Cloud-topped Boundary Layers



- Shallow cumulus (Cu) boundary layers
- Stratocumulus (Sc)-capped BLs
- Cold-air outbreaks
- Shear-driven shallow stratus layers
- Summertime Arctic stratus

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CONSTRAIN

As air advects over warmer seas the Sc transitions to mixedphase cumulus clouds at around 60N, prior to reaching land

MODIS ch4





No-Ice simulation with UKMO LEM



* LEM is UKMO large eddy simulation model



Ice simulation with UKMO LEM



Liquid/ice potential temperature (left) and total water (right) from UM LAM, LEM, SCM







Liquid water potential temperature & total water content, Nc = 10 cm⁻³



Coarser resolution hardly has an effect on the mean state evolution



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Radiation Fog in the Central Valley of California MODIS True Color Image 2115 UTC 17 Jan 2003





Dynamically Forced Fog

Dynamically forced fog is generated primarily by the cooling of moist near-surface air by dynamic processes. Dynamic processes may include advection and vertical mixing processes that lead to changes in the boundary layer temperature or moisture characteristics.

Perhaps most familiar is advection fog, which is a ground-based cloud caused by the cooling of an air mass to the saturation point as it moves over a colder surface. That surface may be cold ground, snow cover, water, or ice. Advective fog events in the marine environment tend to occur most often where a warm air mass moves over much cooler ocean water.

Dynamically Forced Fog

Regions where this is common include:

- Along the west coast of the North and South America and more infrequently along the west coast of Africa during the cool season
- Over the northeast and northwest Atlantic north of the Gulf Stream and over the northwest Pacific north of the Kuroshio Current during the warm season



GOES-8 SST Image Synoptic Pattern Conducive to Fog/Stratus



NOAA/CoastWatch

Daytime Frequency of Sky-Obscuring Fog Over the Ocean, Jun-Jul_Aug, Based On Ship Observations, 1954-1992



Daytime Frequency of Fair-Weather Stratus Over the Ocean, Based On Ship Observations, 1954-1992



Air Parcel Trajectories for Advection Fog Events in NE Atlantic



©The COMET Program

The skew-T sequence shown here is a conceptual example of how the boundary layer of an air parcel evolves over time as it moves over a cooler sea surface environment. A hypothetical air parcel trajectory is shown on the map by the white line. The numbers indicate the location of the air parcel for each highlighted stage.











Main Regions of Upwelling in Coastal Regions Around the World



- Shallow cumulus (Cu) boundary layers
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Figure 2.2: A simplified diagram of different modes of interaction between low clouds and ocean/ice surfaces in the Arctic.

(de Boer 2009)

Arctic Leads



Lead and associated plume. Photo taken on BASE flight 16, October 12, 1994, over the Beaufort Sea.

Outflow from deep convection



Thunderstorm Outflow Boundary

2231 UTC 02 Jun 95