transitional", i.e. when synoptic forcing is relatively strong but variable in time and strength.
- Advection of cold air eastward and southward, advection of warm air northward and westward or turbulent mixing in the vertical all appear to be important during displacements.
- Composite data of cold air displacements point to east-west displacements early in the night, and north-south displacements later in the night.
- No cold air displacements are observed when winds are strong all night (> 10 m/s) or under significant cloud cover with precipitation. (No cold pools form.)

Work is ongoing – focusing on processes forcing the displacements. An update will be given by Matt at AGU.

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